

What is the Mississippian Play Without the Arbuckle and What We Need to Know About It

**BY
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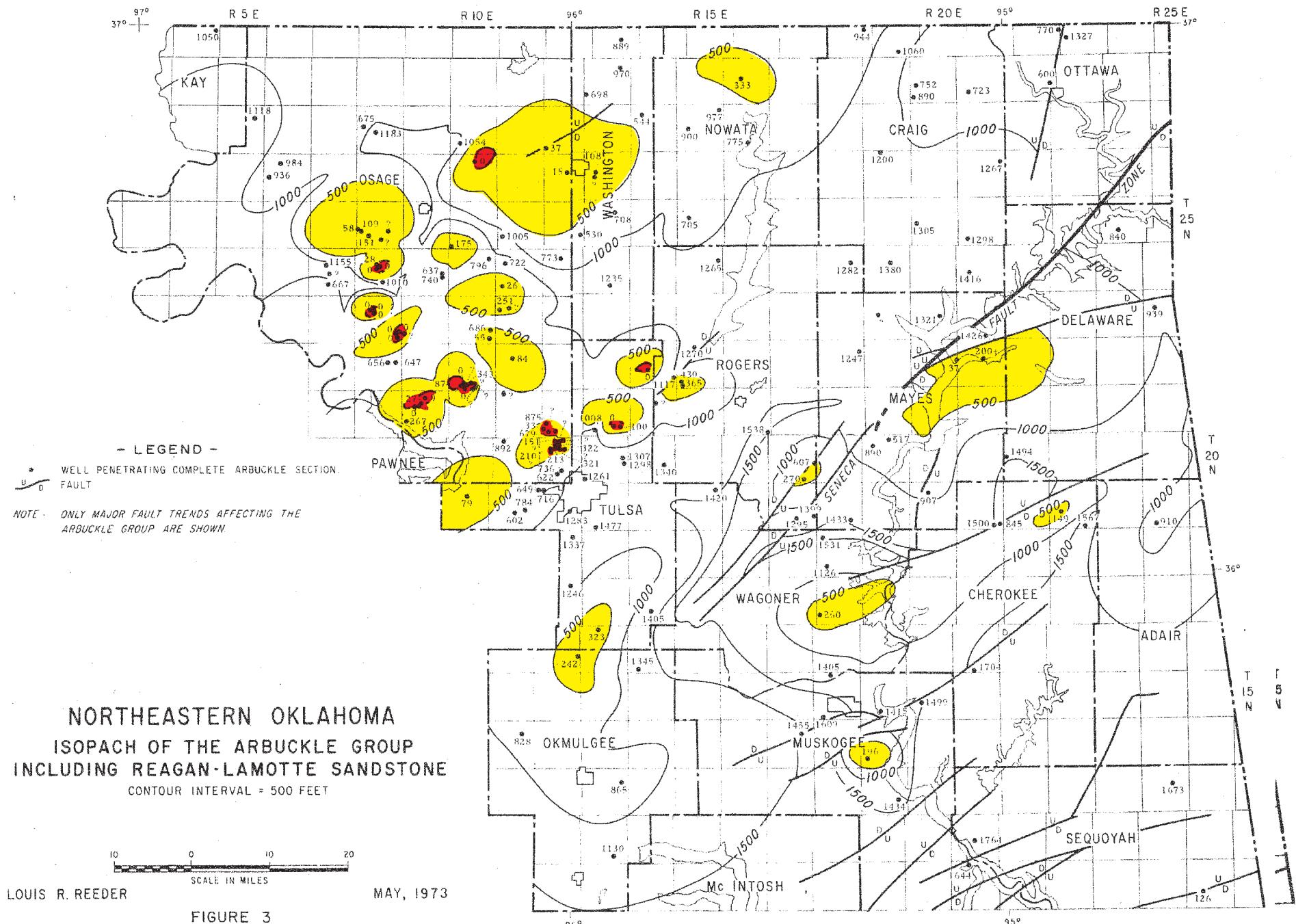


Fig. 2

Louis R Reeder, 1973
Fig. 3, Reprinted by
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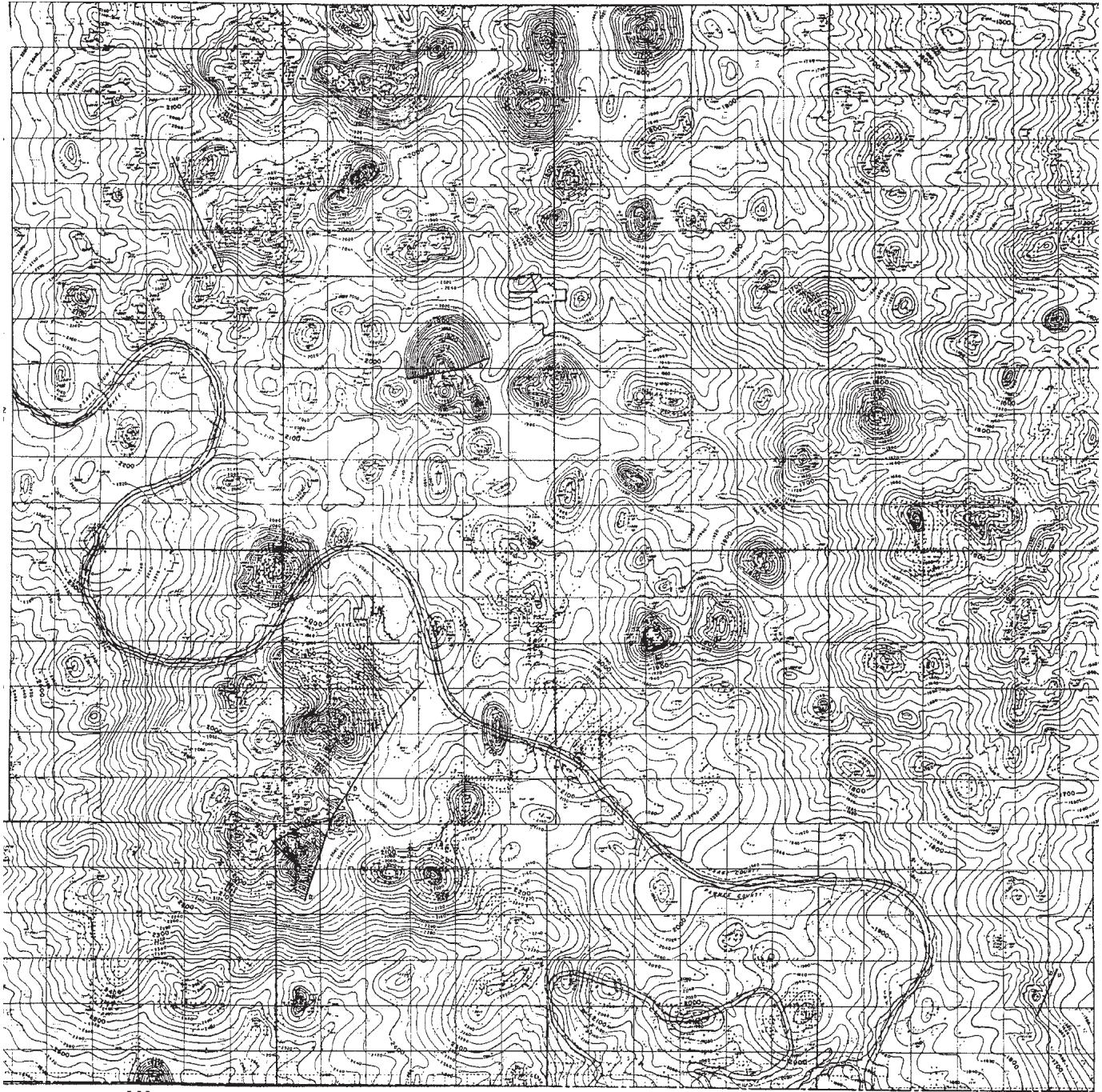


Fig 3

Arbuckle Structure, Osage and
Pawnee Counties, Oklahoma

John H Rountree, 1991
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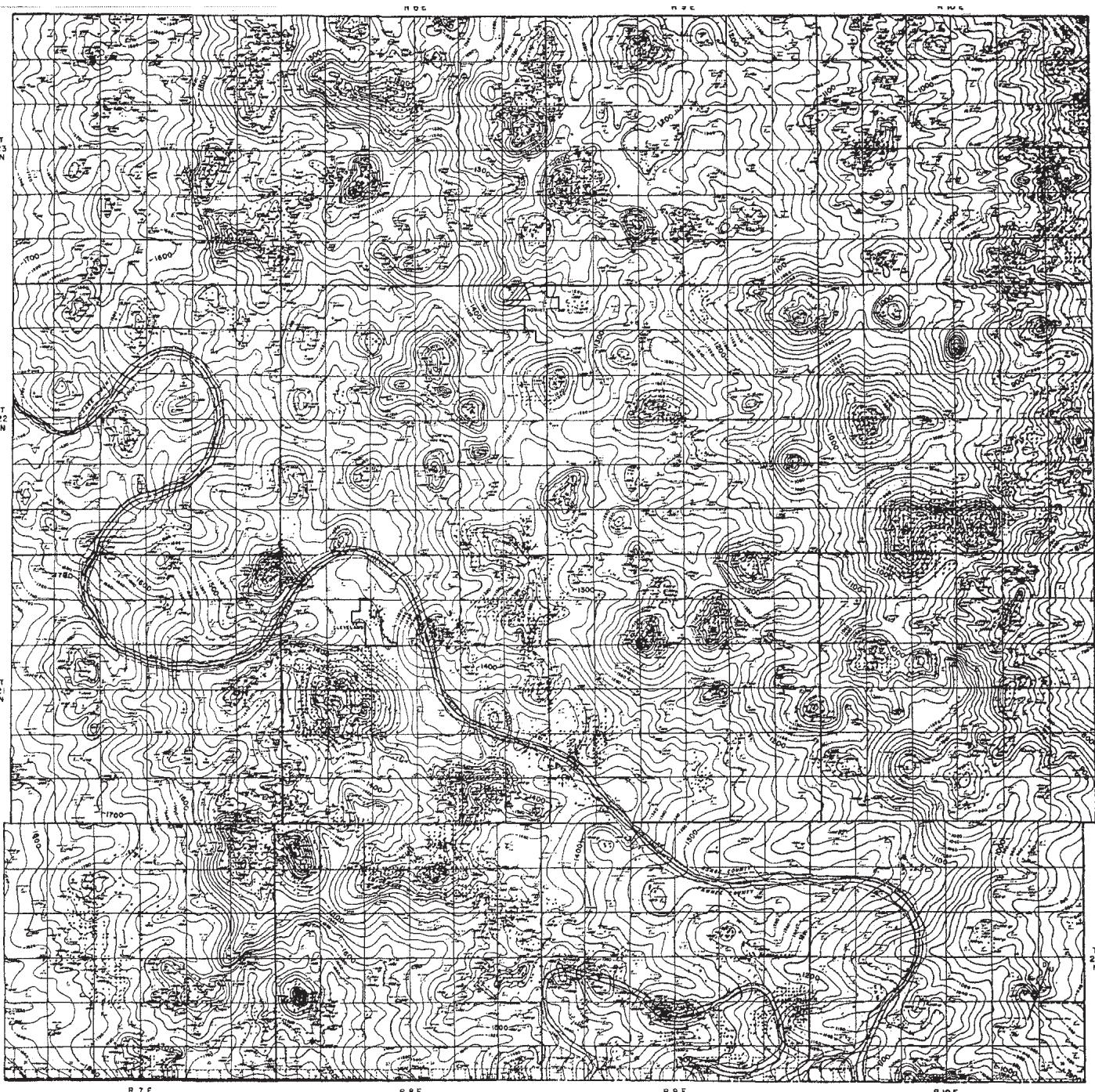


Fig. 4

Pink Lime Structure, Osage and
Pawnee Counties, Oklahoma

John H Rountree, 1991
Plate 1, Reprinted by
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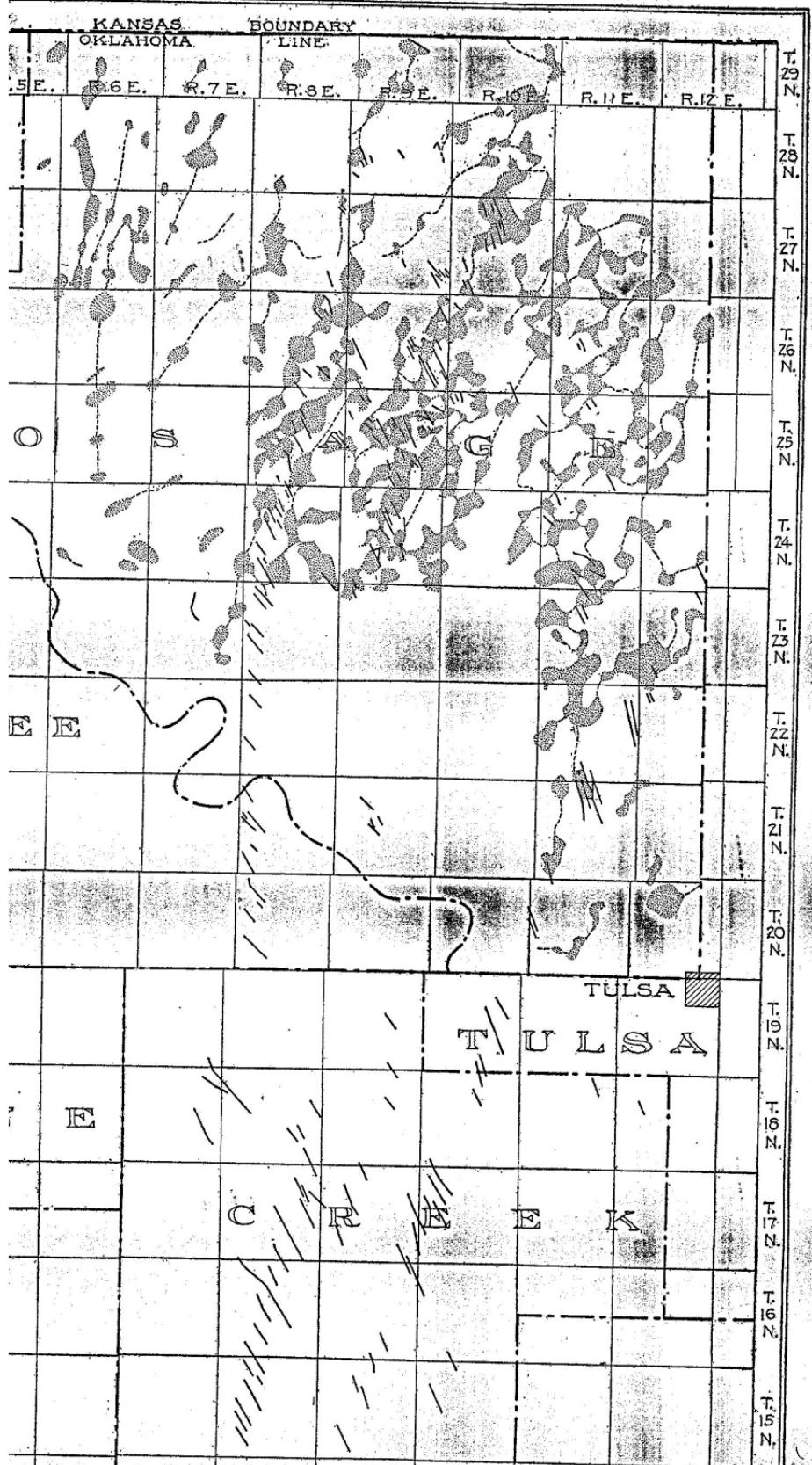


Fig 4 a En Echelon faults and northerly alinement
of anticlines and domes in Creek and
Osage Counties, Oklahoma

Fath, A. E., 1920,
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Geological Survey

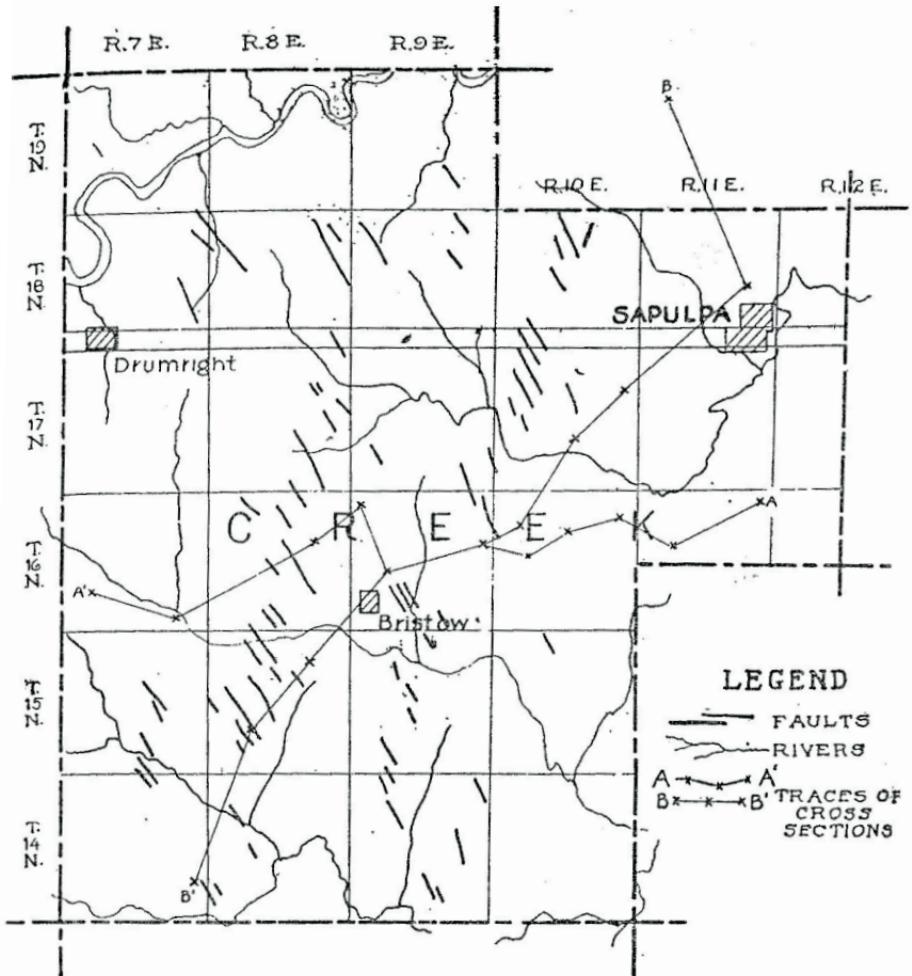
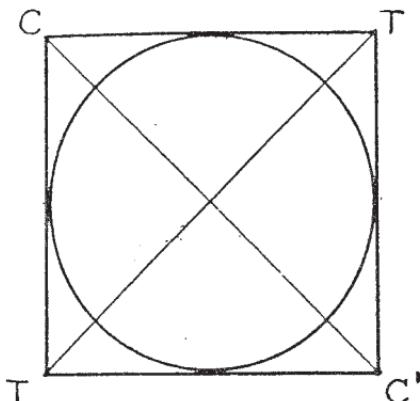


Fig 4 b

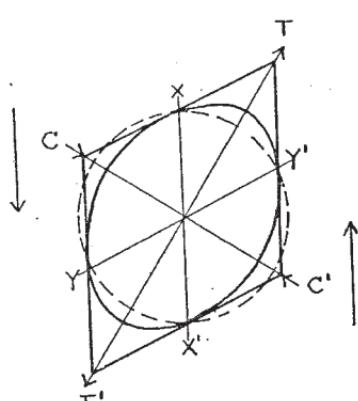
Zones of parallel faults
in Creek County, Oklahoma

Merritt and McDonald, 1926
Fig 2, Reprinted by
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Geological Society



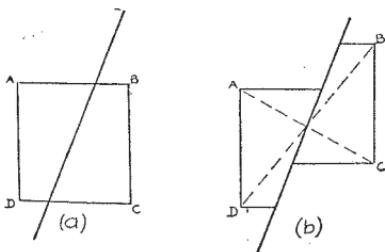
A

Illustrating the strain ellipse as developed by Leith. It is assumed that the square **CTC'T'** is a section of incompetent rocks, with the circle (representing the undeformed position of the strain ellipse), and the diagonal lines drawn on it for illustrative purposes.



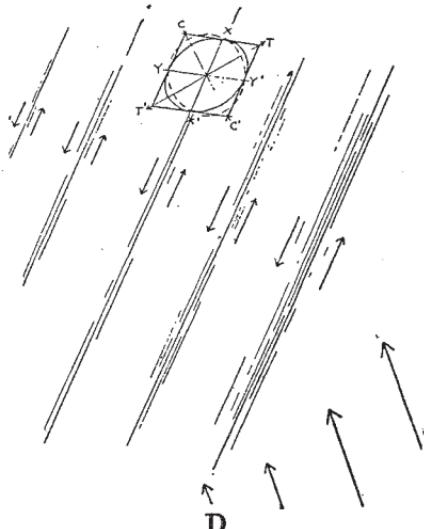
B

If sketch **A** is deformed by forces acting as indicated by the arrows the circle becomes an ellipse, which means compression between **C** and **C'** and tension between **T** and **T'**. **XX'** and **YY'** are planes of no distortion or planes of maximum shear. If **CTC'T'** be a section of incompetent sediments overlying shearing planes in the crystalline rocks, it is clear that NE-SW compression folds would tend to develop between **CC'** and tension faults striking NW-SE would develop between **TT'**.



C

Sketches **(a)** and **(b)** picture the effect of shearing in crystalline rocks. **(a)** may be considered as an undeformed section of the rocks and **(b)** the same section faulted. Notice that the distance between **AC** in **(b)** is shortened, implying compression, and that **BD** is lengthened, implying tension. There is more lengthening than shortening. In northeastern Oklahoma there are more tension than compression faults.



D

Hypothetical sketch showing the probable lines of weakness and shearing in the basement crystalline rocks. The large arrows to the SE indicate the direction and intensity of the parent thrust coming from the Ouachita-Arbuckle orographic element. The small arrows along the shear lines show the relative movement of the shearing and its intensity.

Figure 6. Illustrating the principles of the strain ellipse and its application to the interpretation of the structural features and showing the effect of shearing in the basement rocks, the origin of the parent forces, and the probable lines of weakness in the pre-Cambrian basement crystalline rocks.

Fig 4 c

Merritt and McDonald, 1926

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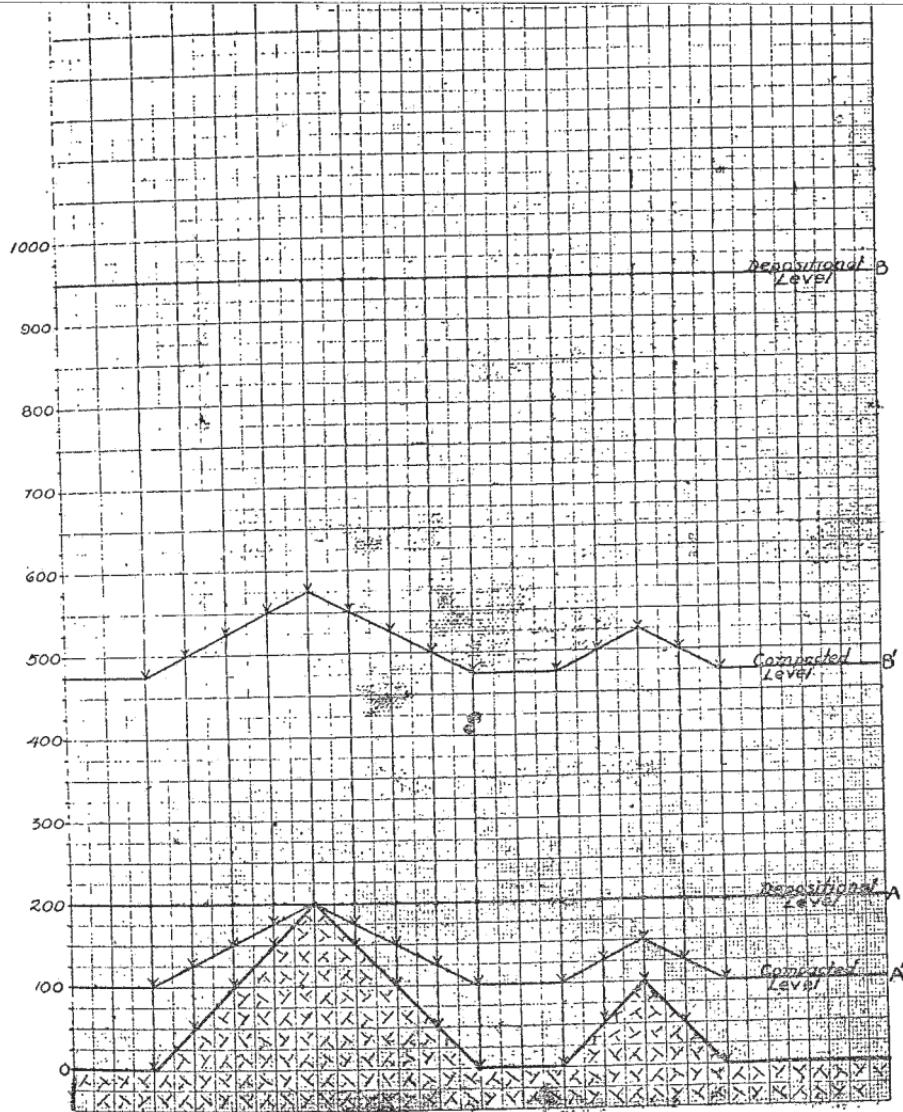


Fig 4 d

**Scenario of Compaction
after deposition**

Merritt and McDonald, 1926
Fig 4, Reprinted by
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Geological Society

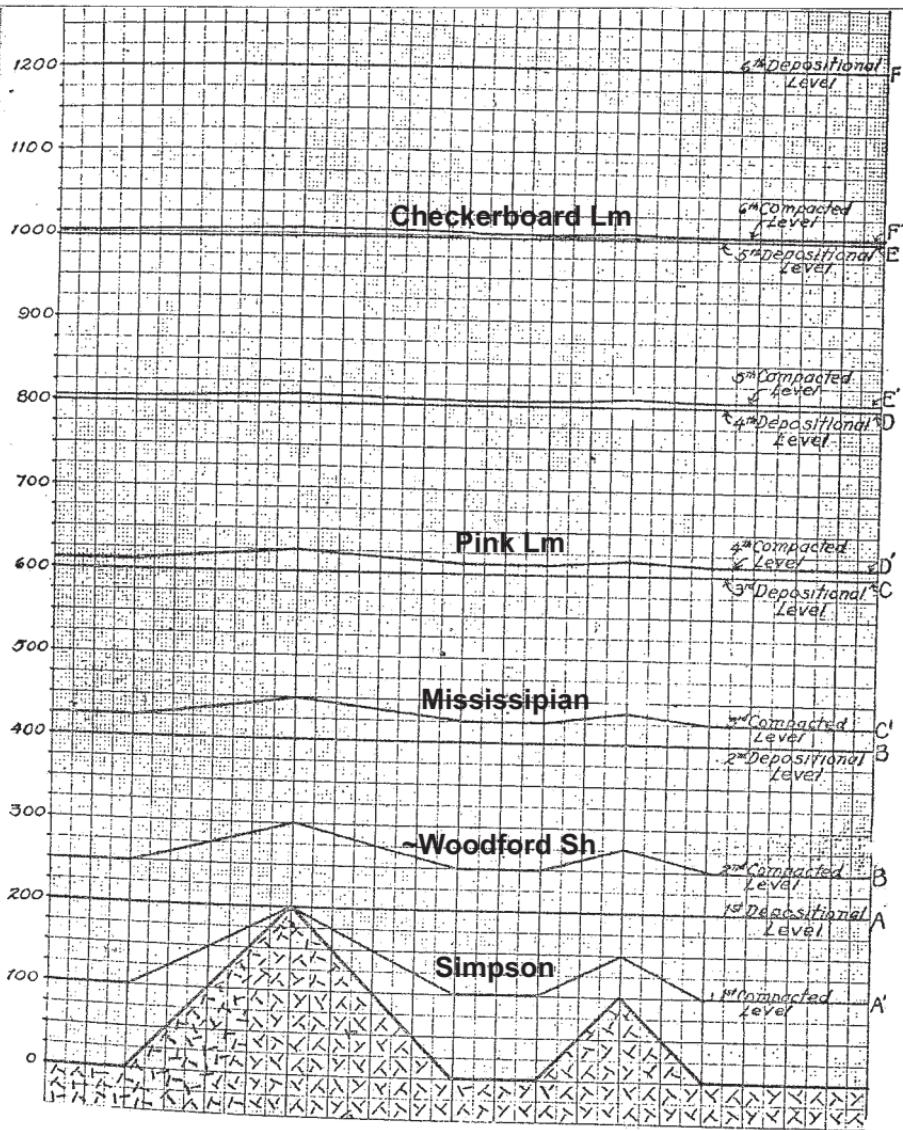


Fig 4 e

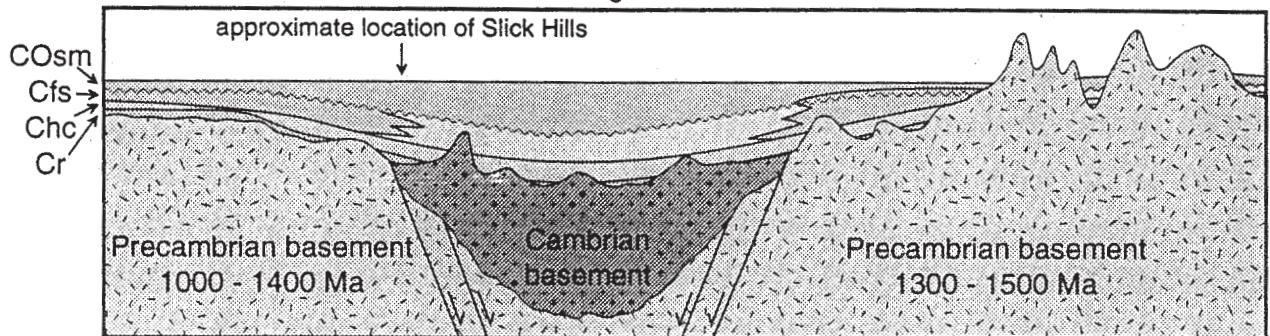
**Scenario of compaction
during deposition**

Merritt and McDonald, 1926
Fig 2, Reprinted by
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SW

Southern Oklahoma
Aulacogen

Tulsa Mountains NE

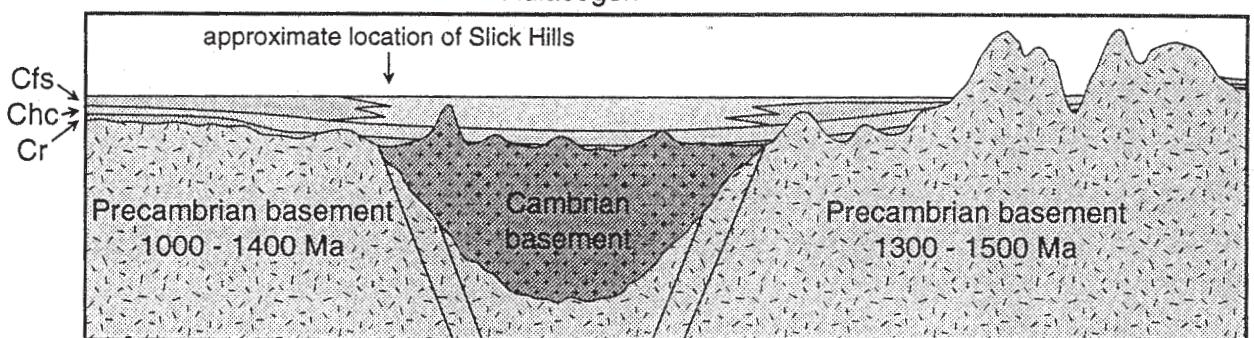


STAGE III - Reactivation of basin margin faults and deposition of Cambro-Ordovician
Signal Mountain Formation in deeper waters? (Trempealeauan - Canadian)

SW

Southern Oklahoma
Aulacogen

Tulsa Mountains NE



STAGE II - Deposition of Cambrian Timbered Hills Group and Cambrian Fort Sill
Formation. (Franconian)

SW

Southern Oklahoma
Aulacogen

Tulsa Mountains NE

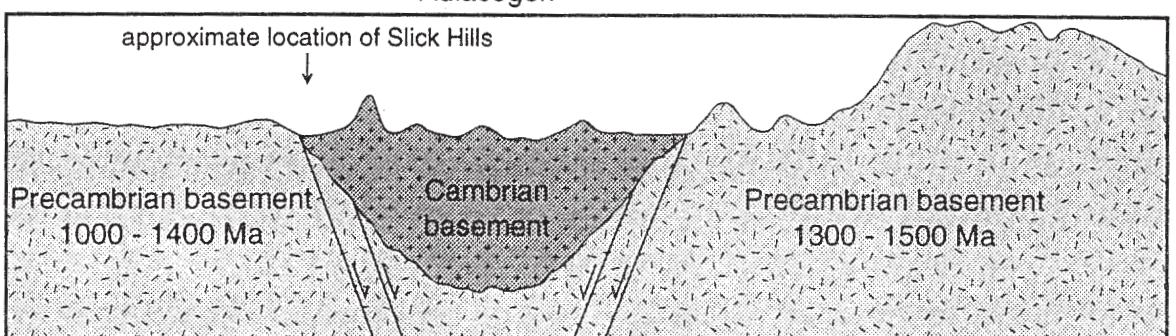


Fig 5

Pre-Cambrian
Unconformity surface

Hosey and Donovan, 2000
Fig 13, Reprinted by
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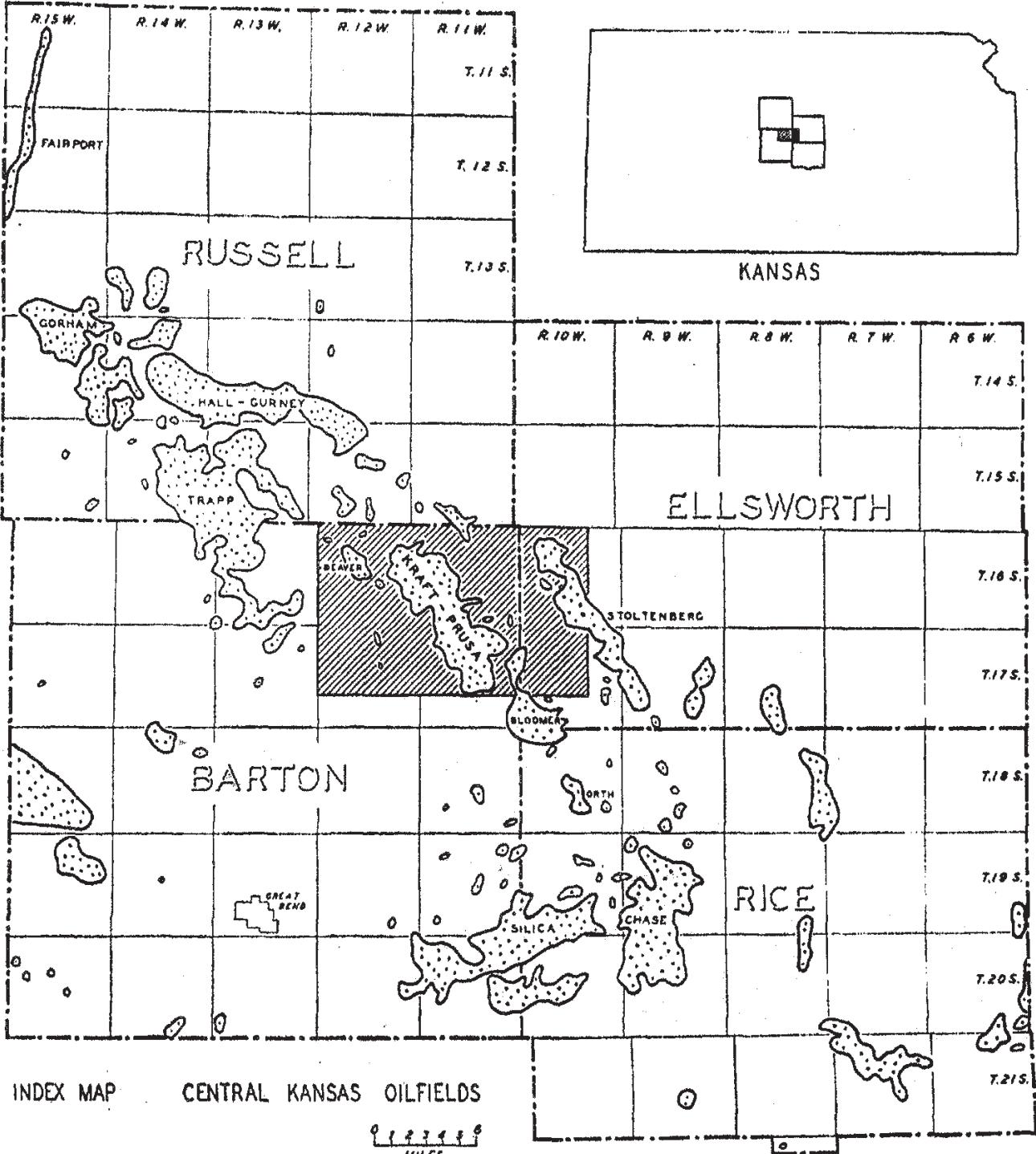


Fig. 6

Exposed Pre-Cambrian “hills”
in central Kansas

Robert Walters, 1946
Plate 2, Reprinted by
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Association of Petroleum
Geologists

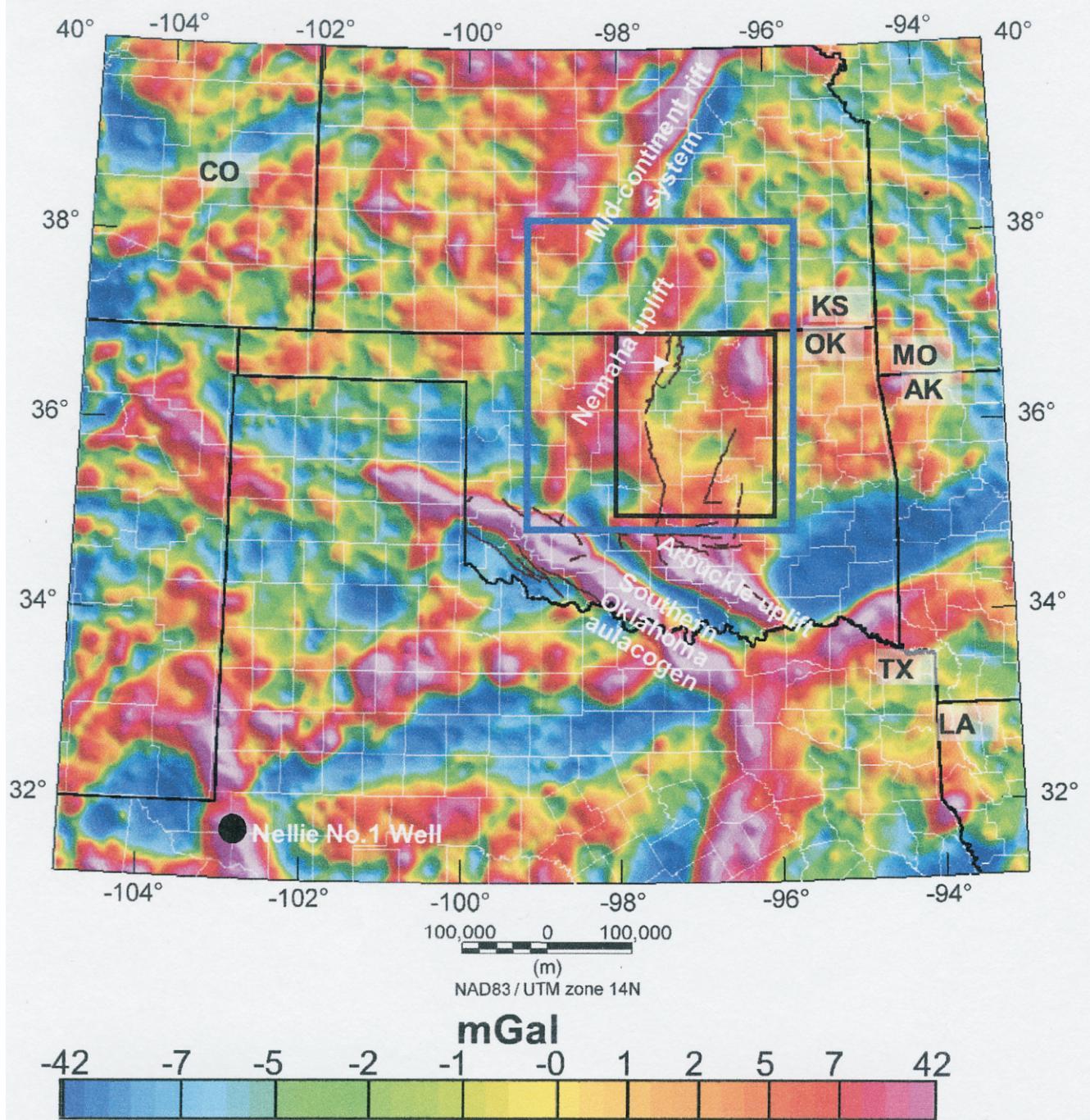


Fig 7

Simple residual Bouguer anomaly map of the mid-continent region

Elebiju et al, 2011
Plate 2, Reprinted by
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Association of Petroleum
Geologists

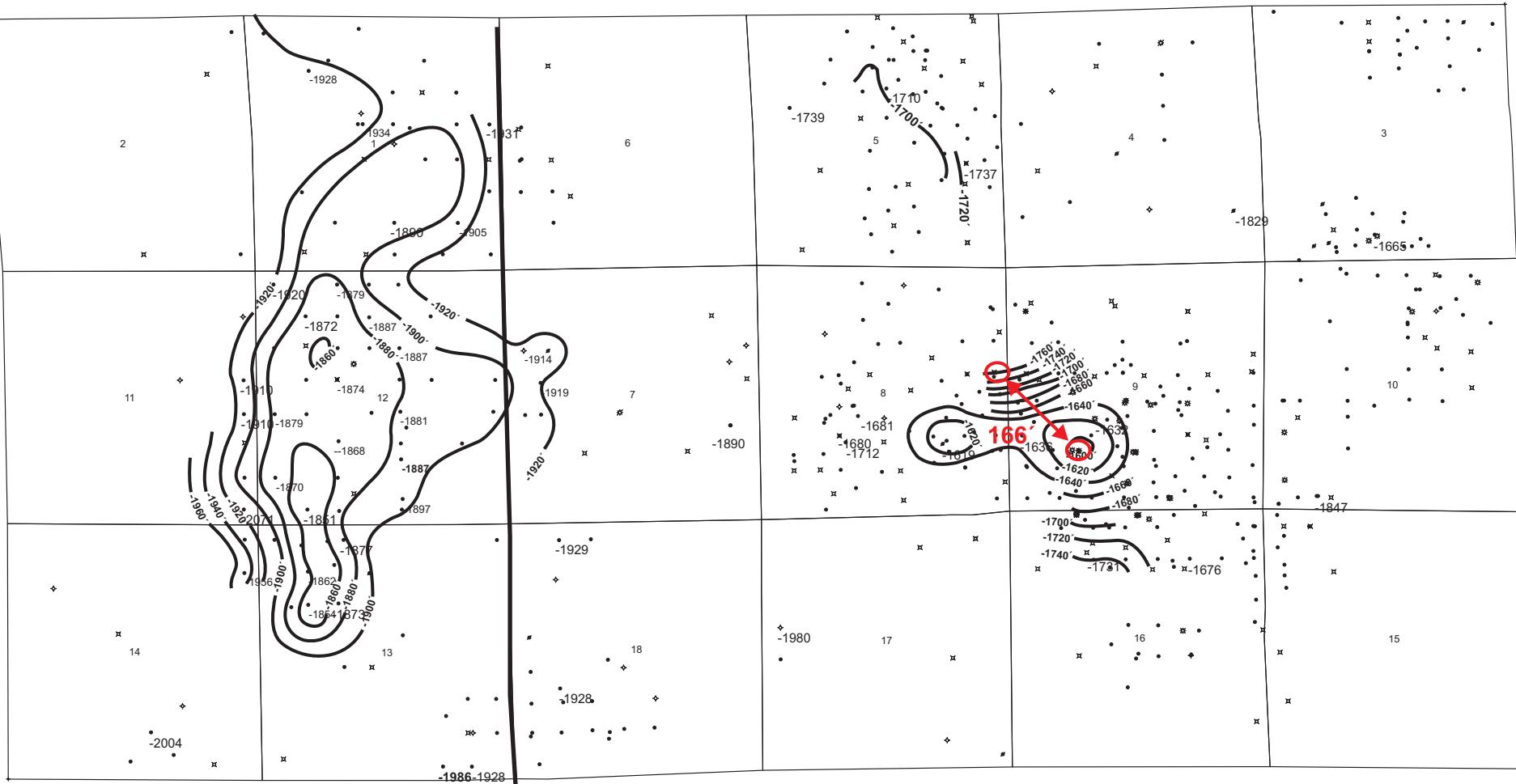


Fig. 9a

Arbuckle Structure

Barker Field Study Osage County, OK

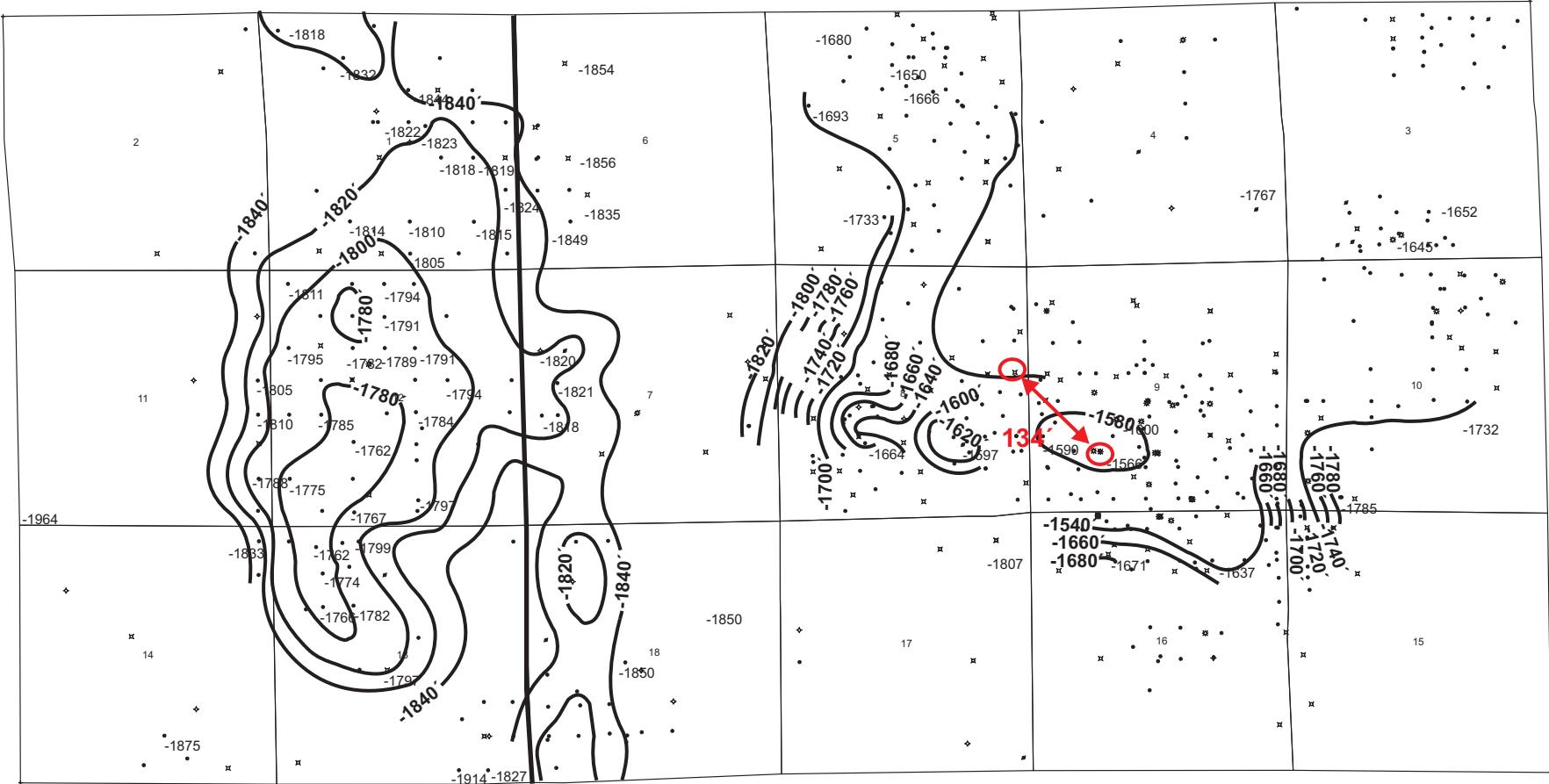


Fig. 9b

Simpson Structure

Barker Field Study Osage County, OK

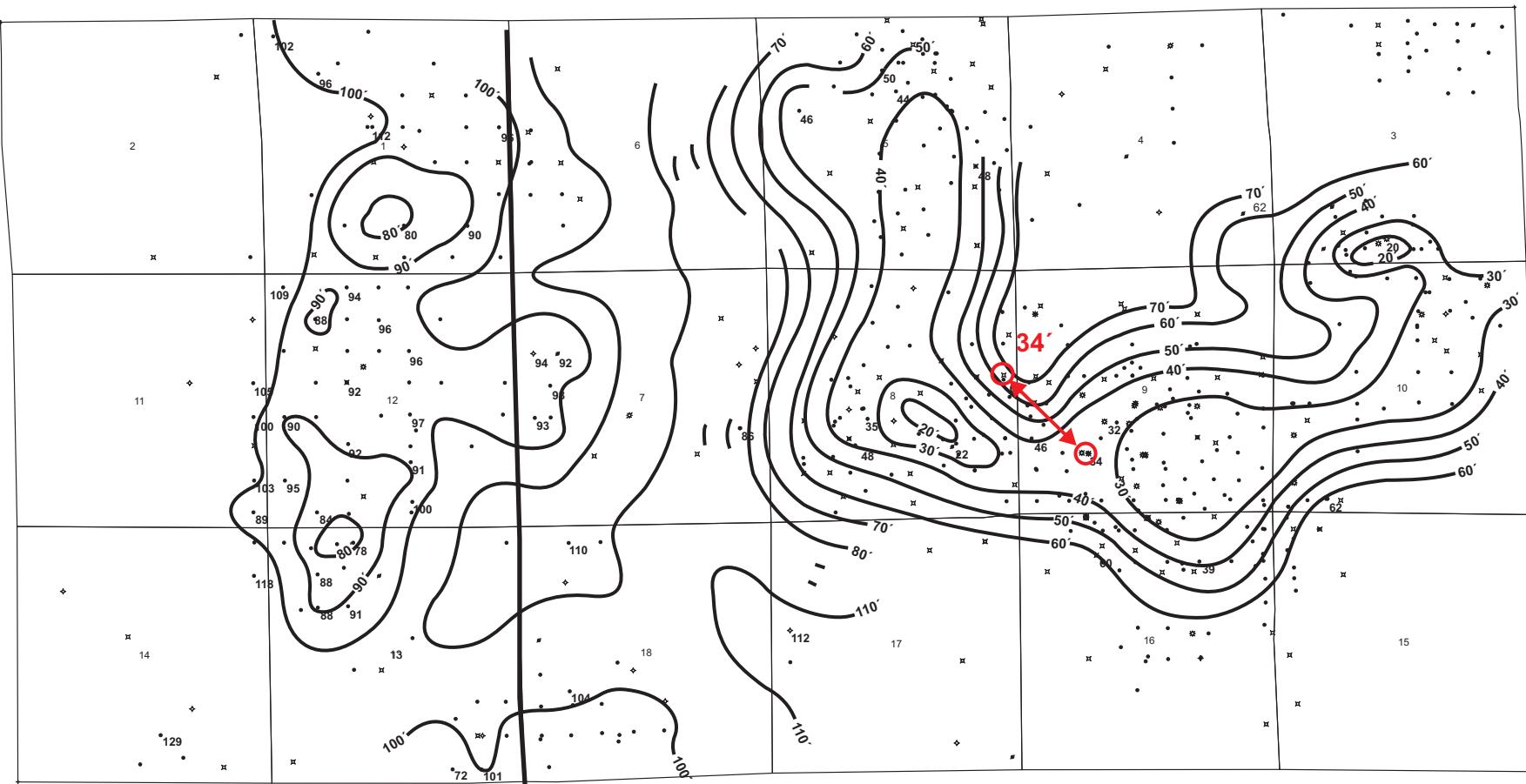
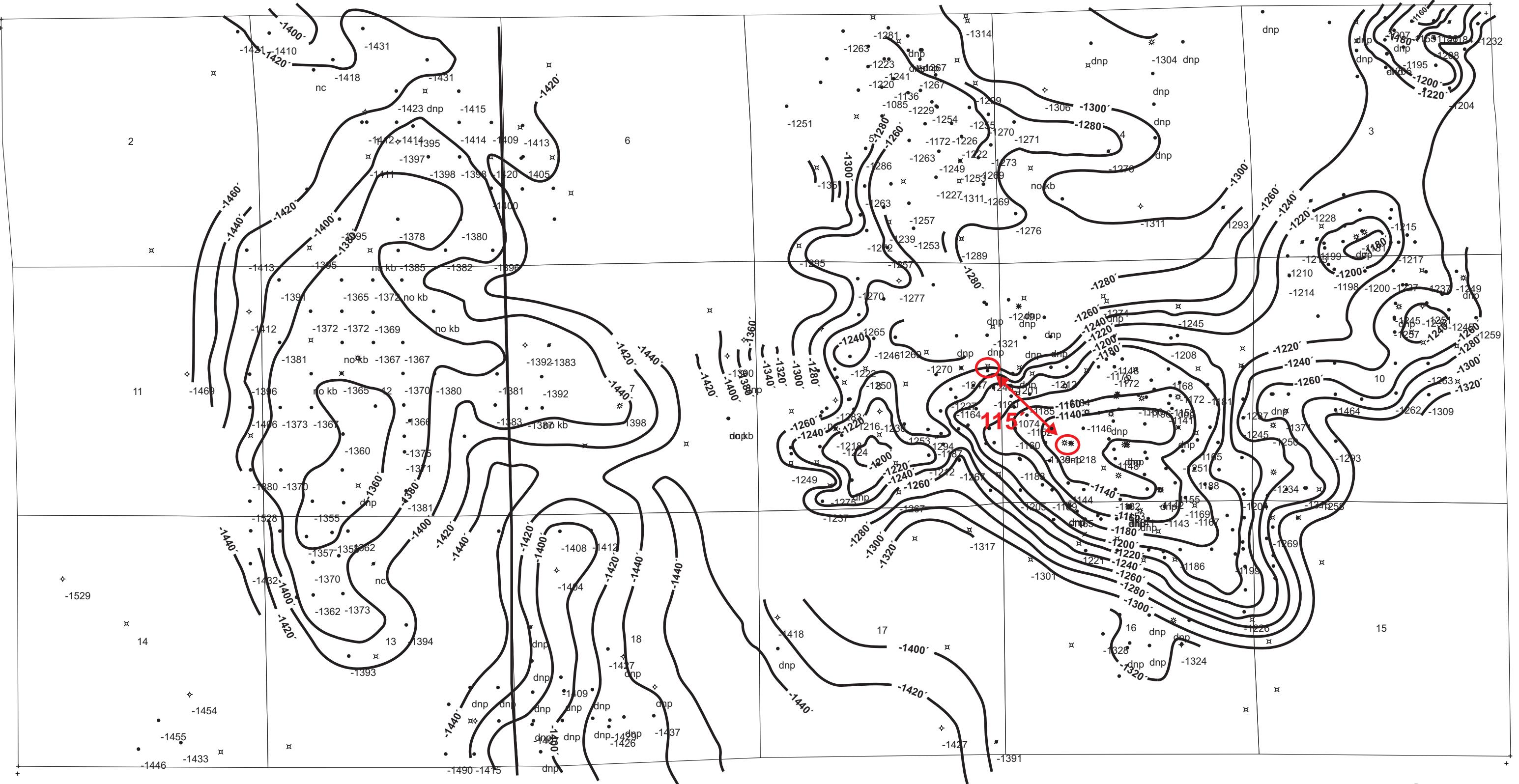


Fig. 9c

Simpson Isopach

Baker Field Study
Osage County, OK



Pink Structure

Fig 9d

Barker Field Study
Osage County, OK

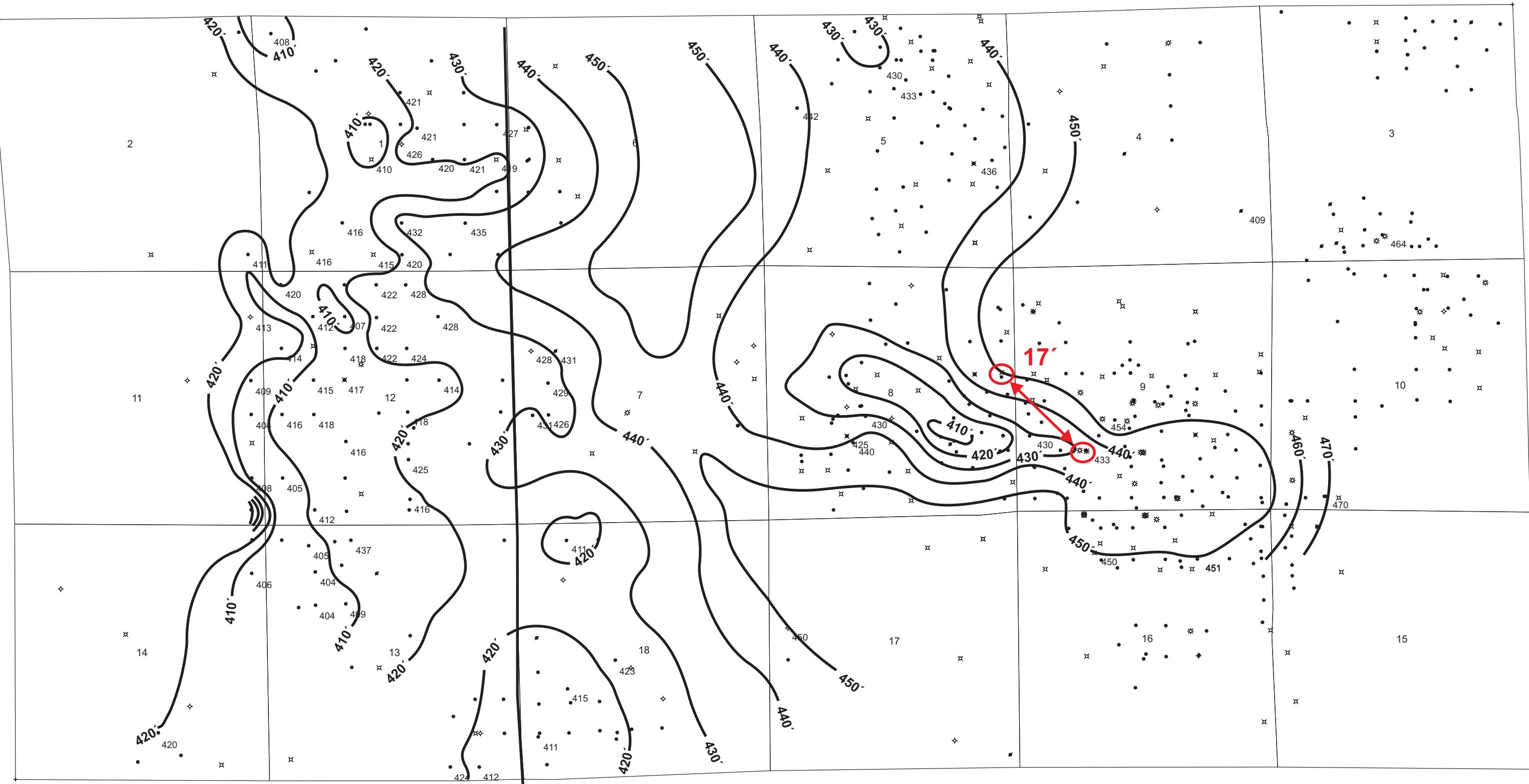


Fig. 9e

Pink-Simpson Isopach

Barker Field Study
Osage County, OK

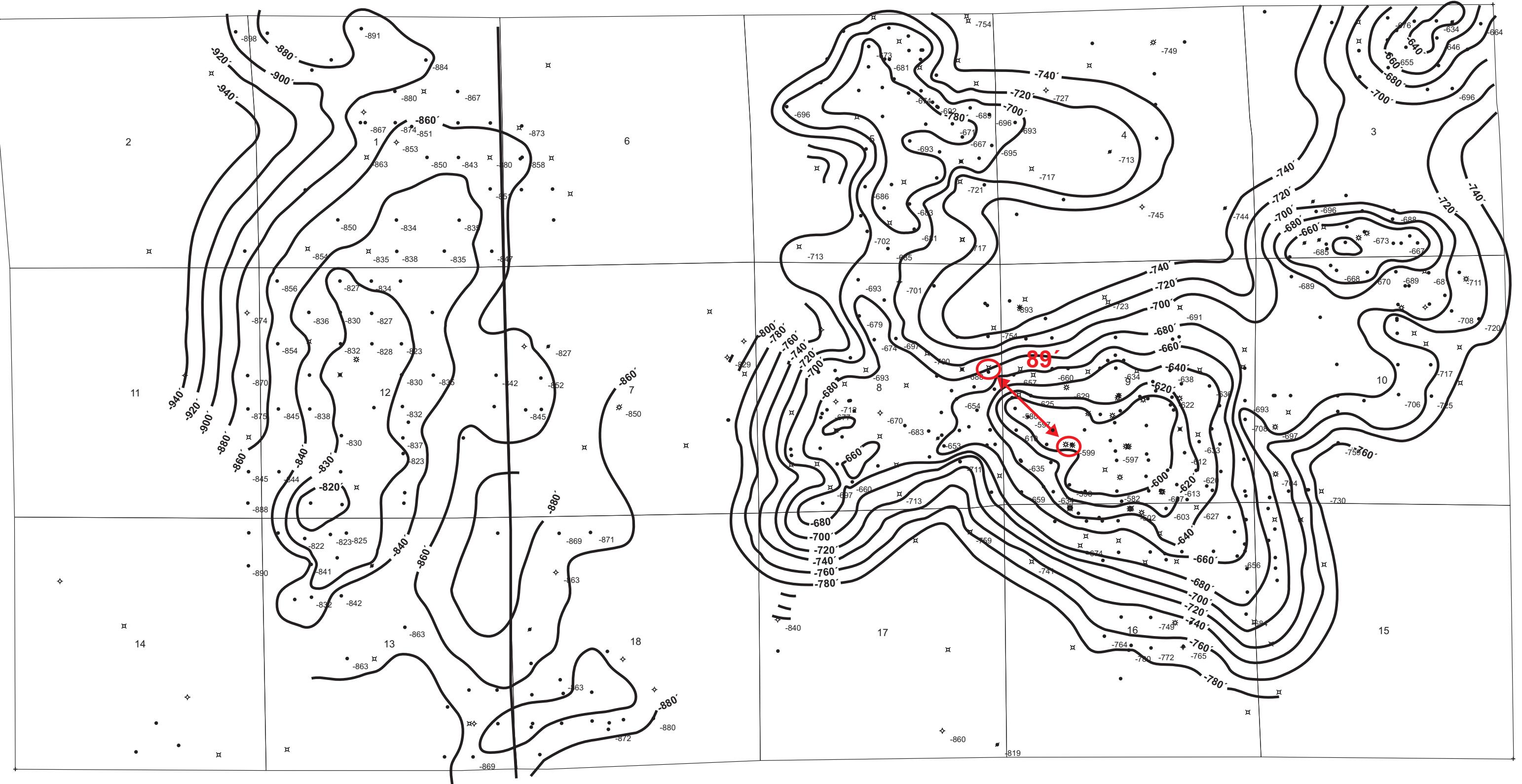


fig. 9f

Checkerboard Structure

Barker Field Study
Checkerboard Structure

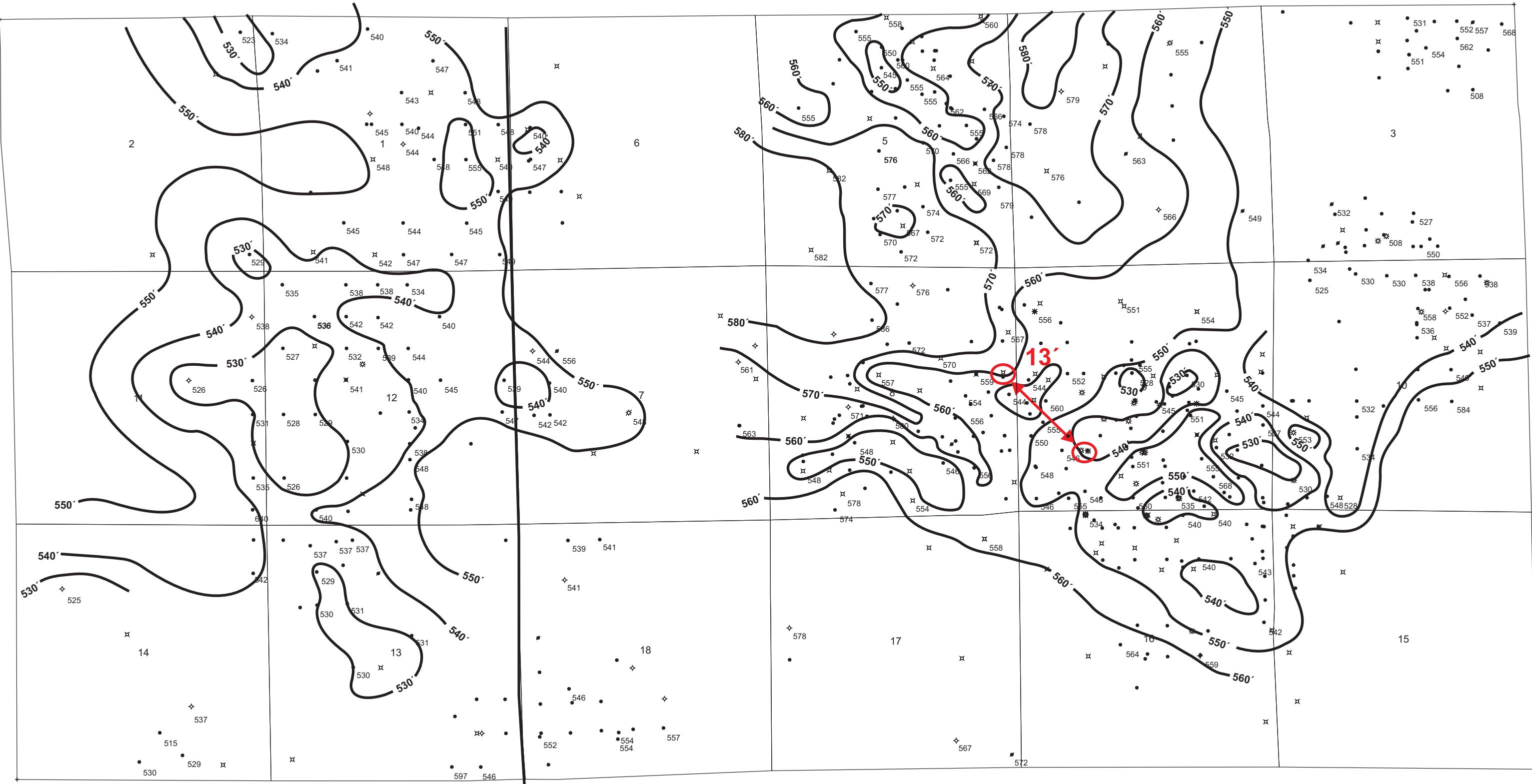


Fig 9g

Checkerboard-Pink Isopach

Barker Field Study
Osage County, OK

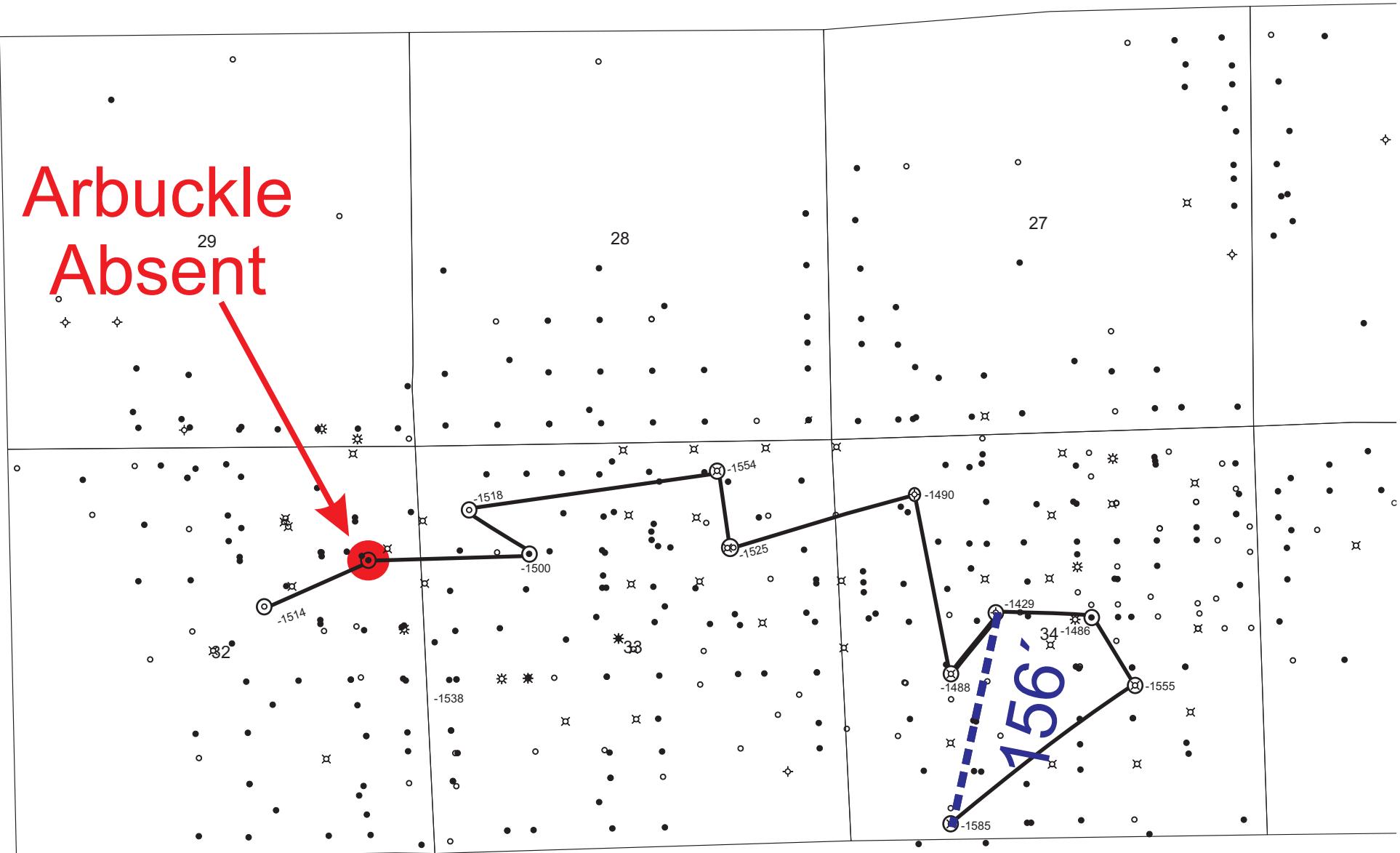


Fig. 10

Arbuckle Structure

Wildhorse Field
Osage County, OK

**Arbuckle
Absent**

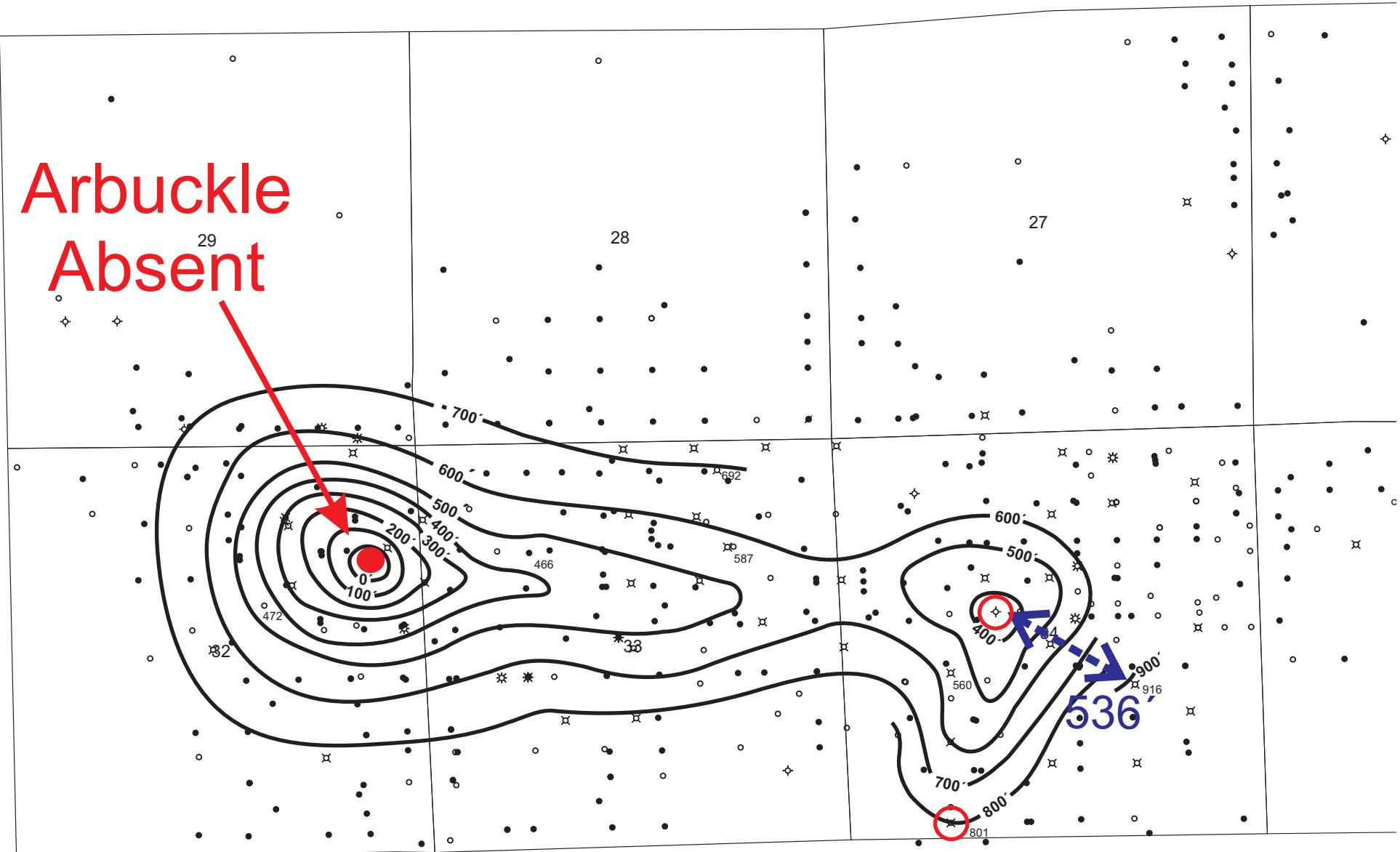
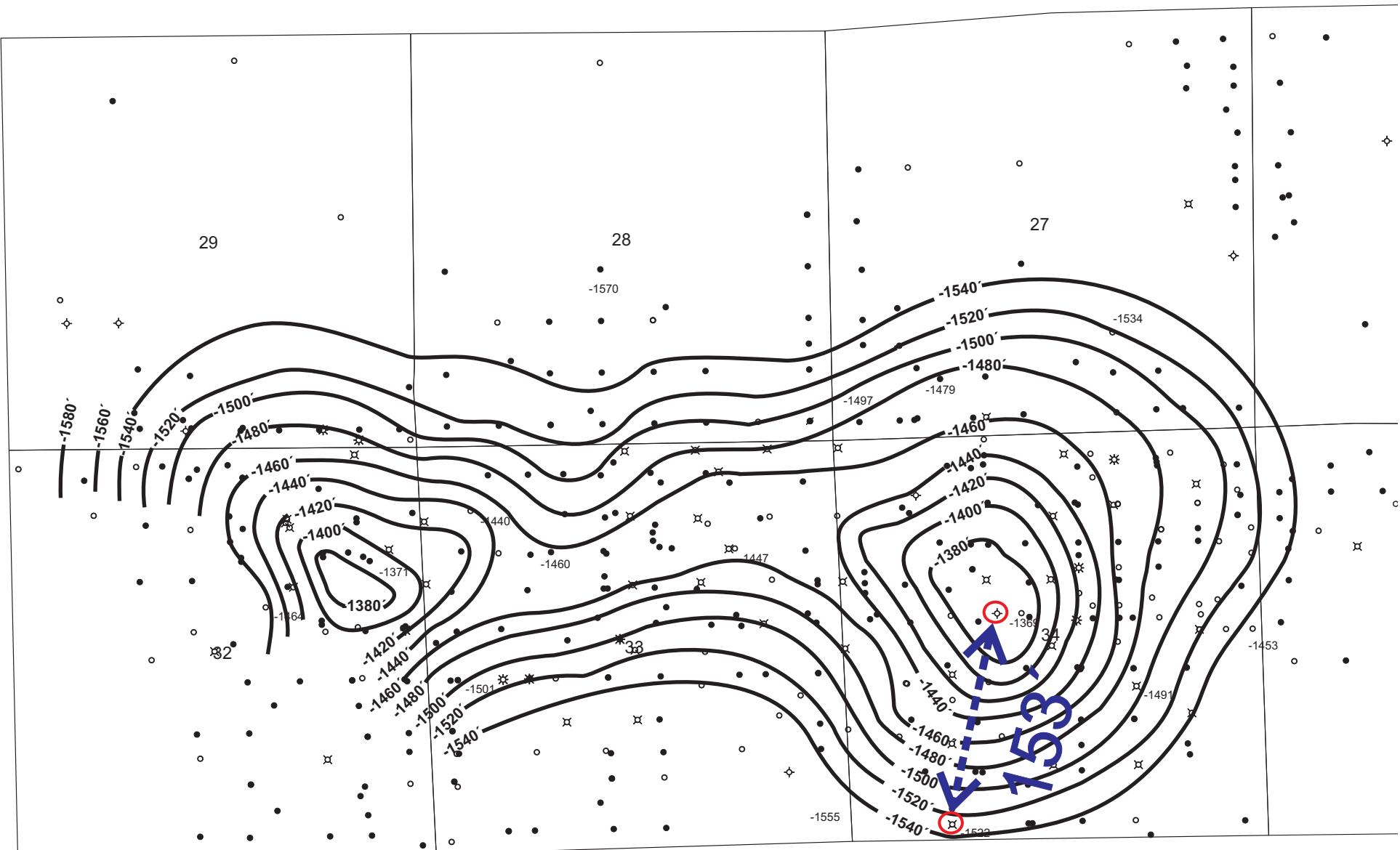


Fig. 10a

Arbuckle Isopach

**Wildhorse Field
Osage County, OK**



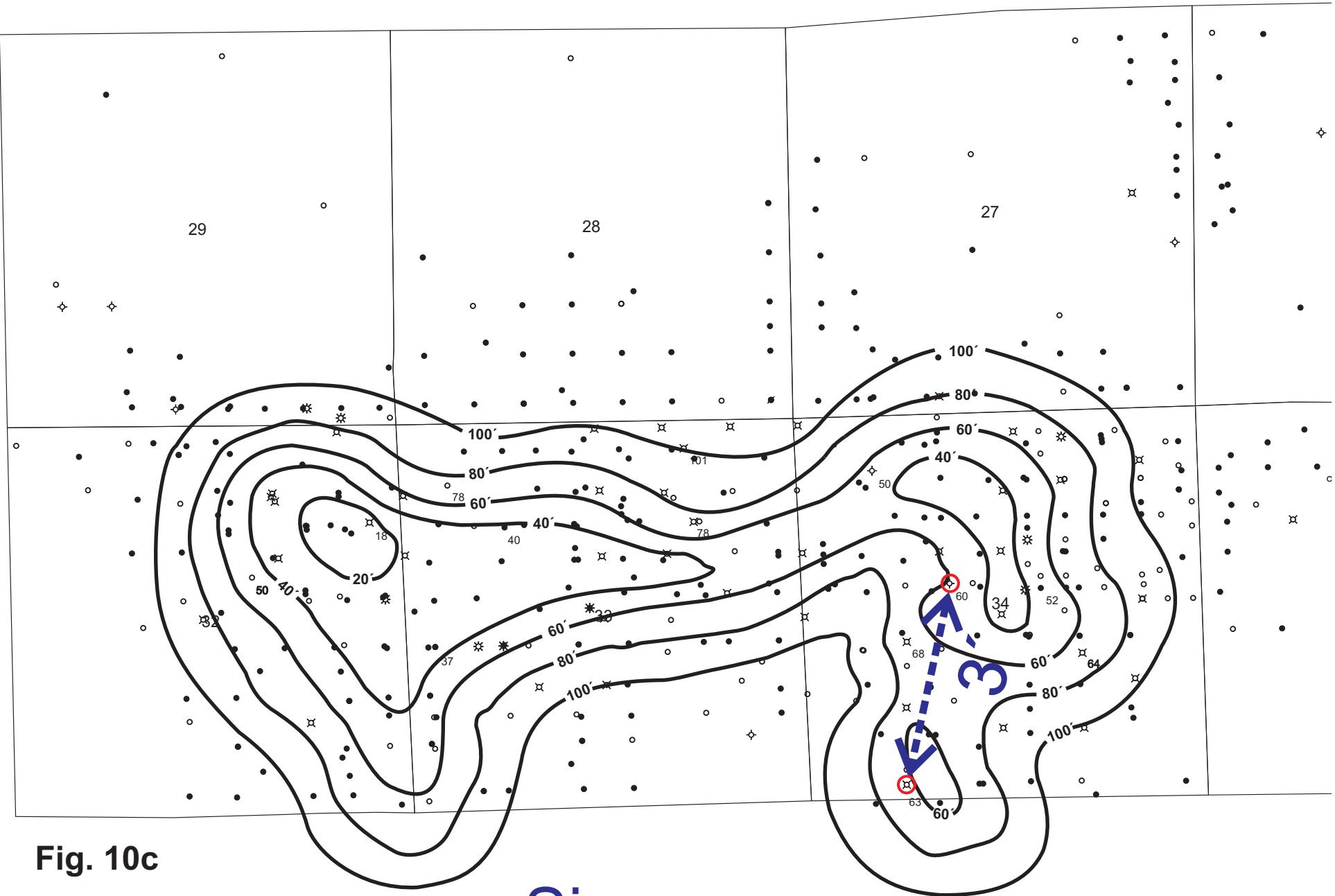


Fig. 10c

Simpson Isopach

Wildhorse Field Study
Osage County, OK

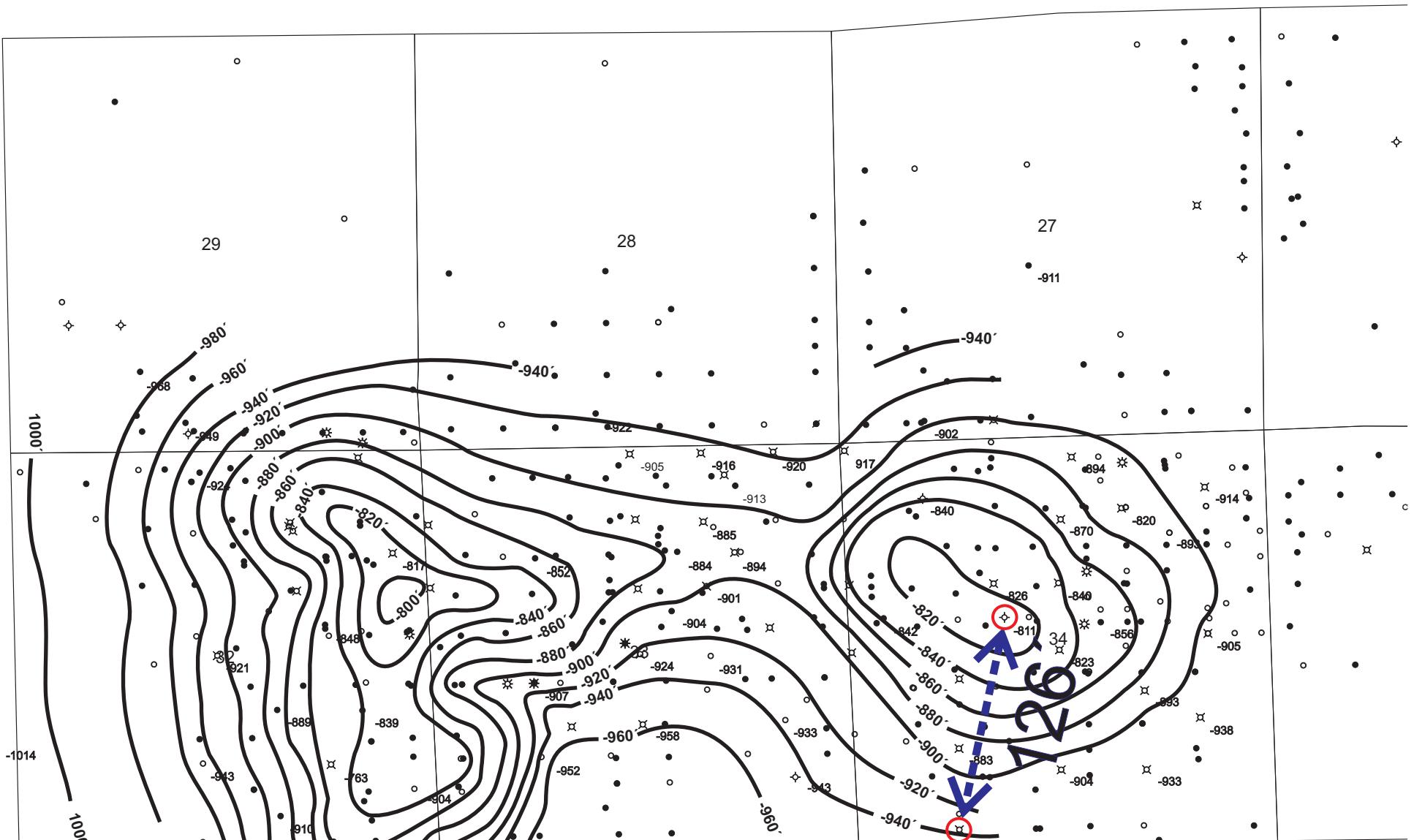


Fig. 10d

Pink Structure

Wildhorse Field
Osage County, OK

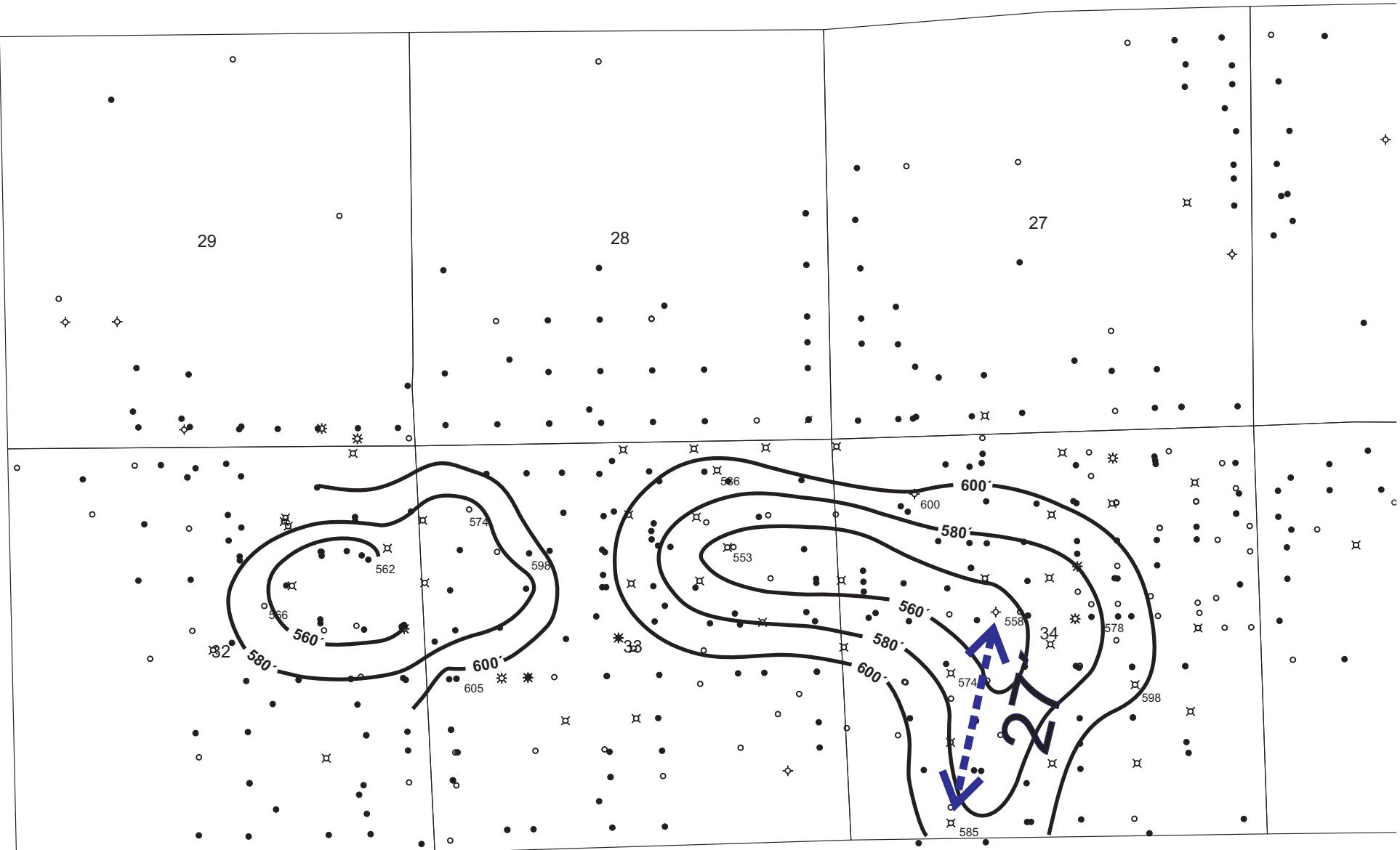


Fig. 10e

Pink-Simpson Structure

Wildhorse Field
Osage County, OK

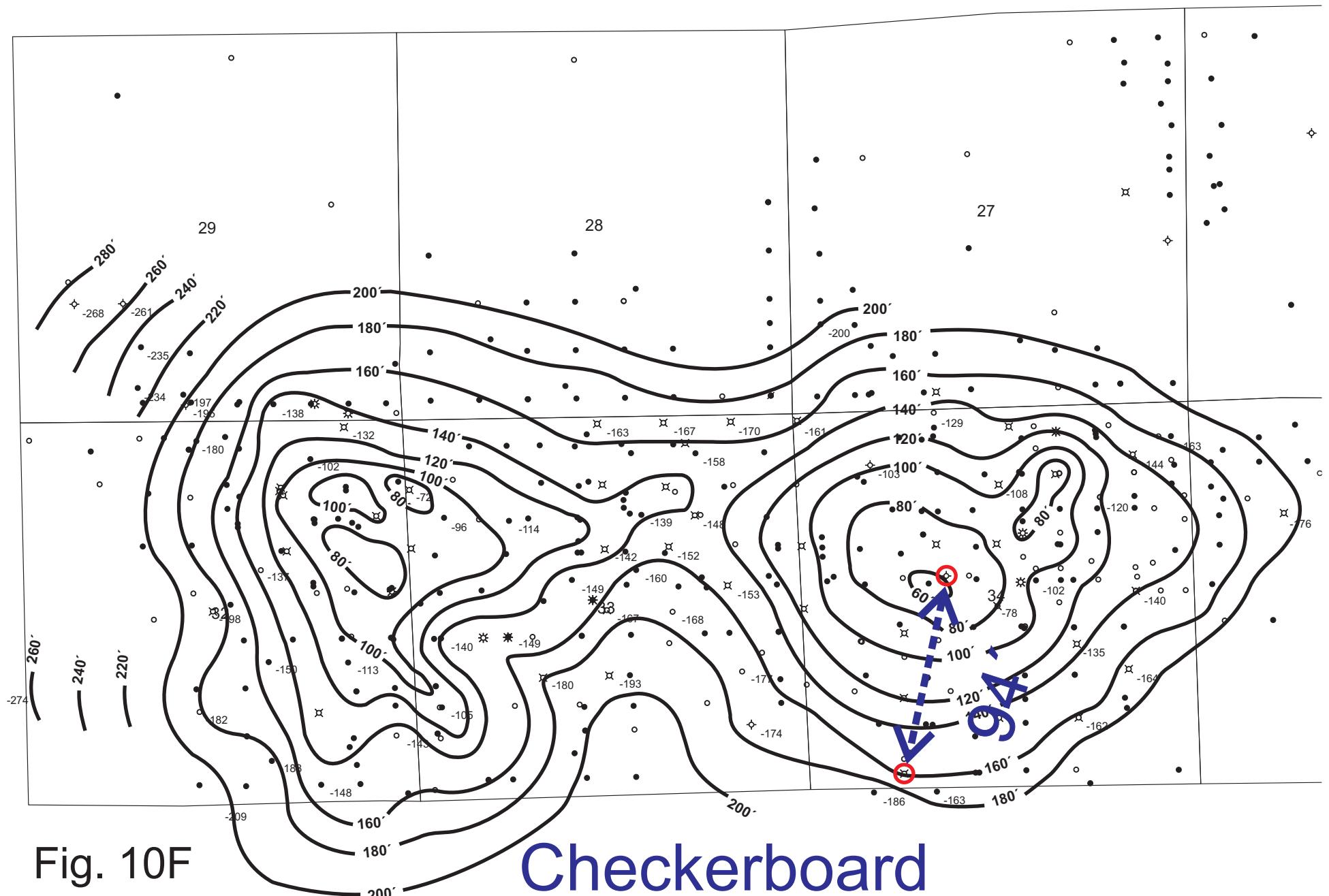


Fig. 10F

Checkerboard Structure

Wildhorse Field Osage County, OK

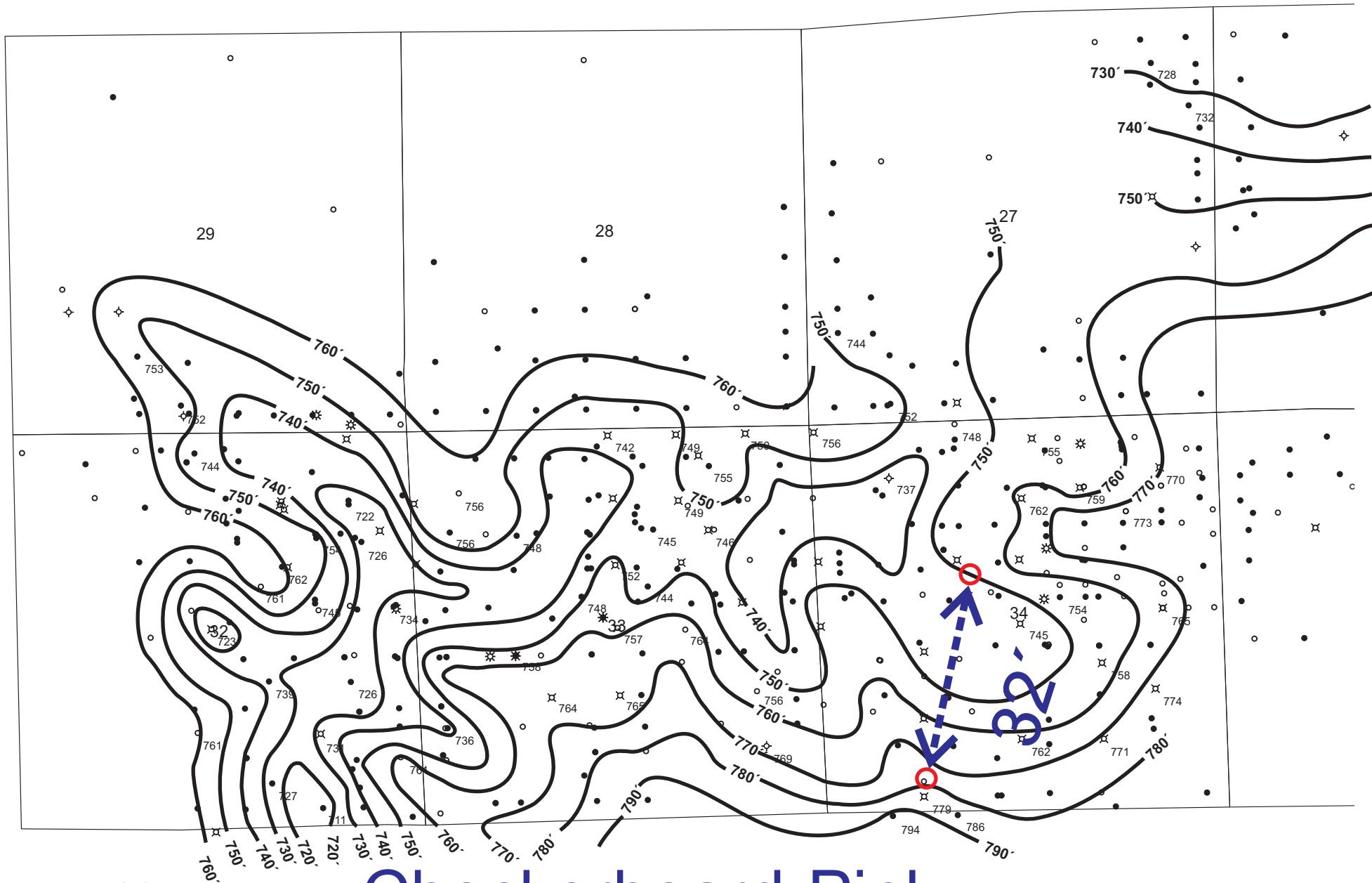


Fig. 10g

Checkerboard-Pink Isopach

Wildhorse Field
Osage County, OK

Assumptions

1. Tops of the Checkerboard and Pink Ls represent time stratigraphic units.
2. Checkerboard and Pink were deposited essentially parallel to sea level. Water depth was similar depth was similar depth for these zones.
3. Compaction was predominately responsible for isopach variation. Little or no structural uplift occurred.
4. Deposition remained consistent with subsidence.
5. Oil/gas production relationship suggest migration was pre-Woodford erosion. Oil/water contacts in the Penn suggest post Checkerboard oil migration.
6. Structural compaction was zero over Arbuckle “0’s”.
7. Arbuckle predominately uniform in composition with little lateral variation which influence compaction.

Fig 10 h

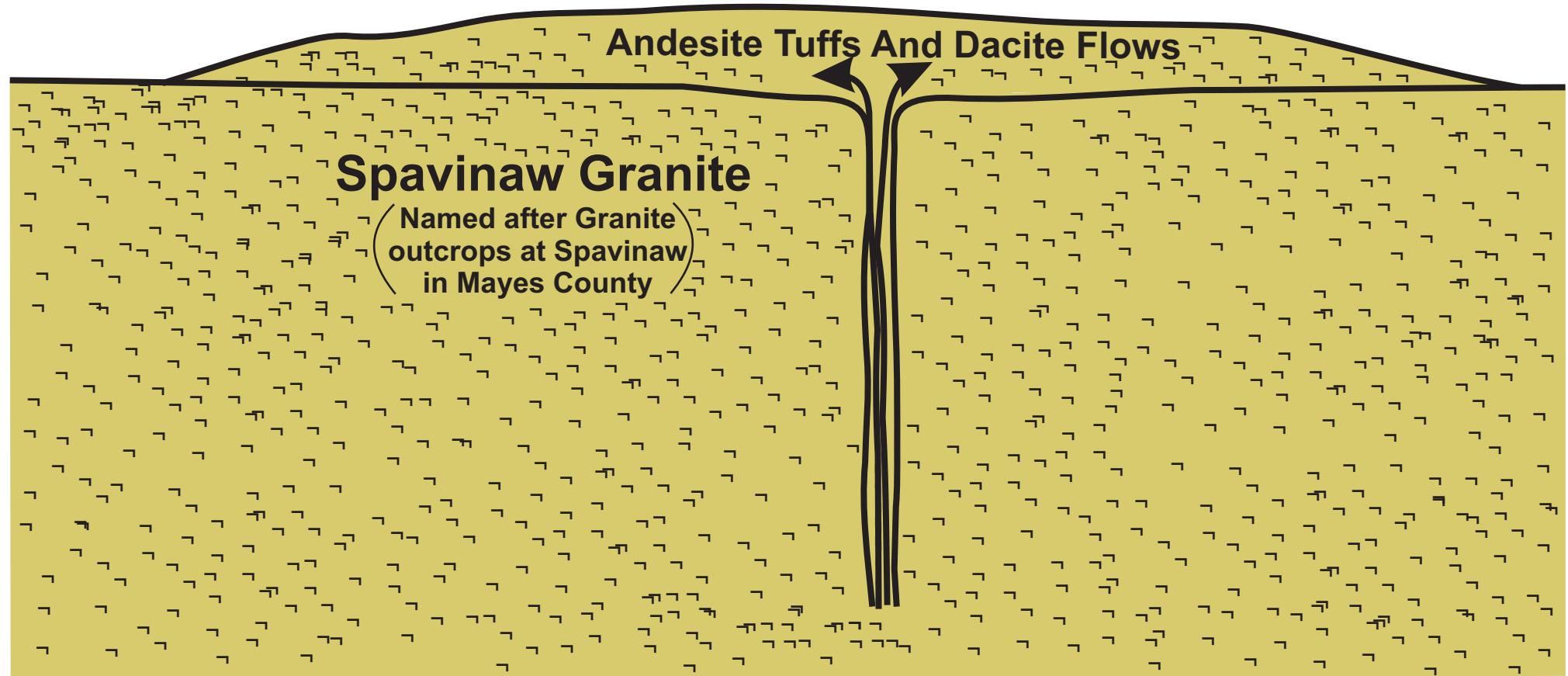


Fig. 12 Paleozoic formation

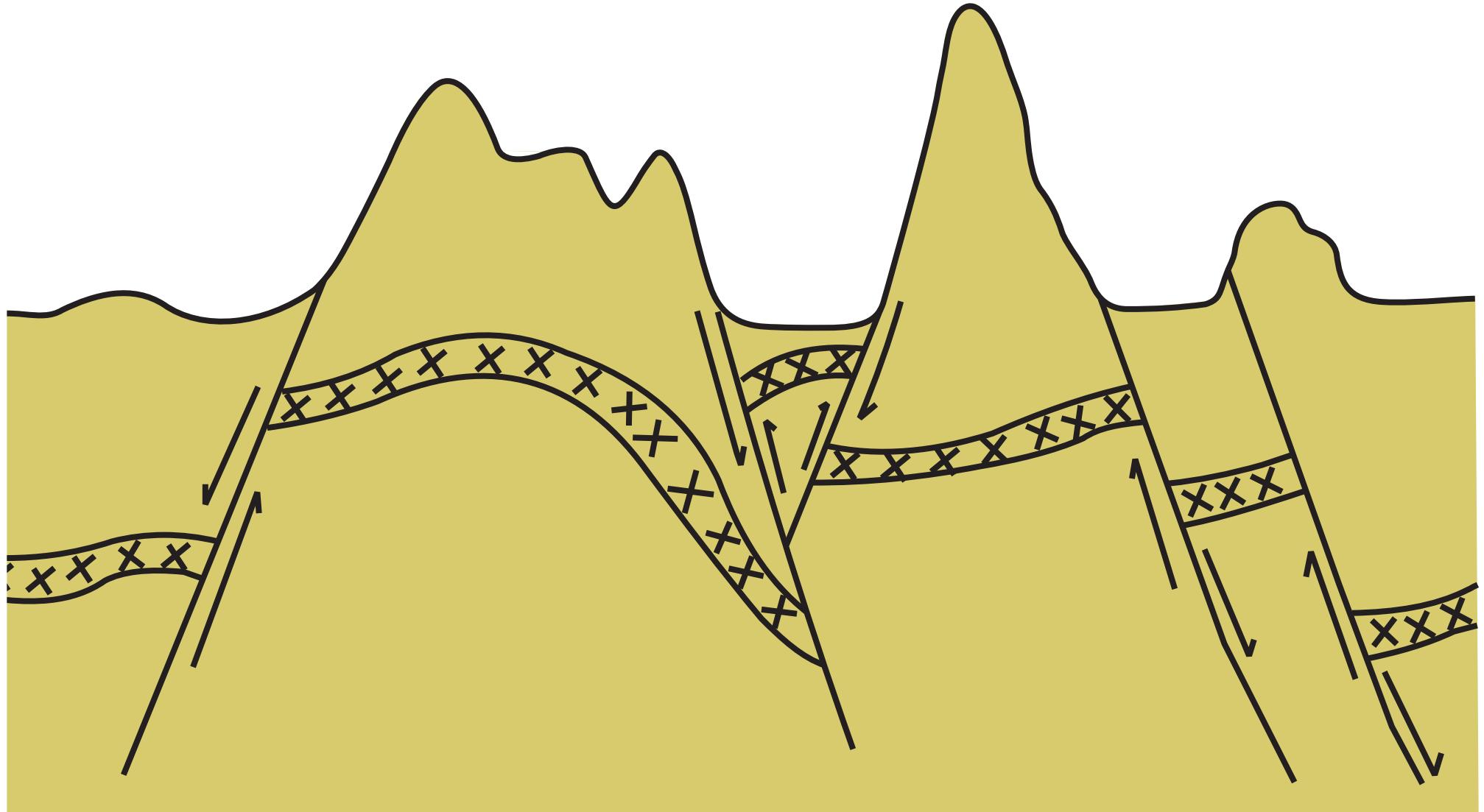
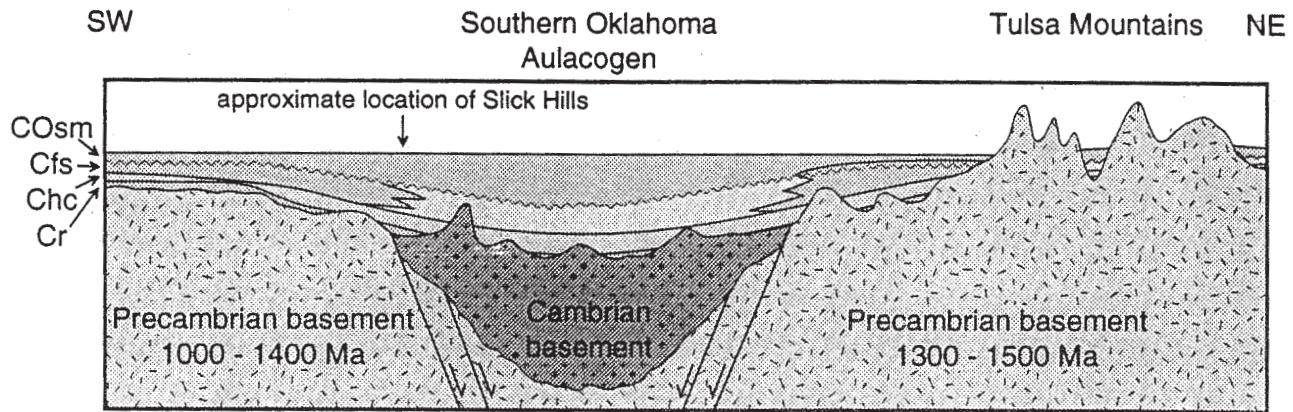
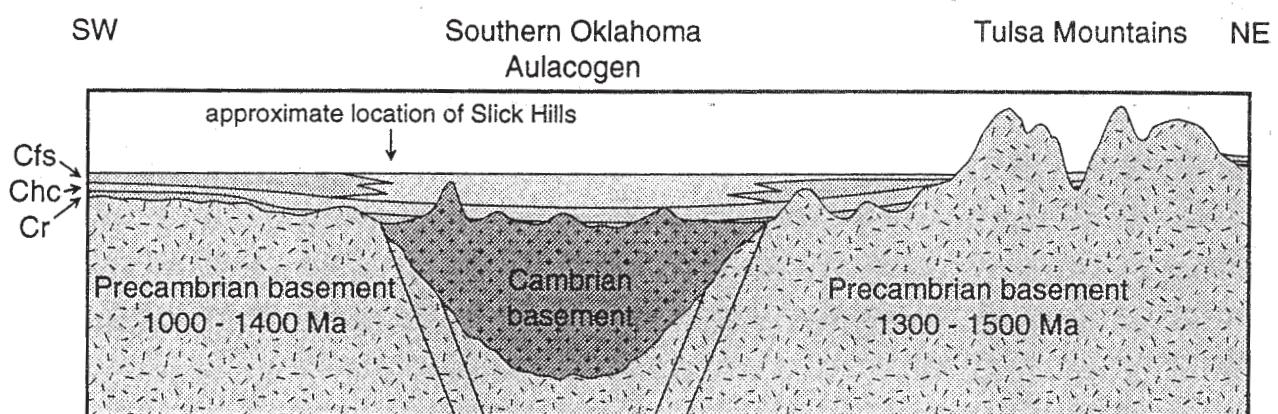


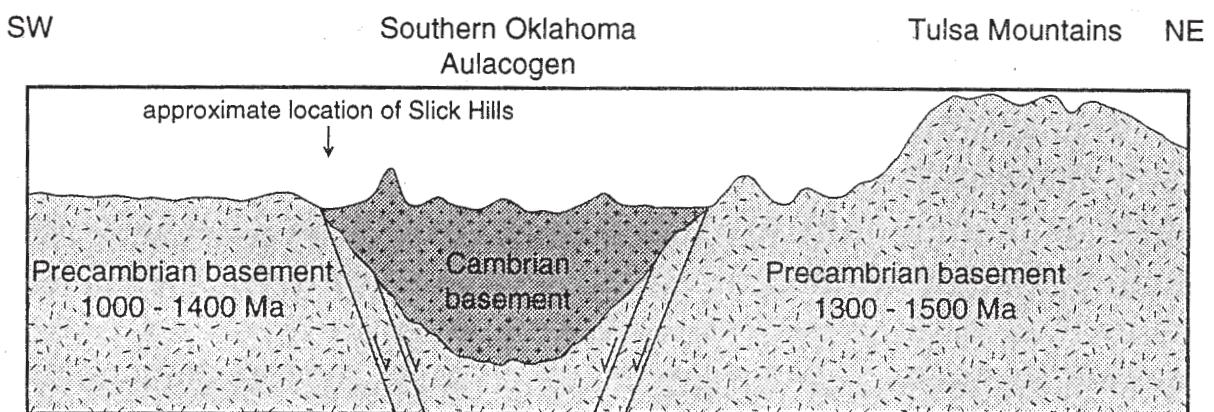
Fig 13 Faulting, folding and erosion of the Paleozoic



STAGE III - Reactivation of basin margin faults and deposition of Cambro-Ordovician Signal Mountain Formation in deeper waters? (Trempealeauan - Canadian)



STAGE II - Deposition of Cambrian Timbered Hills Group and Cambrian Fort Sill Formation. (Franconian)



STAGE I - Development of rift basin. (Middle Cambrian)

Hosey and Donovan, 2000
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Fig. 14 Arbuckle deposition

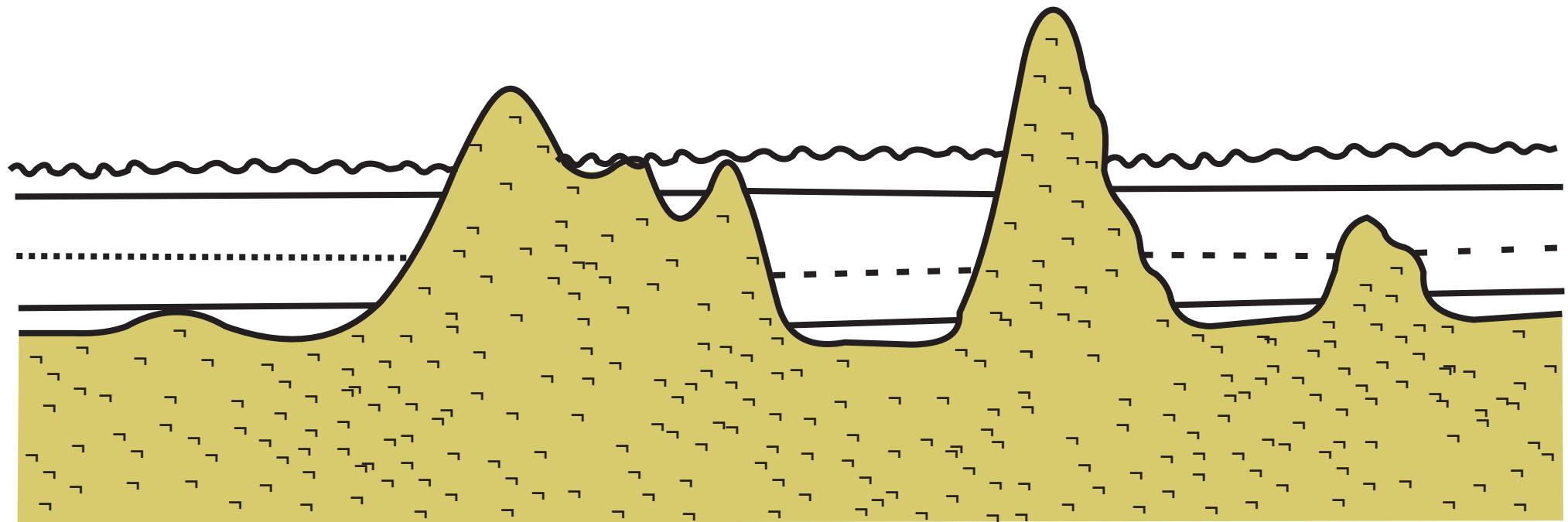


Fig 15 Deposition of the Arbuckle

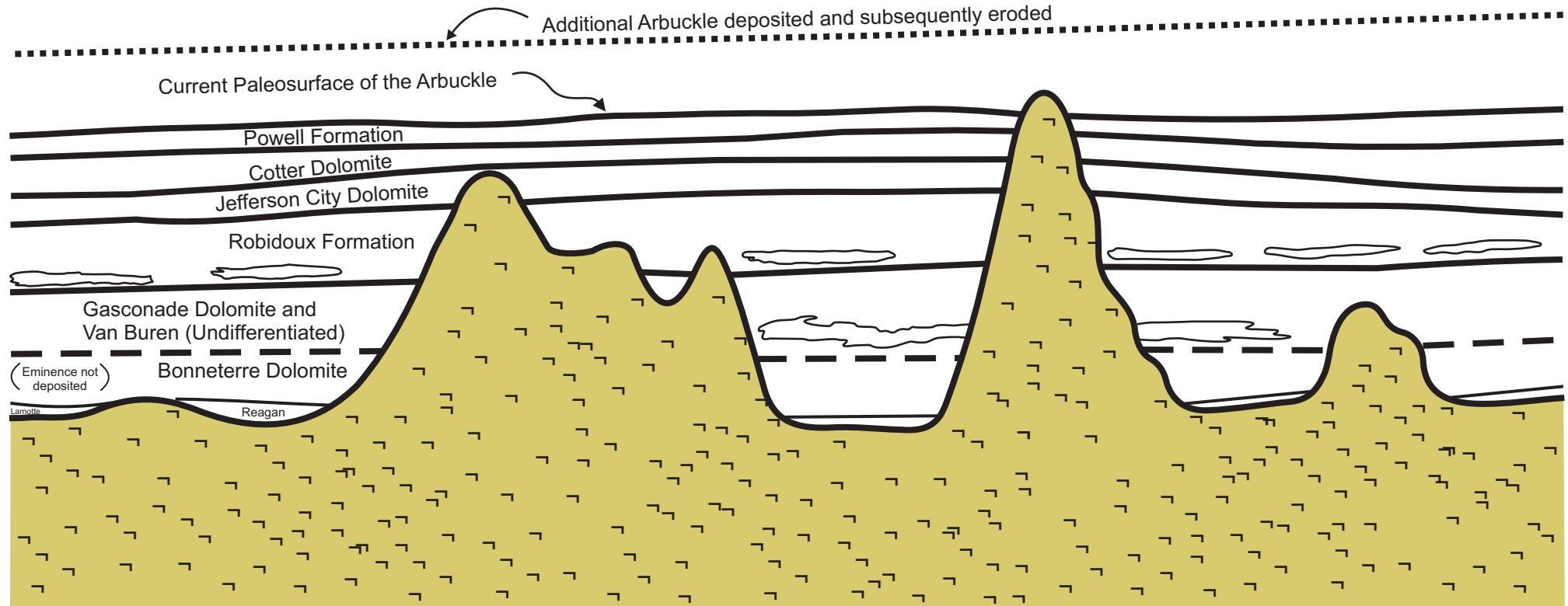


Fig 16 Continued deposition of the Arbuckle

Deposition of Simpson and Subsequent Strata

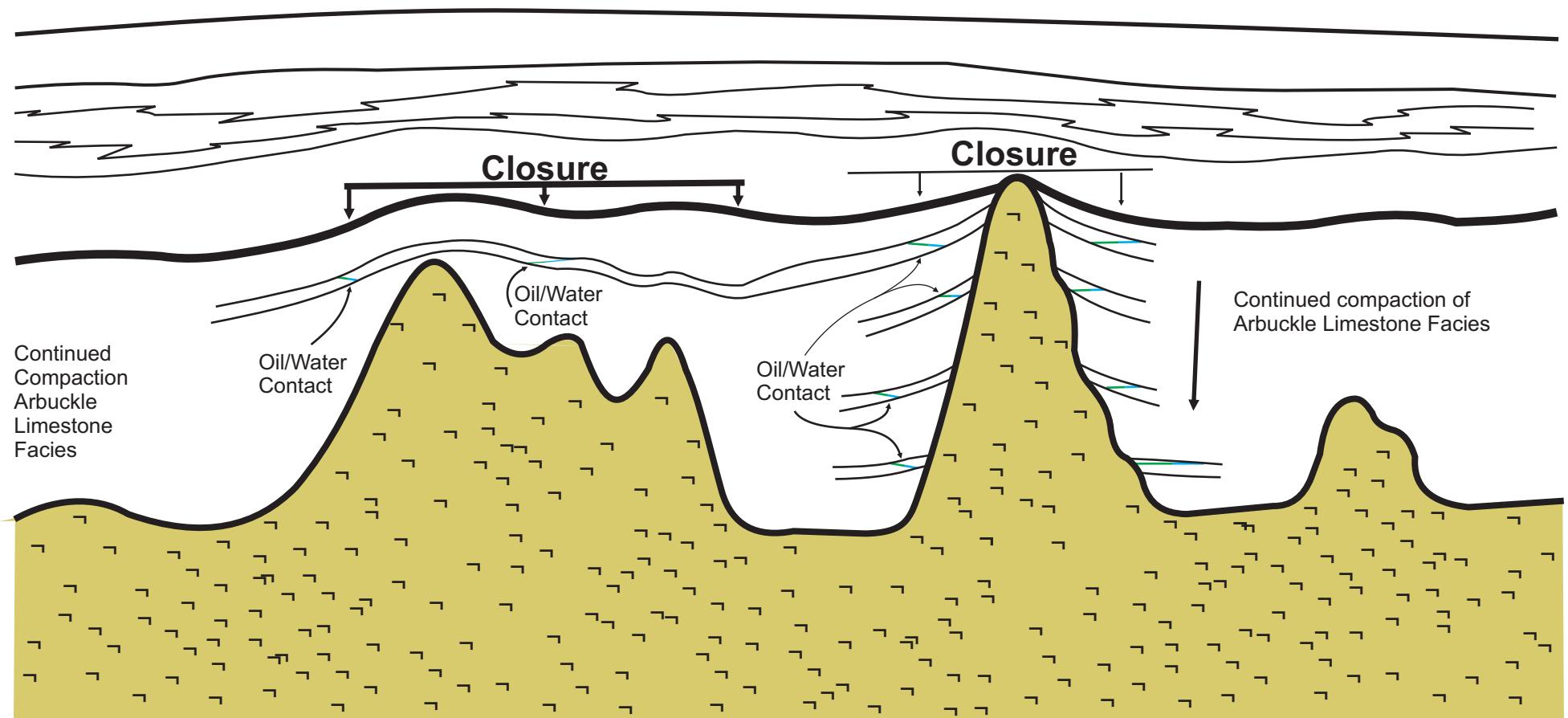


Fig 17 Compaction of Arbuckle sediments

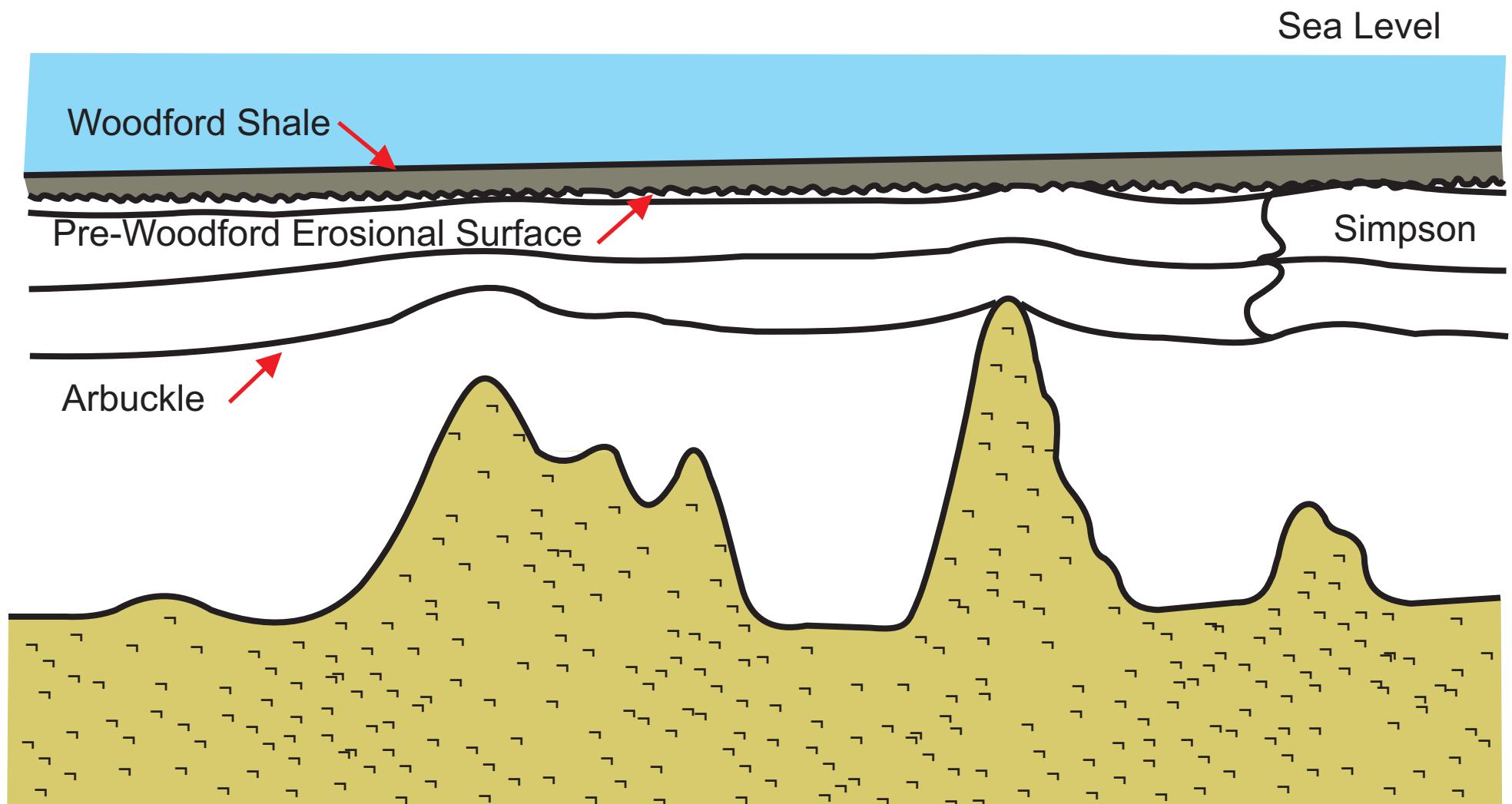


Fig 18

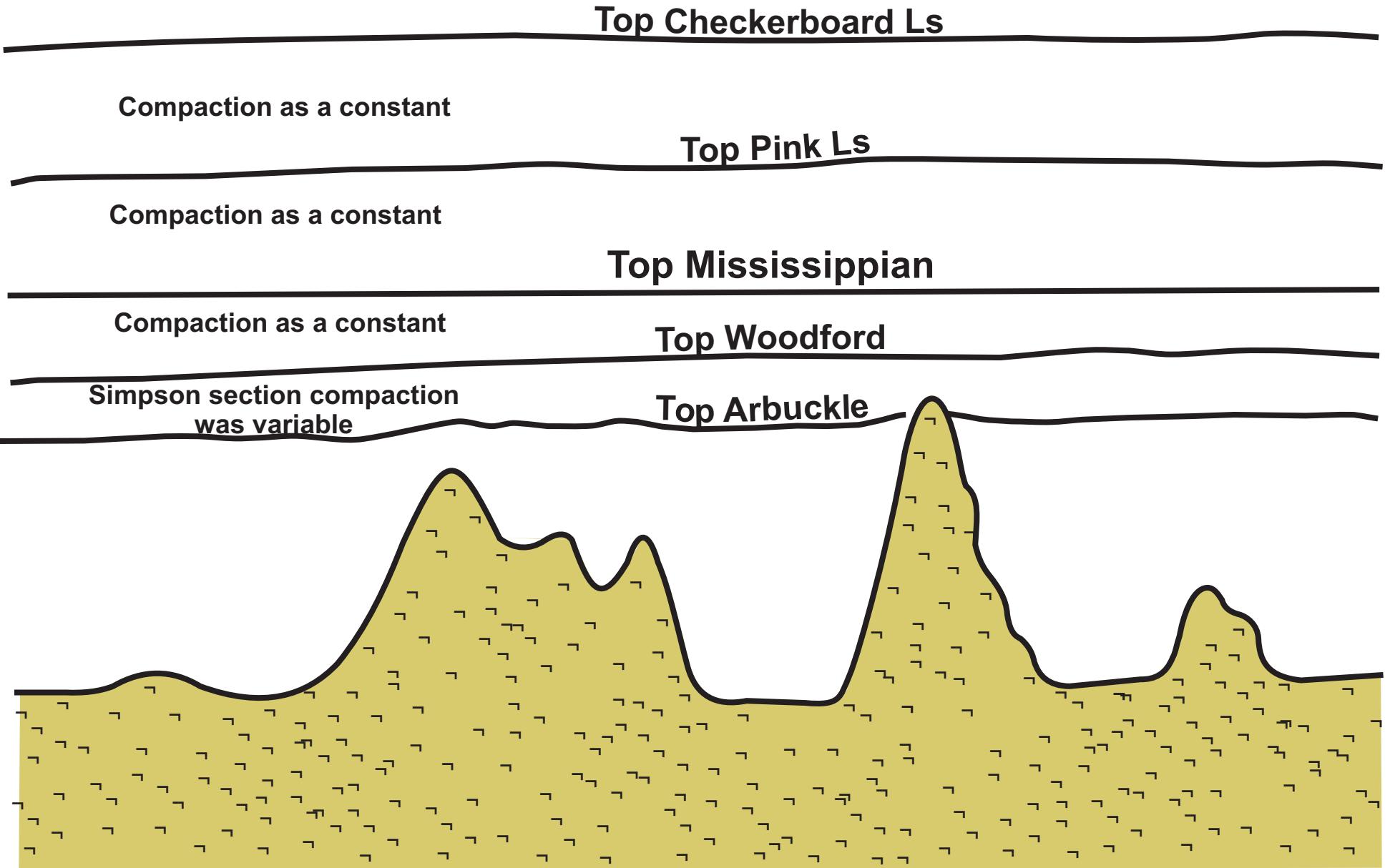


Fig 19 Deposition of Pink Ls and Checkerboard Ls

“The Dolomite Problem”

Dolomite is one of the commonest sedimentary materials. No geologic evidence that its formation was under unusual conditions of temperature or pressure. Efforts to prepare dolomite in the lab under simulated sedimentary conditions have failed. No dolomite is observed to be forming in nature in ordinary sedimentary environments.

Crystallography

Dolomite is a double carbonate of Mg and Ca which the anions are CO_3^{2-} and cations are REGULARLY alternation Ca^{++} and Mg^{++} .

This is a special, highly ordered crystal structure. When attempts are made to precipitate dolomite in the laboratory, the usual result is a mixture of calcite and hydro-magnesite. These are similar compounds found in shells or some marine organisms and are not dolomite but structure with random distribution of Ca and Mg, similar to what is found readily in the laboratory.

Dolomite is readily prepared artificially at temperatures somewhat over 100 degrees. Minerals formed at successively lower temperature result in crystal lattices that have increasingly disordered or random distribution of Ca and Mg in the lattice.

Geological

Not much evidence that dolomites, in older strata, were formed as primary precipitates, except possible evaporate deposits.

Characteristic poor preservation of fossils in dolomites, the coarseness of grains, commonly observed cavities and pore spaces suggest that dolomites formed subsequent to deposition.

Fig 19 a

Source: Introduction to Geochemistry
by Konrad Krauskopf

BURIAL DOLOMITIZATION OF THE UPPER DEVONIAN MIETTE BUILDUP, JASPER NATIONAL PARK, ALBERTA¹

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McGill University, Montreal, Quebec H3A 2A7

ABSTRACT

The periphery of the well exposed Miette buildup in the Front Ranges of the Rocky Mountains is heavily dolomitized. Dolomitization is most abundant in a narrow zone less than 1 km wide and decreases rapidly into the buildup interior. Detailed mapping, petrographic analysis, and chemical studies facilitated the recognition of five types of dolomites in the following paragenetic sequence: (1) diffuse microcrystalline dolomite, an early phase of microcrystalline dolomite in lime mudstones and wackestones; (2) 'mosaic' dolomite which replaced large volumes of precursor limestone matrix and forms at least 70 percent of the dolomite in the buildup; (3) pressure solution related dolomite, coarse to fine crystalline, associated with pressure solution interfaces; (4) white sparry dolomite which forms irregular, coarsely crystalline pods in the massive dolomite facies at the buildup margin; and (5) late stage dolomite cements composed of coarse, clear, scalenohedral dolomite crystals forming isopachous crusts which line leached void spaces and fractures associated with brecciation at the buildup margin. Vuggy and intercrystalline porosity is associated with type 2 dolomites, but most reservoir porosity has been destroyed.

The Miette buildup was subaerially exposed only for brief periods and evaporites or evaporite-related features are absent. Petrographic observations supported by geochemical and isotope data indicate that pervasive dolomitization beginning with type 2 dolomites was a late diagenetic phenomenon postdating cementation and lithification of buildup interior deposits. Most of the dolomites in the buildup probably formed by a combination of three processes: (1) migration of brines from adjacent basin muds undergoing compaction, (2) pressure solution, and (3) mixing of near-surface fluids with deep burial brines along fracture controlled conduits. Most of the dolomitizing waters are believed to have been derived from the dewatering of basinal strata adjacent to the buildup and from strata underlying the buildup.

The deep burial regime has been underestimated as an environment conducive to massive dolomitization. Significant kinetic obstacles which hinder the formation of ordered dolomite at or near the surface may be largely overcome with increasing burial depths within a sedimentary basin.

CALCITE CEMENTATION

Stage 1

Solution event I

Stage 2

Solution event II

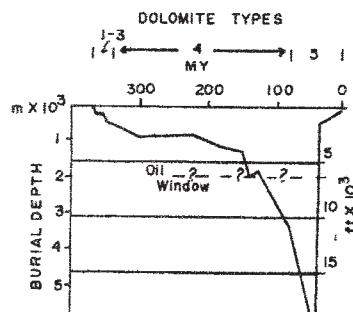
Stage 3

Stage 4

Major solution event III - succeeding the onset of dolomitization

Stage 5

Solution event V - succeeding thrusting and exposure



DOLOMITE FORMATION Type I

??
Type 2

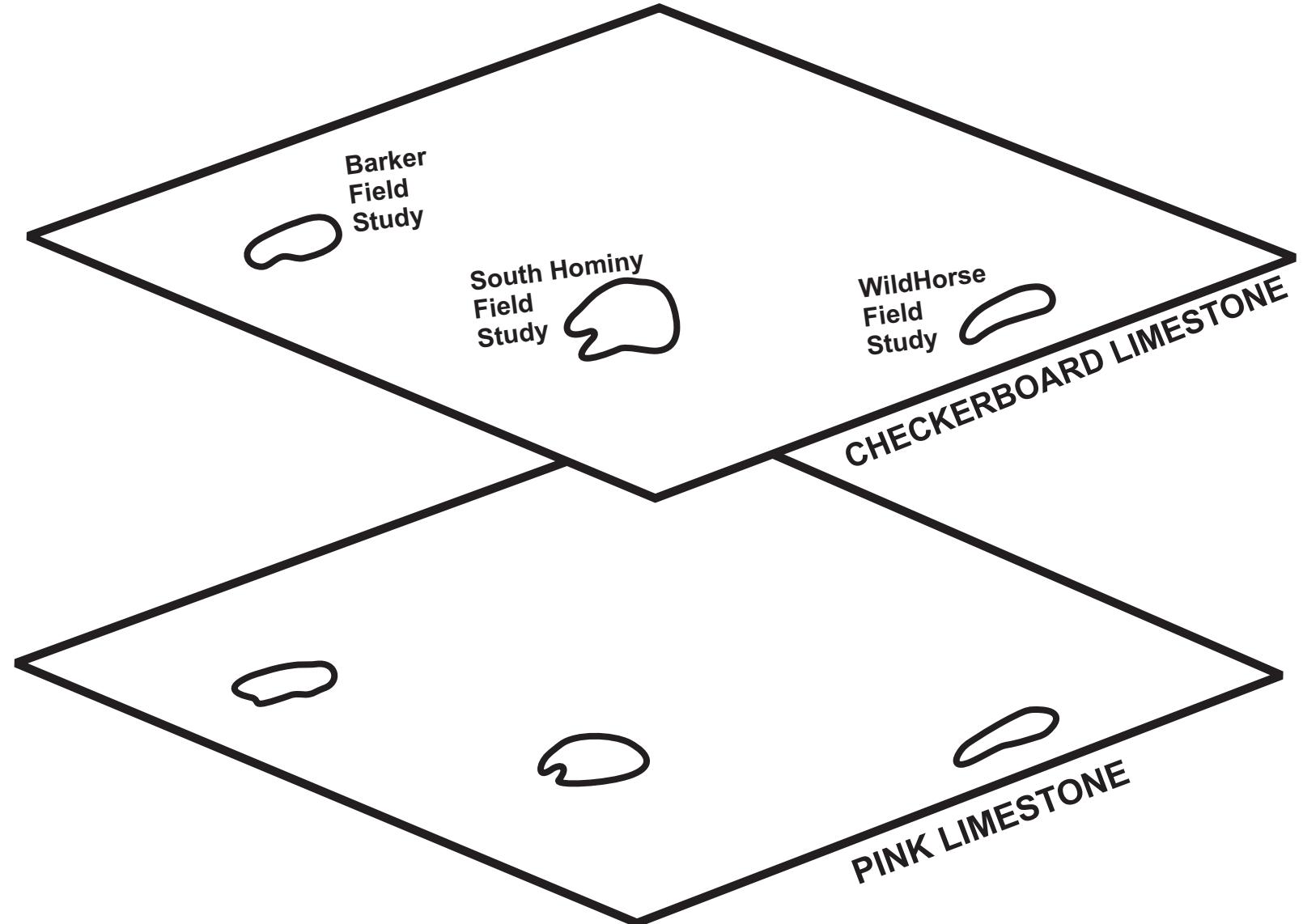
Type 3

Type 4

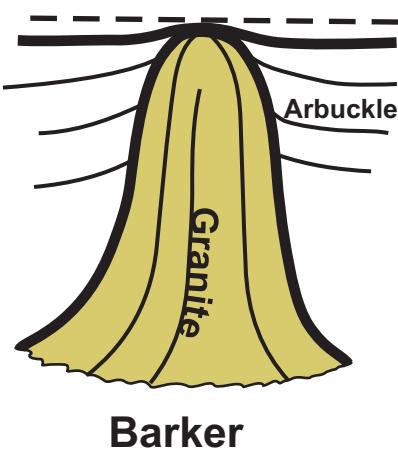
Type 5

INCREASING BURIAL

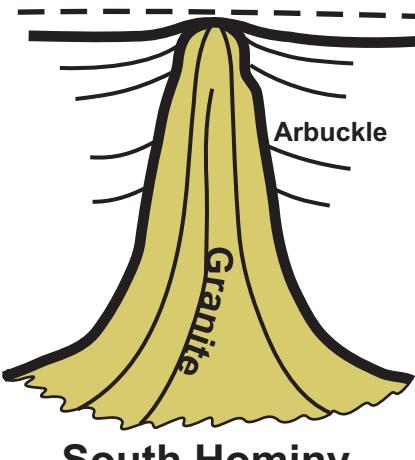
INCREASING TIME



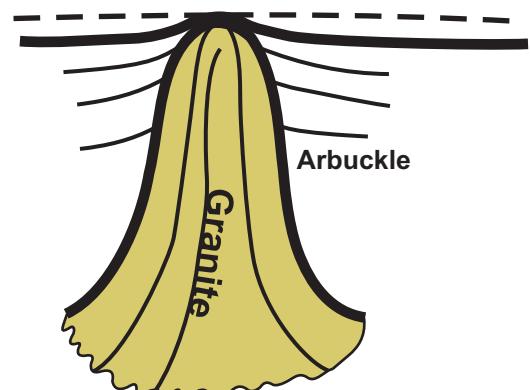
Pre-Structural Compaction



Barker



South Hominy



WildHorse

Fig 20

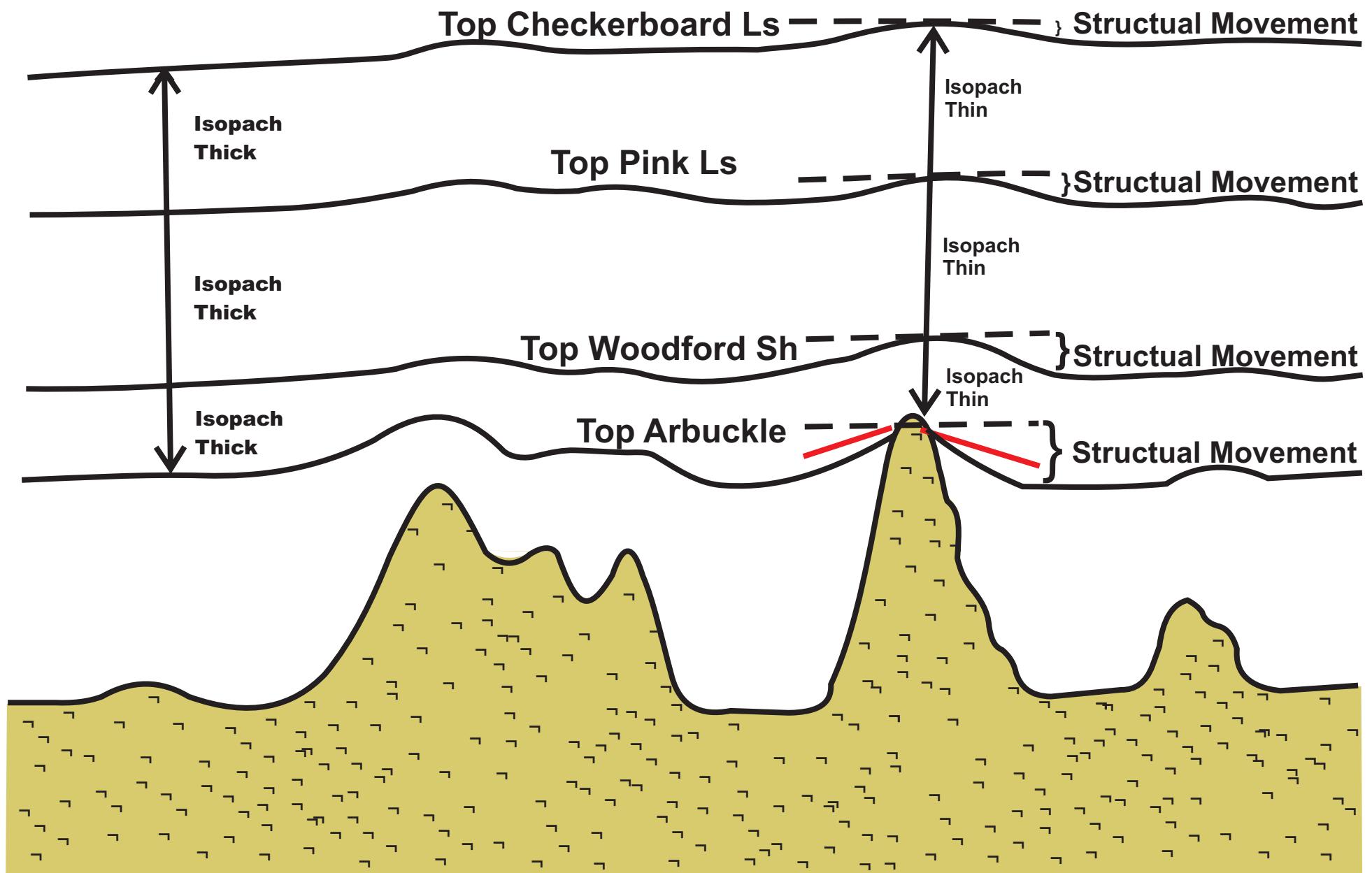


Fig 21

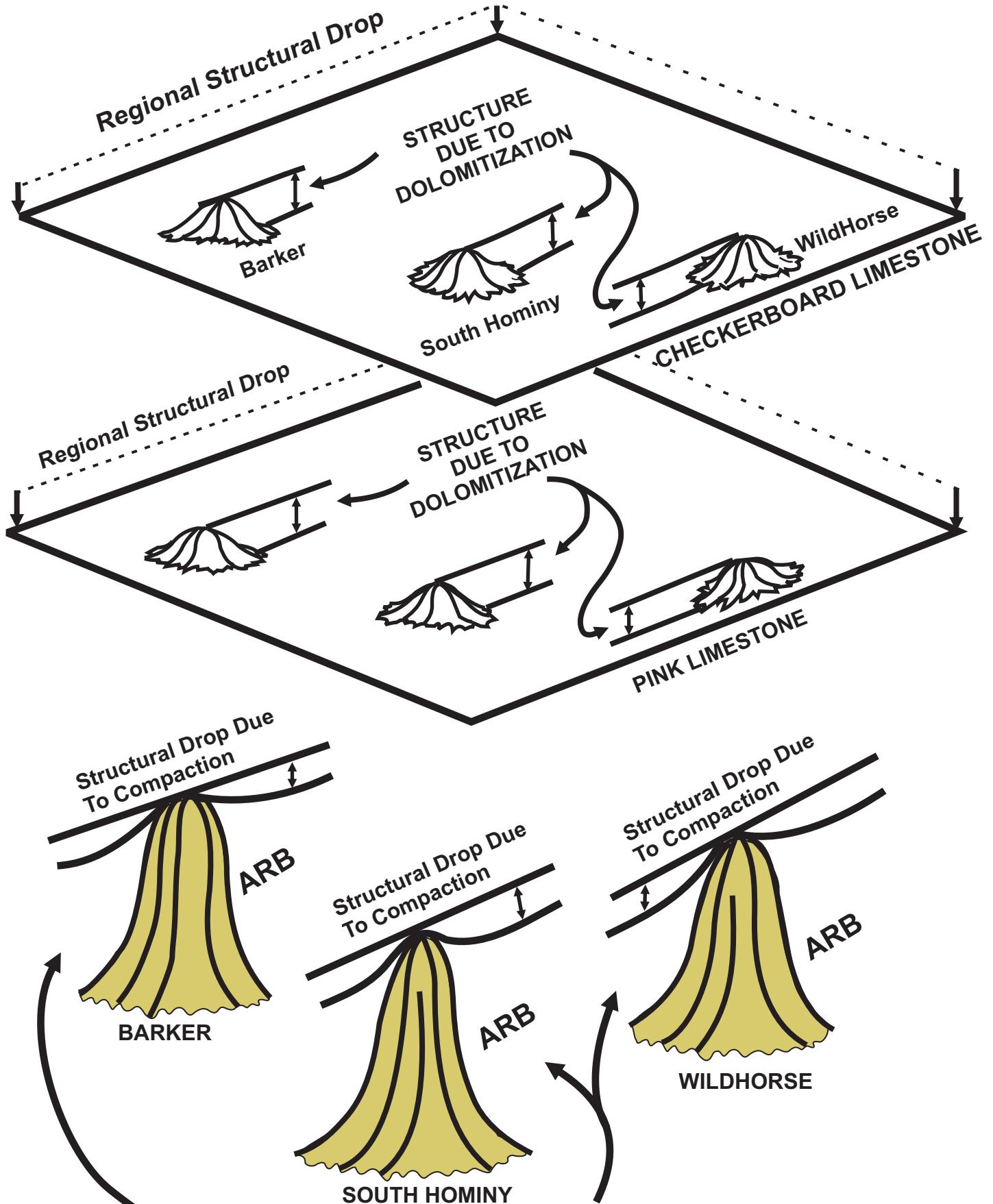


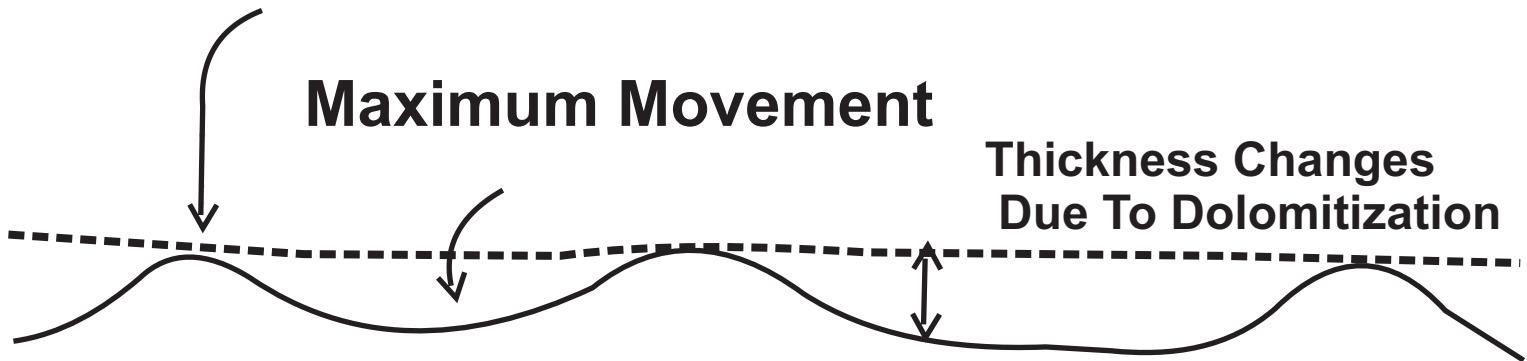
Fig 22

DIAGENETIC CHANGE

No Movement

Maximum Movement

**Thickness Changes
Due To Dolomitization**



No Compaction

Maximum Compaction

**Thickness Changes
Due To Dolomitization**

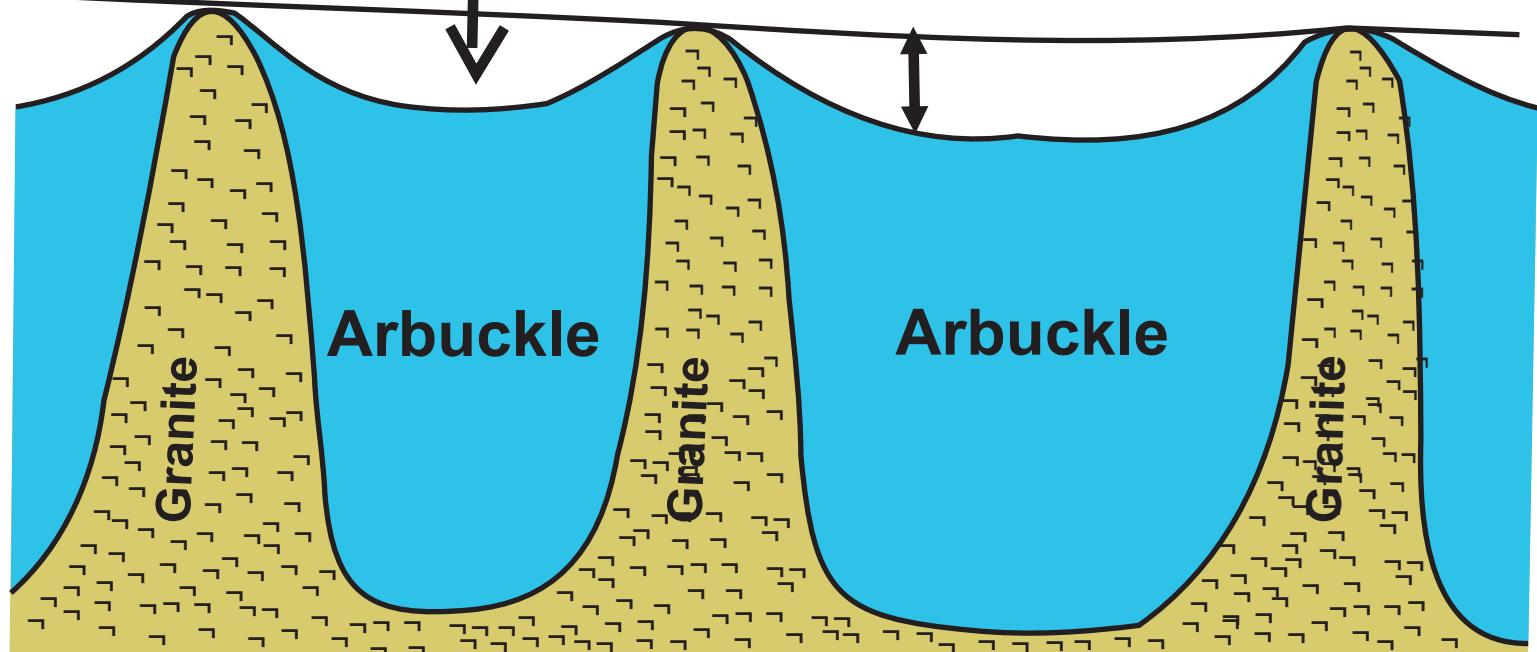


Fig 24

12 23-7

21 24-7

25 Miles

Service Well

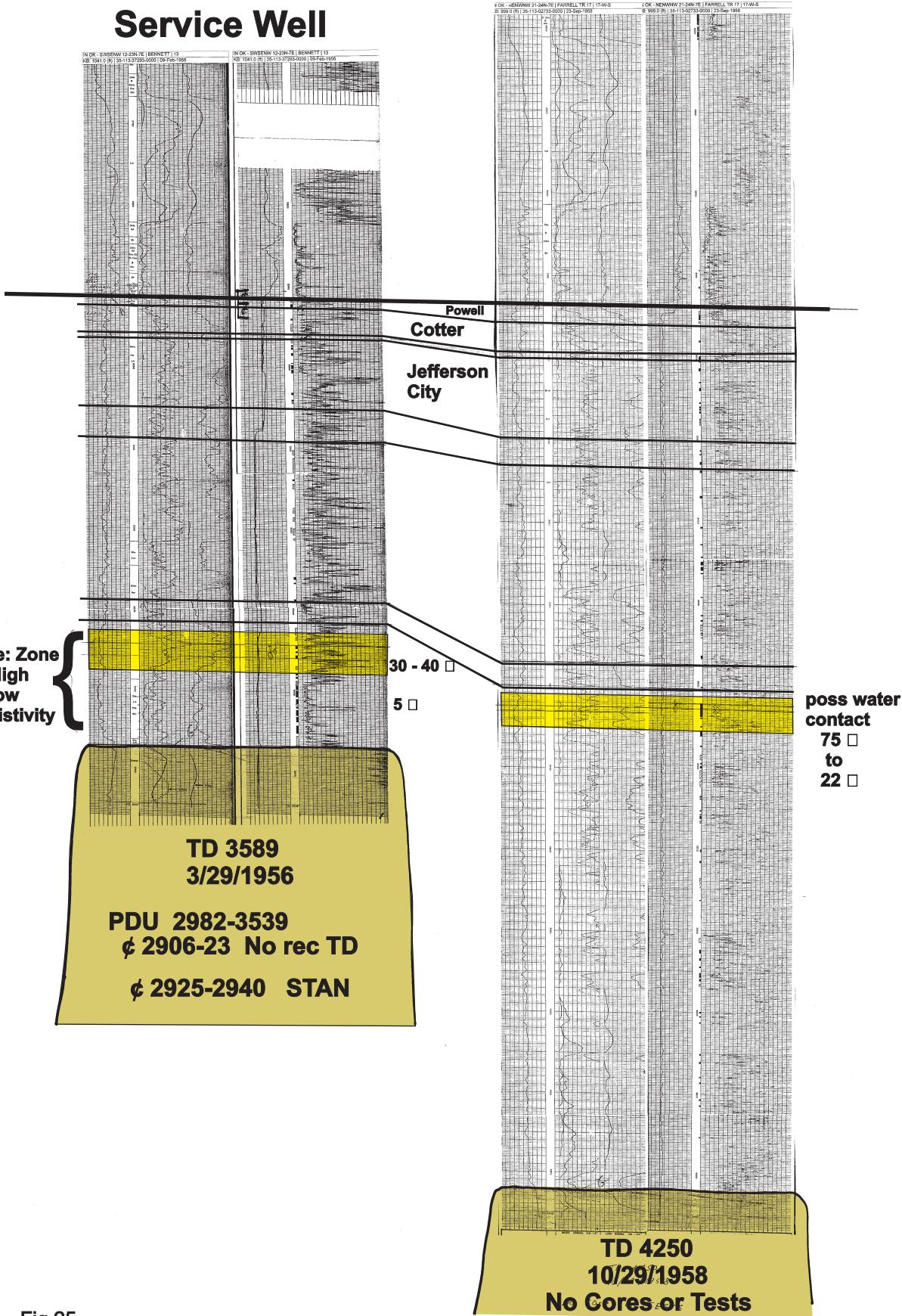
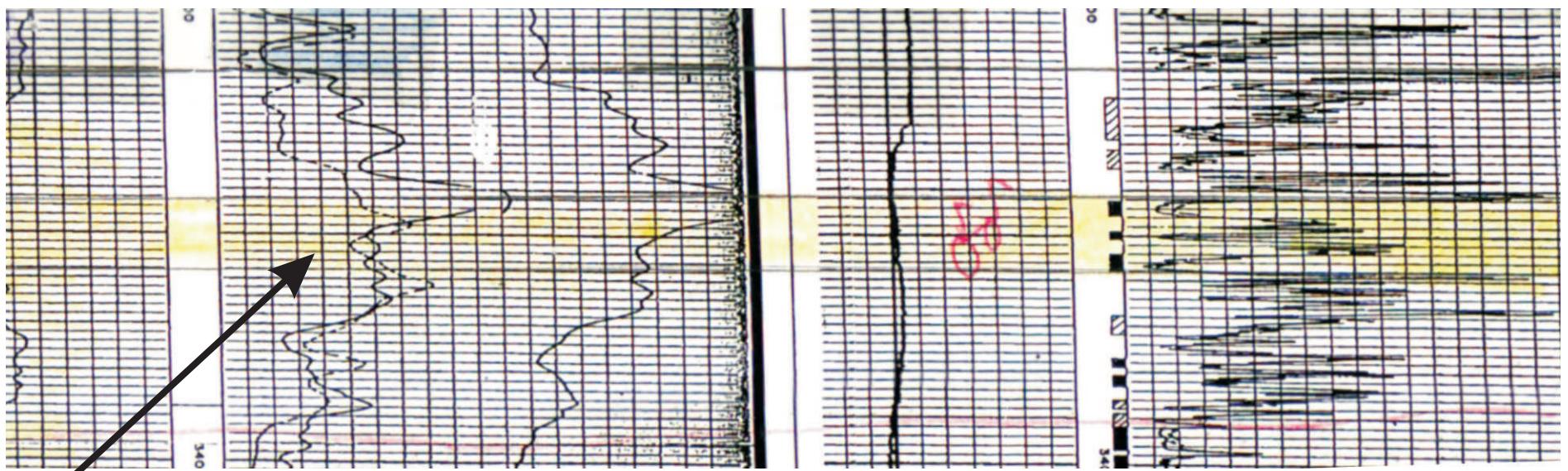


Fig 25



30-40 OHMS RESISTIVITY

Fig 26

16-24-7

1-23-8

9-22-8

8-25-10

10 Miles

8 Miles

22 Miles

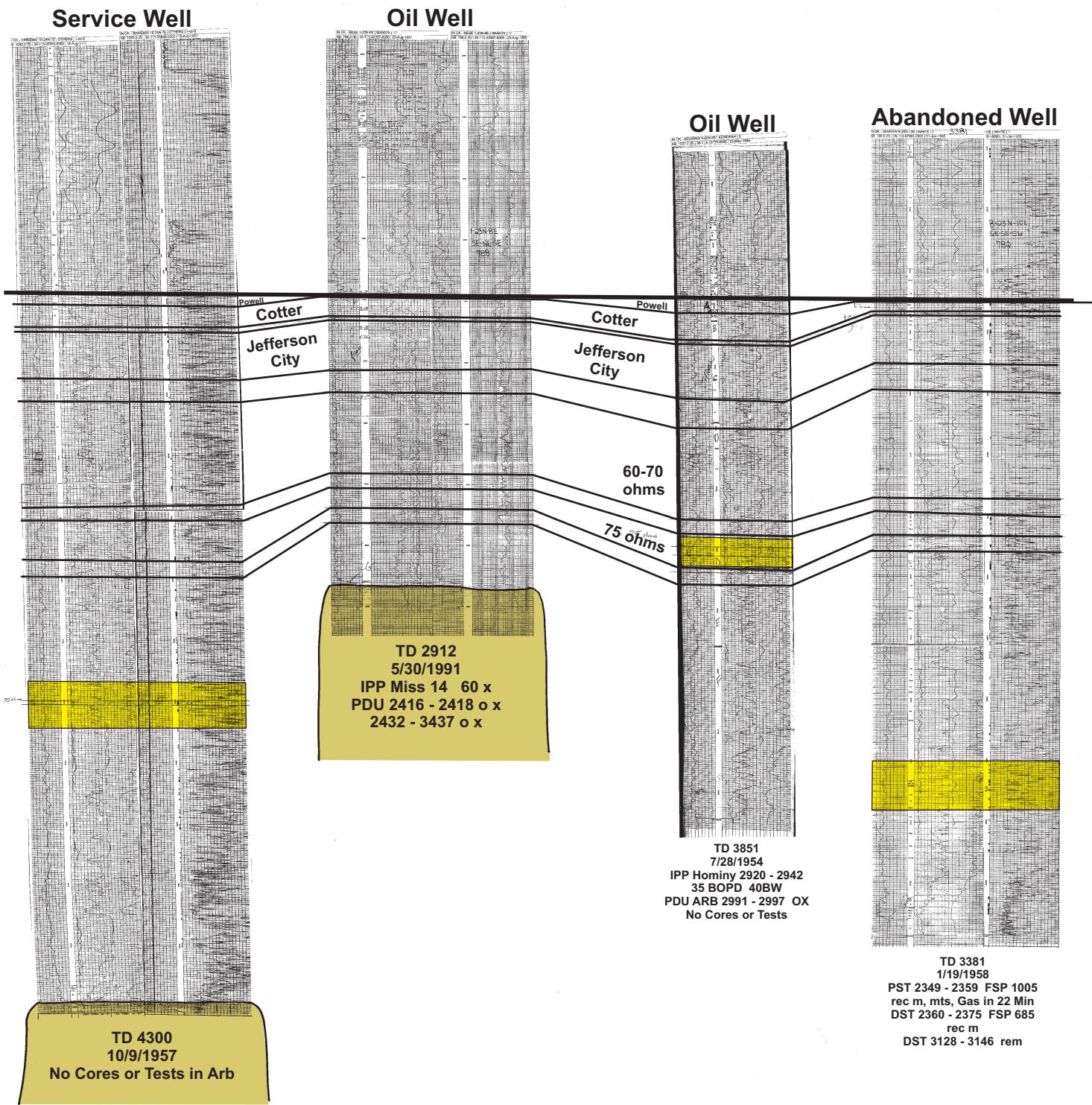


Fig 27

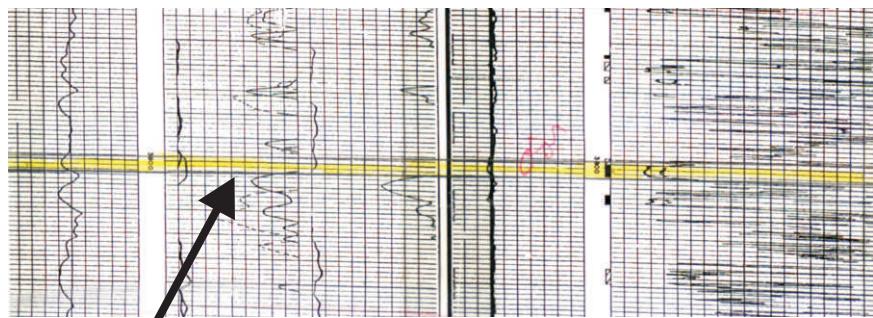


Fig 28a

60-70 OHMS RESISTIVITY

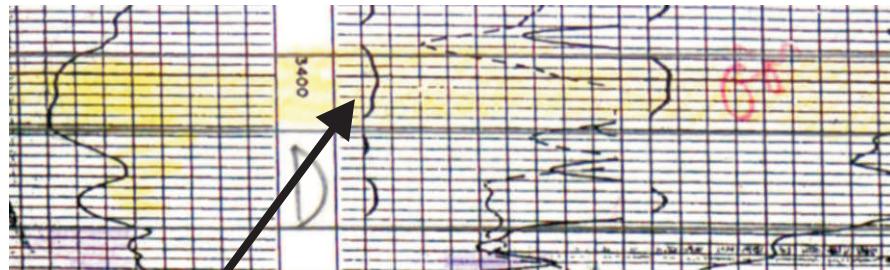
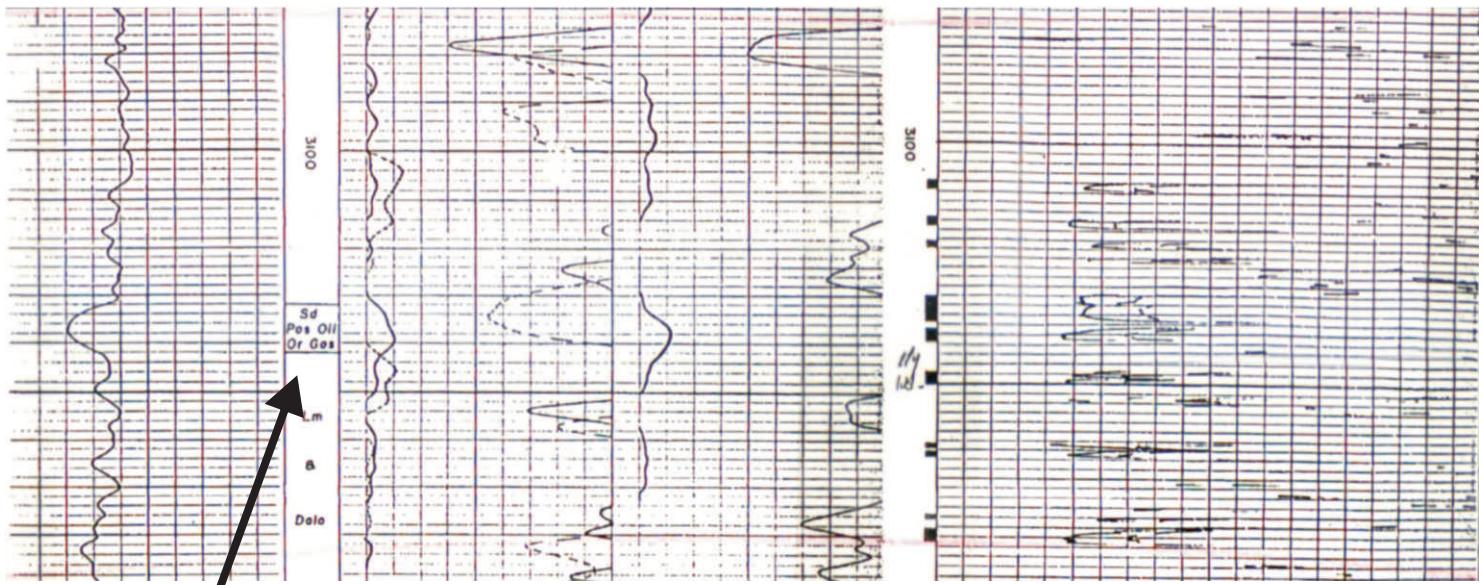


Fig 28b

~70 OHMS RESISTIVITY



LOGGER MAKE NOTES OF POTENTIAL PAY

Fig 28d

T28- R24- S7

T29- R23- S8

Oil Well

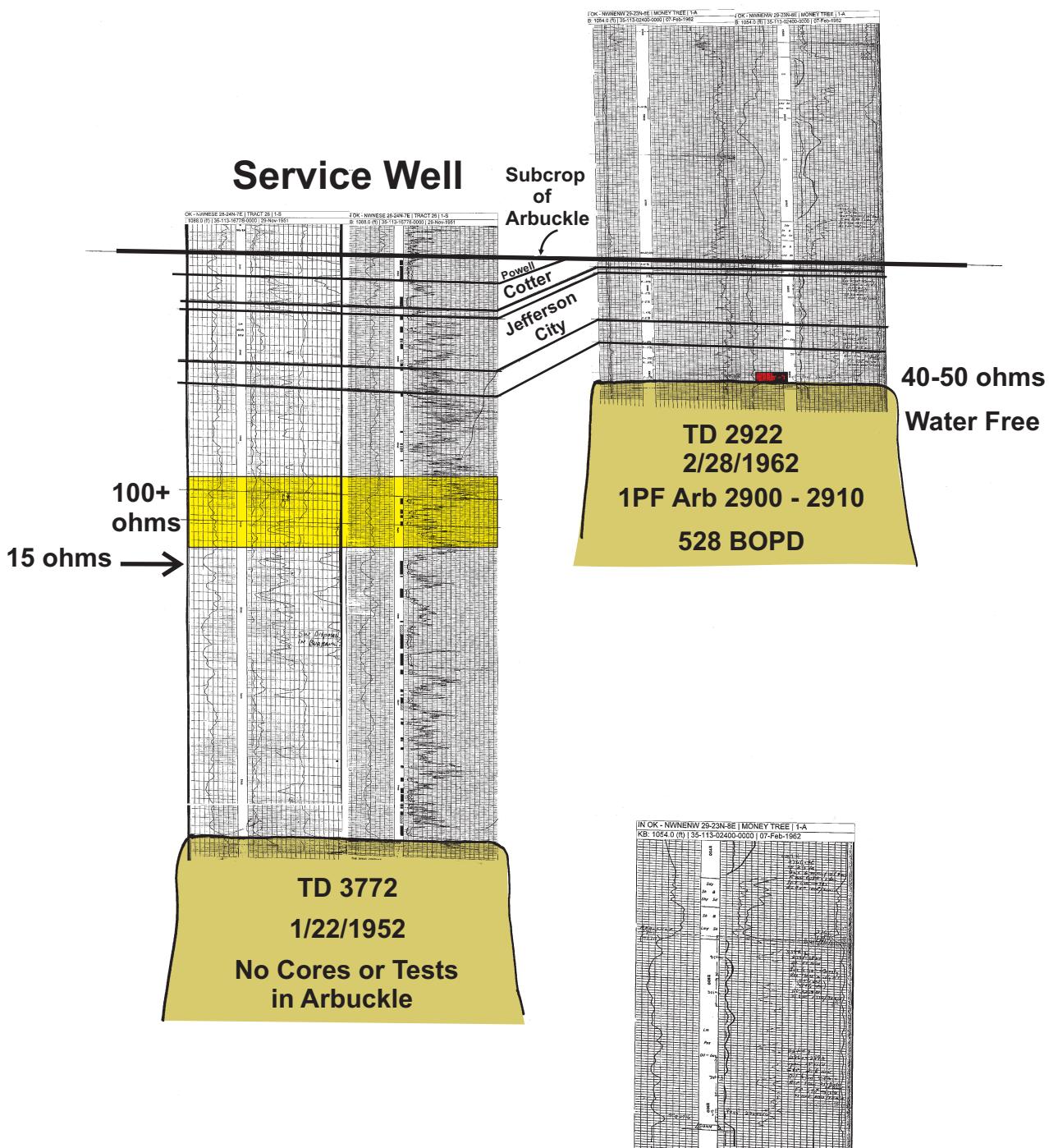
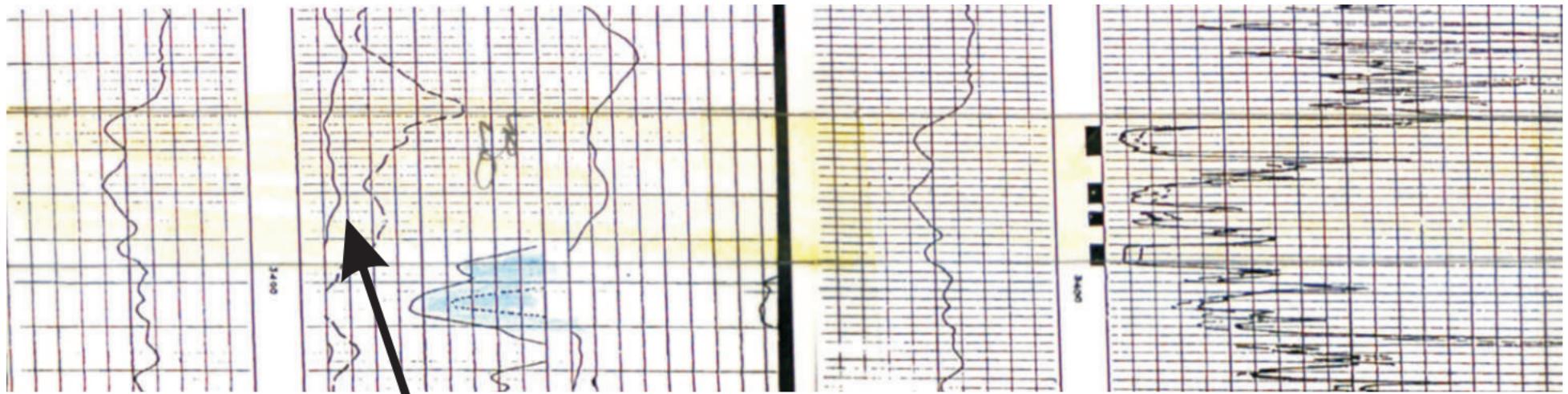


Fig 29



50-75 OHMS RESISTIVITY

Fig 30

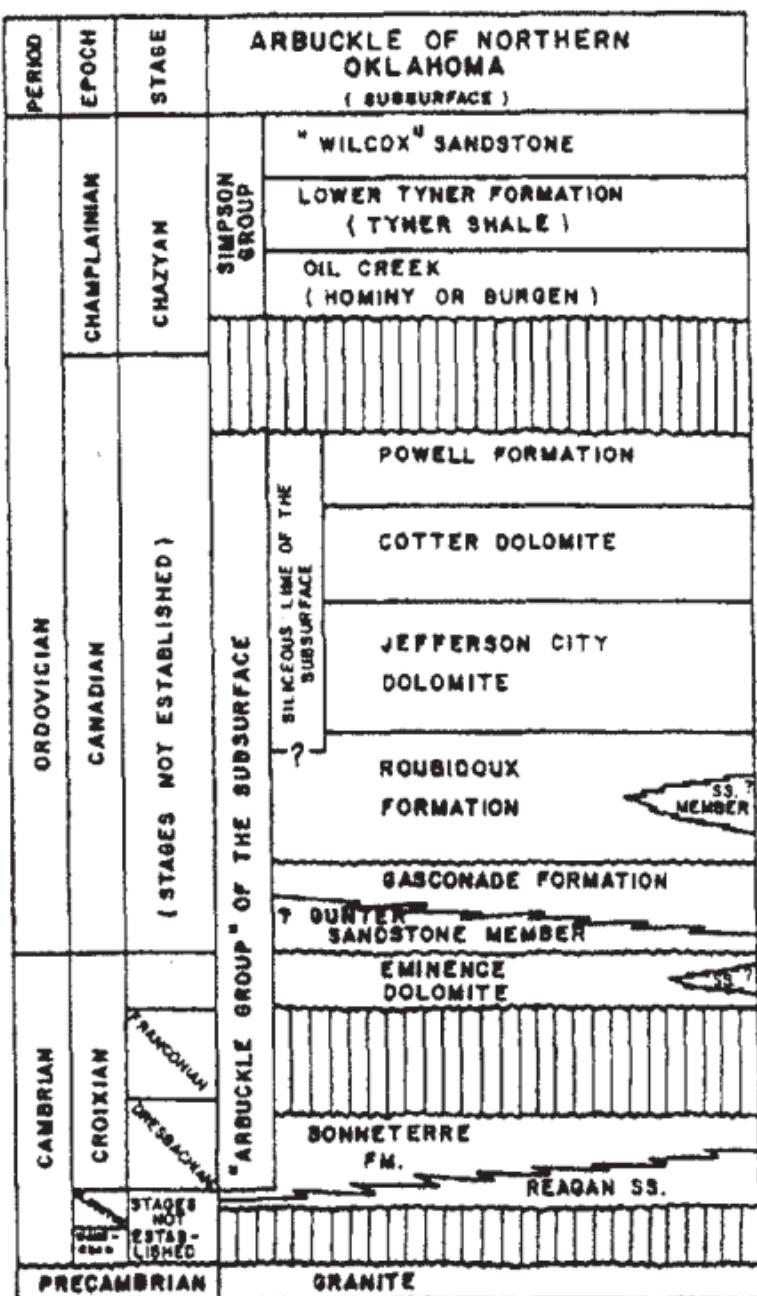


Fig. 31

Chenowith, P.A., 1968
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Geologists

23-25-8

19-25-9

22-27-8

Service Well

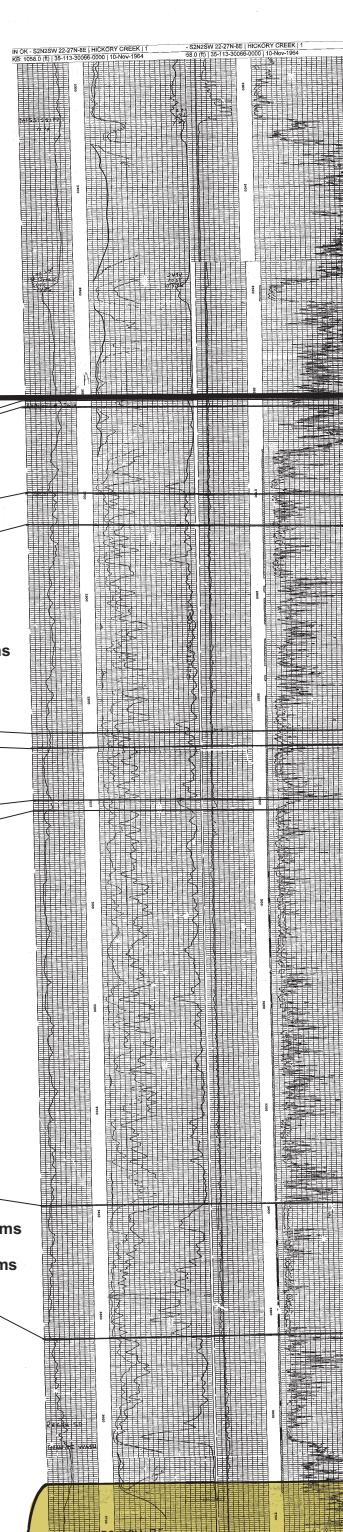
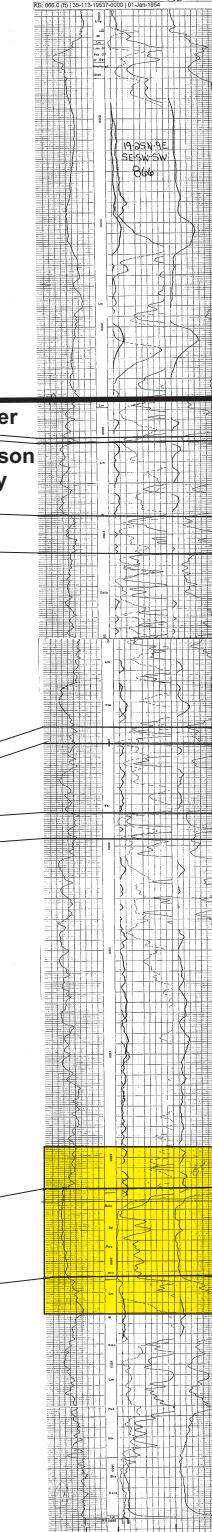
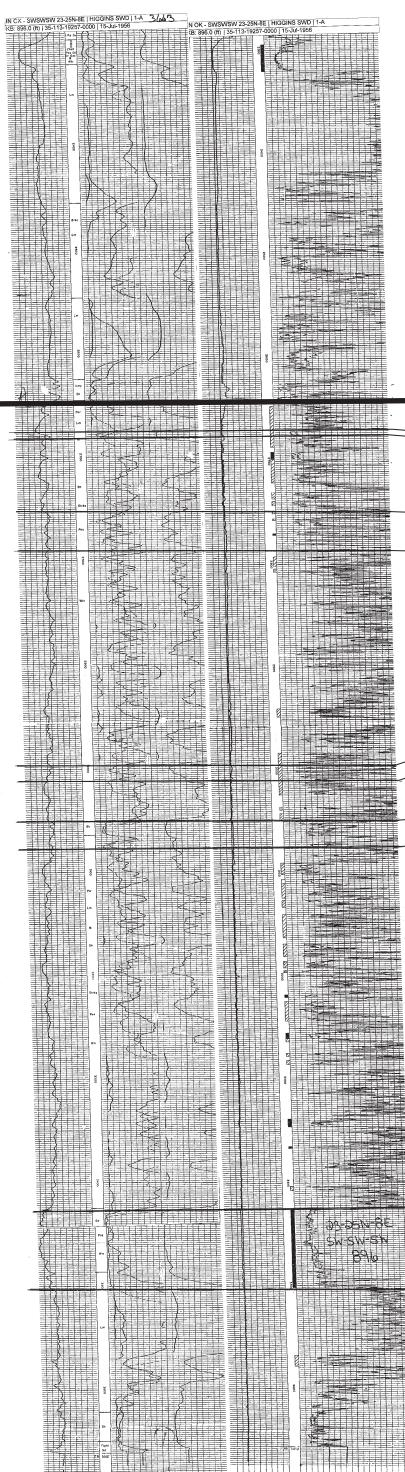
SWD

2 Miles

Service Well

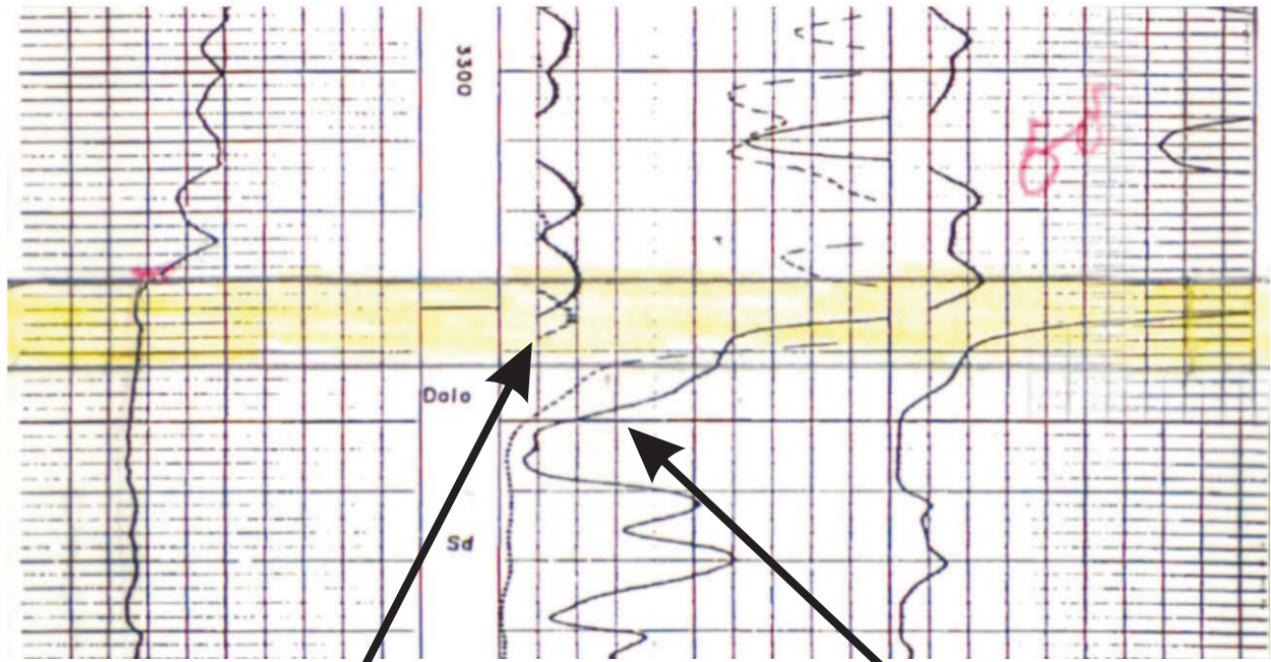
SWD

N 20 Miles



**22-27N-8E
S2-N2-SW
1058**

Fig 32



100 OHMS RESISTIVITY

POSSIBLE OIL/GAS / WATER CONTACT

Fig 33

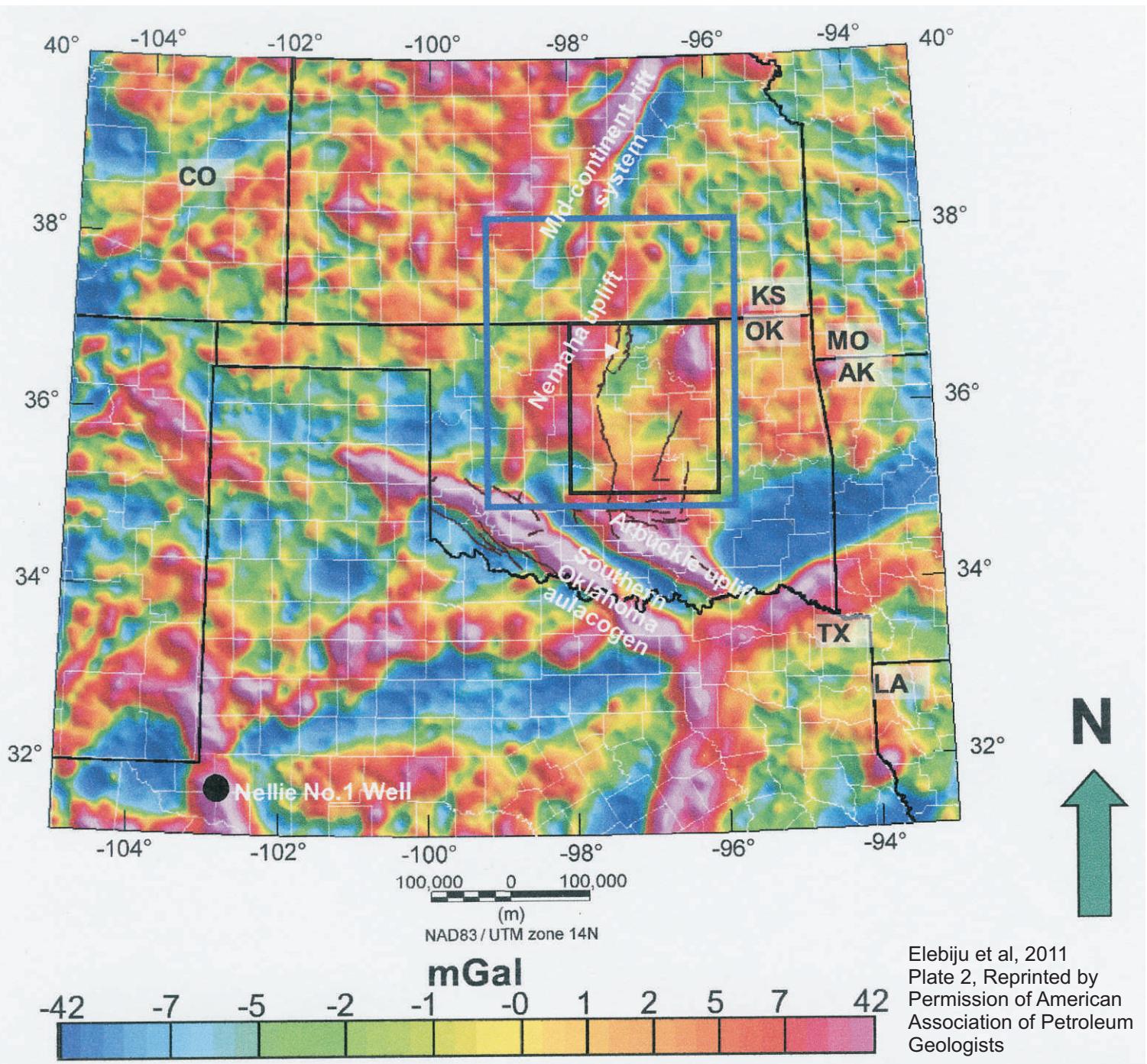
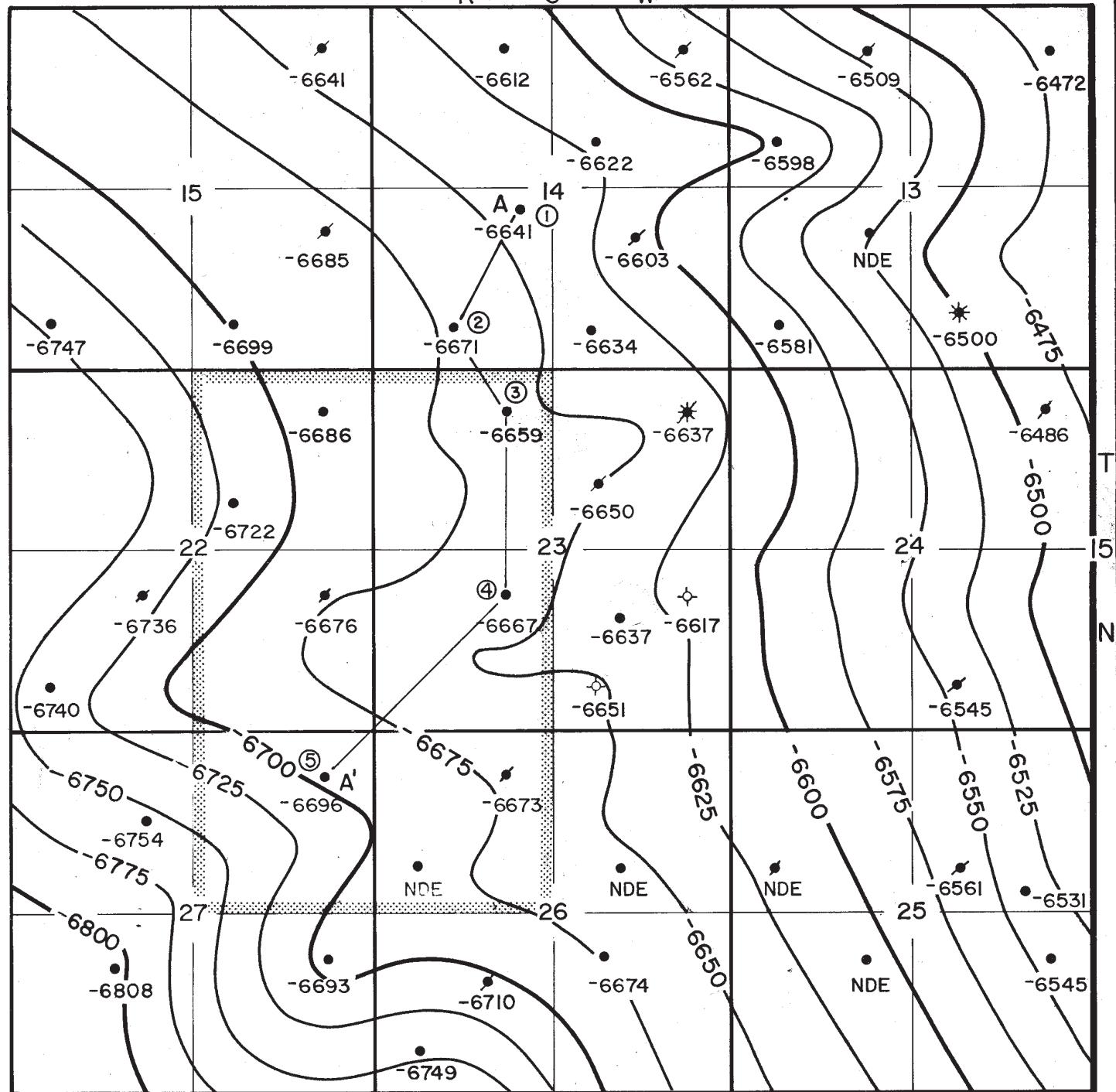


Fig 34



- L E G E N D -

NDE = NOT DEEP ENOUGH

AMI ■■■■■

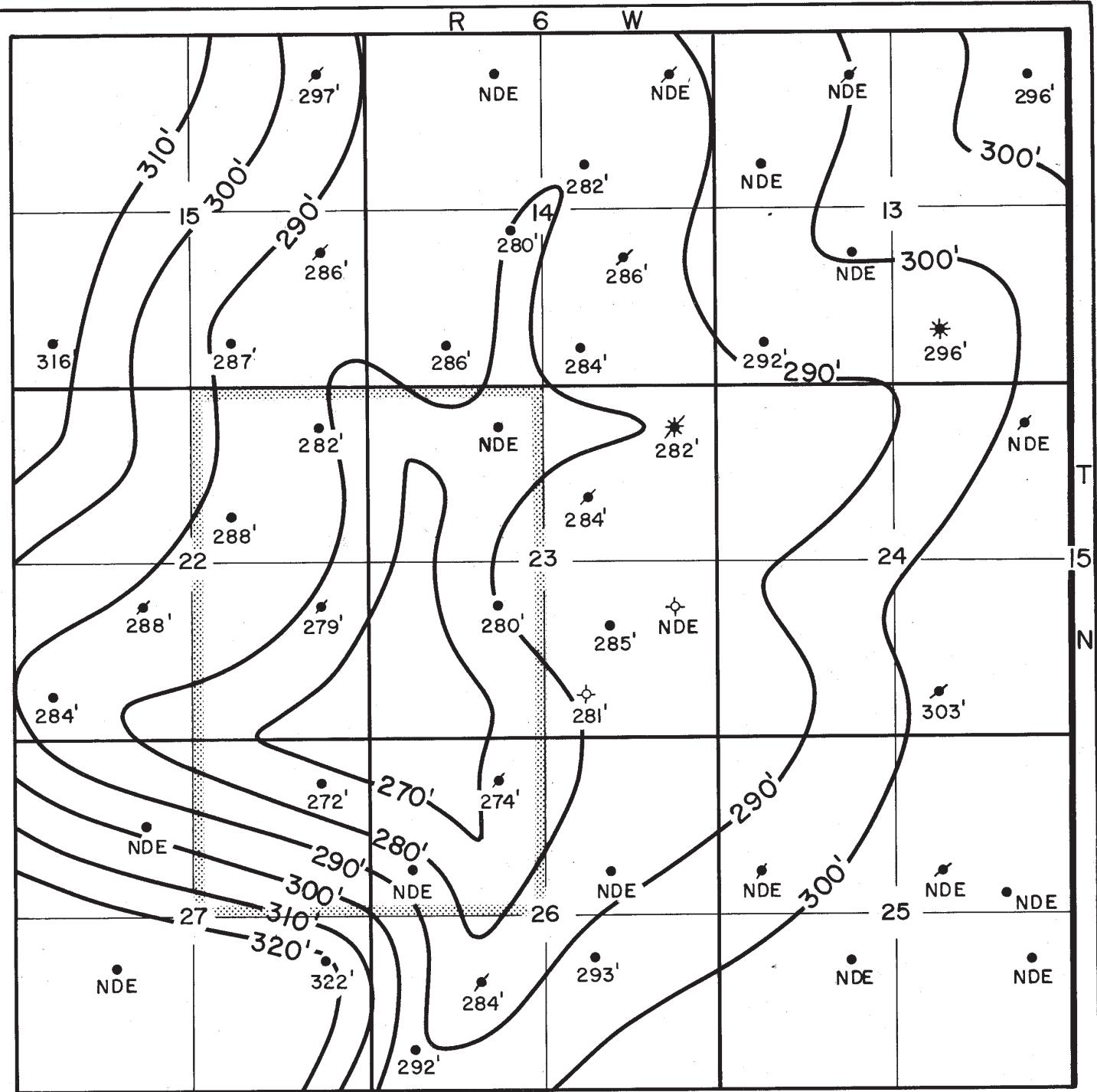
SUNFLOWER PROSPECT
KINGFISHER CO., OK.

HUNTON STRUCTURE
C.I. = 25'

GEOL: CONNIE ALLEN

SCALE: 1" = 2000'

Fig 35



- LEGEND -

NDE = NOT DEEP ENOUGH

AM 1 =

SUNFLOWER PROSPECT

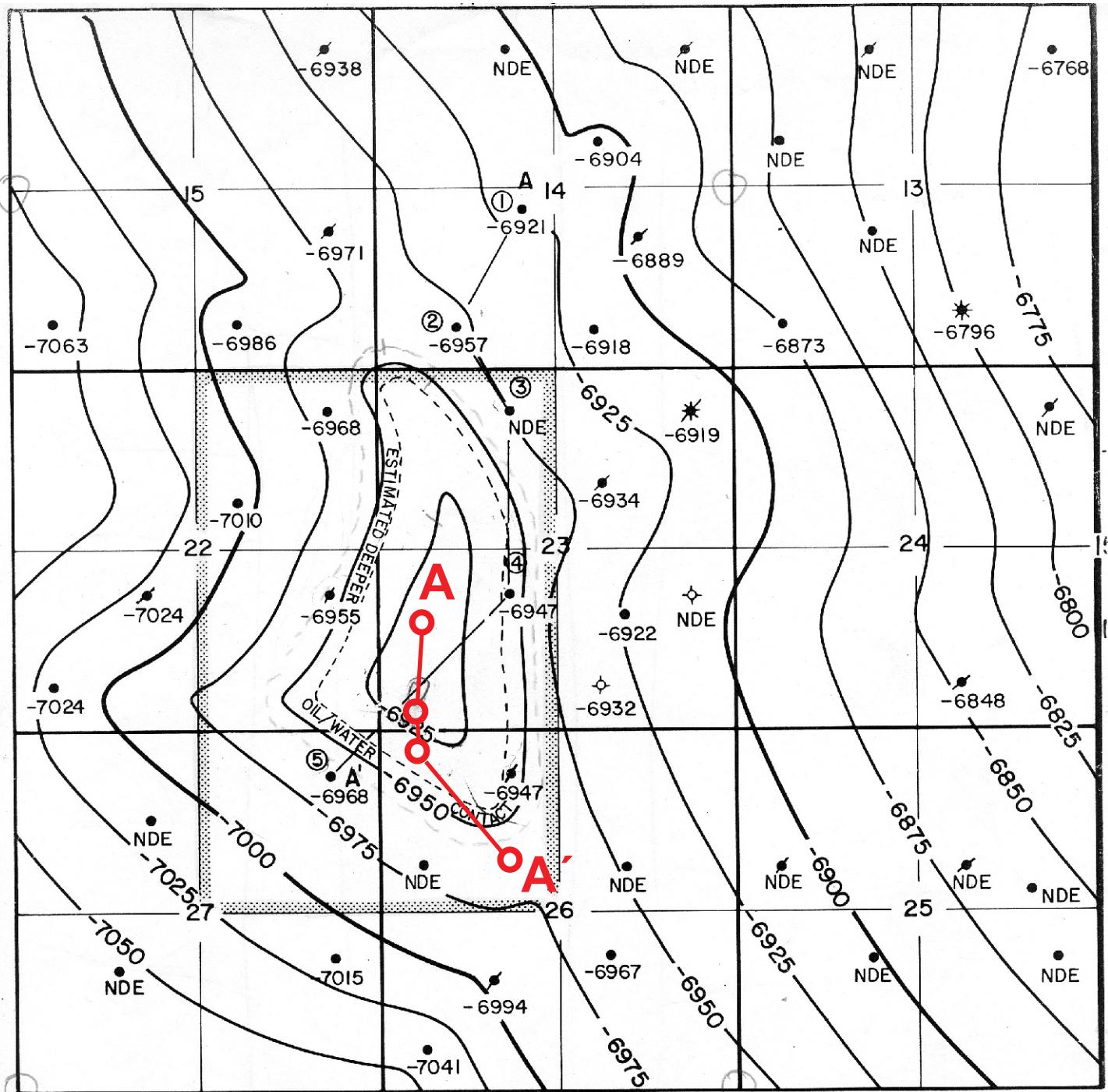
KINGFISHER CO., OK.

HUNTON ISOPACH
C.I. = 10'

GEOL: CONNIE ALLEN

SCALE: 1" = 2000'

Fig 36



- LEGEND -

NDE = NOT DEEP ENOUGH

SUNFLOWER PROSPECT

KINGFISHER CO., OK.

SYLVAN STRUCTURE

C.I. = 25'

GEOL: CONNIE ALLEN

SCALE 1" = 200

Fig 37

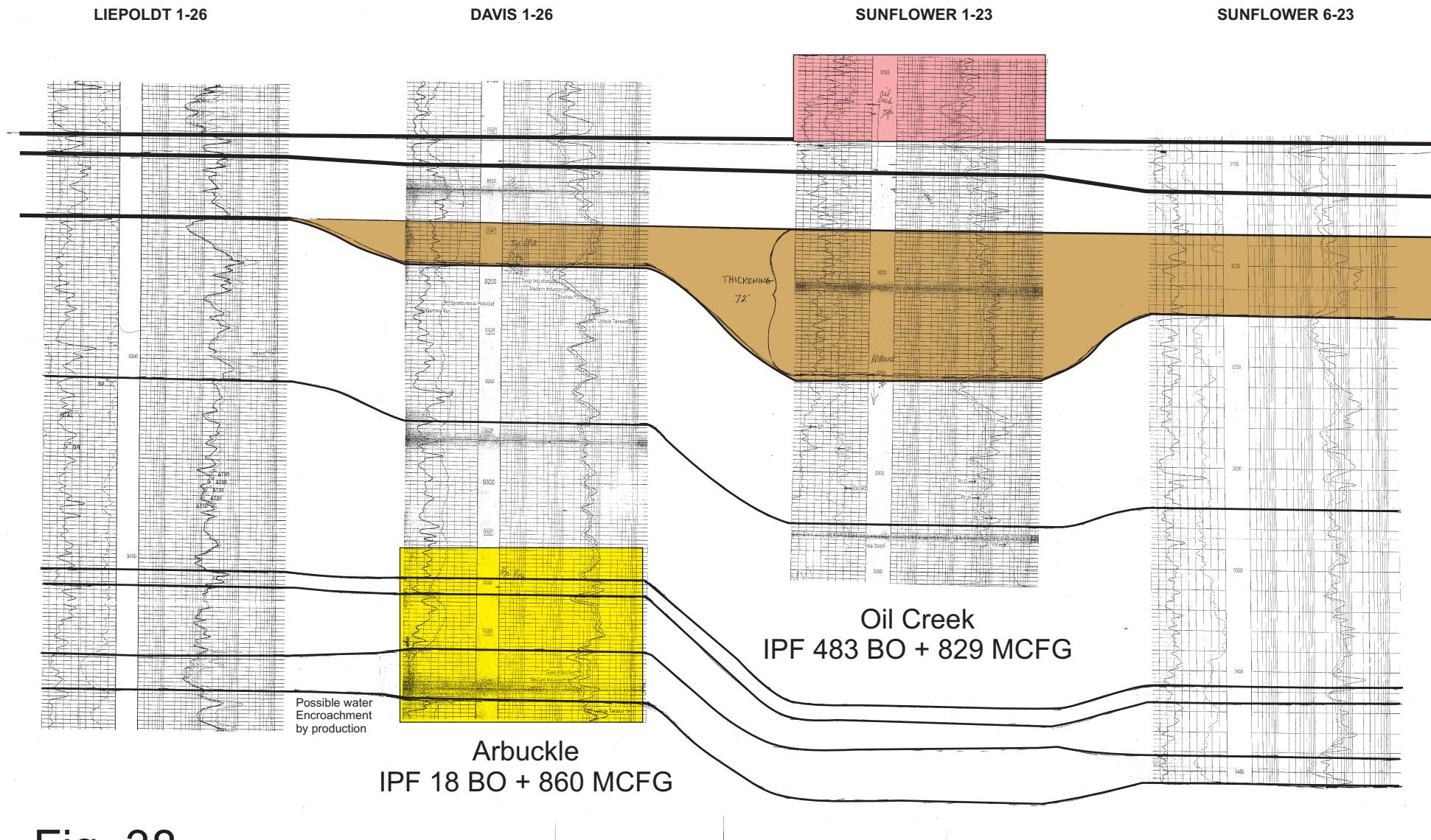


Fig. 38



Fig. 39 Pink Structure, Kingfisher County, OK

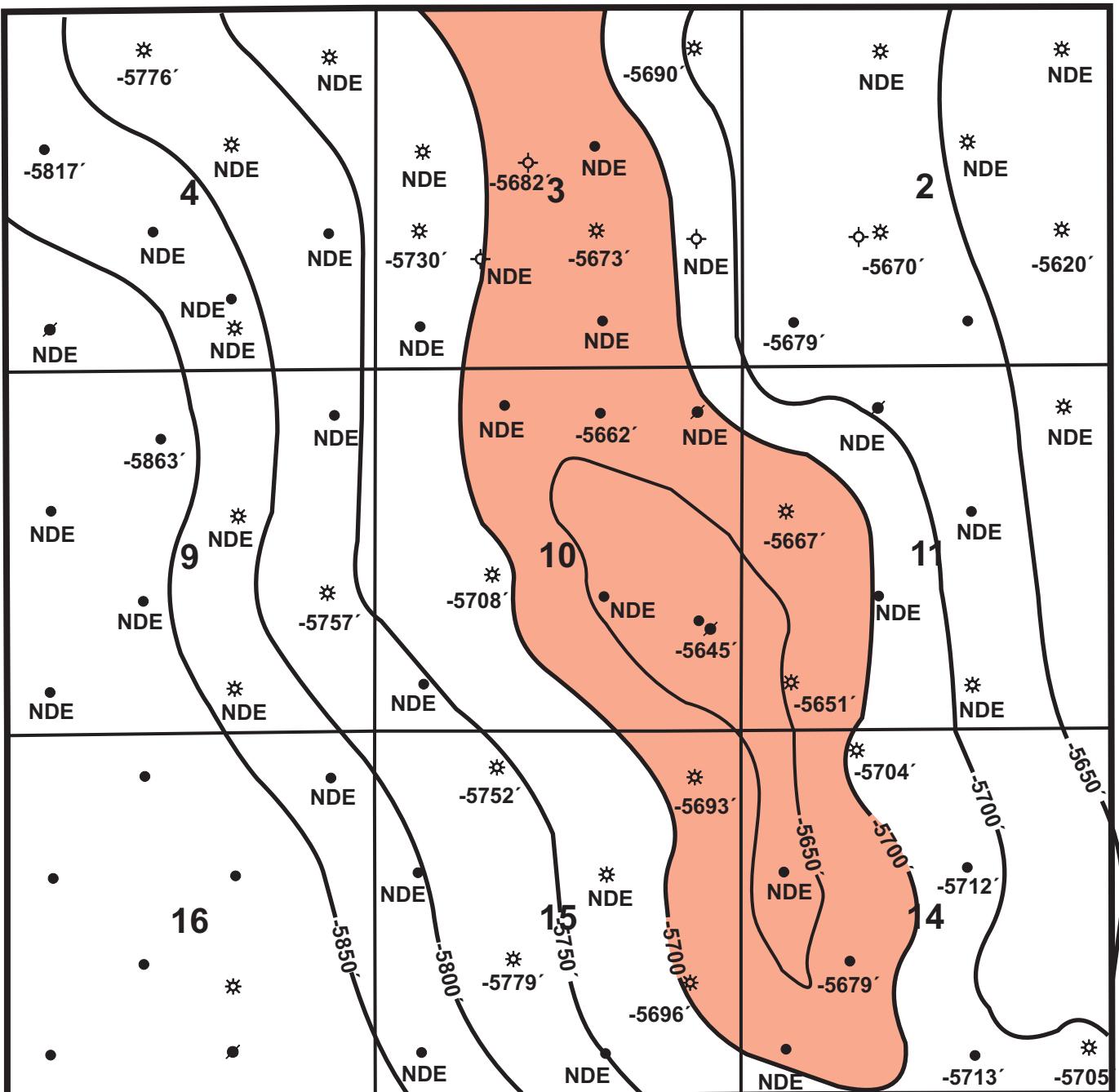


Fig 40 Sylvan Structure

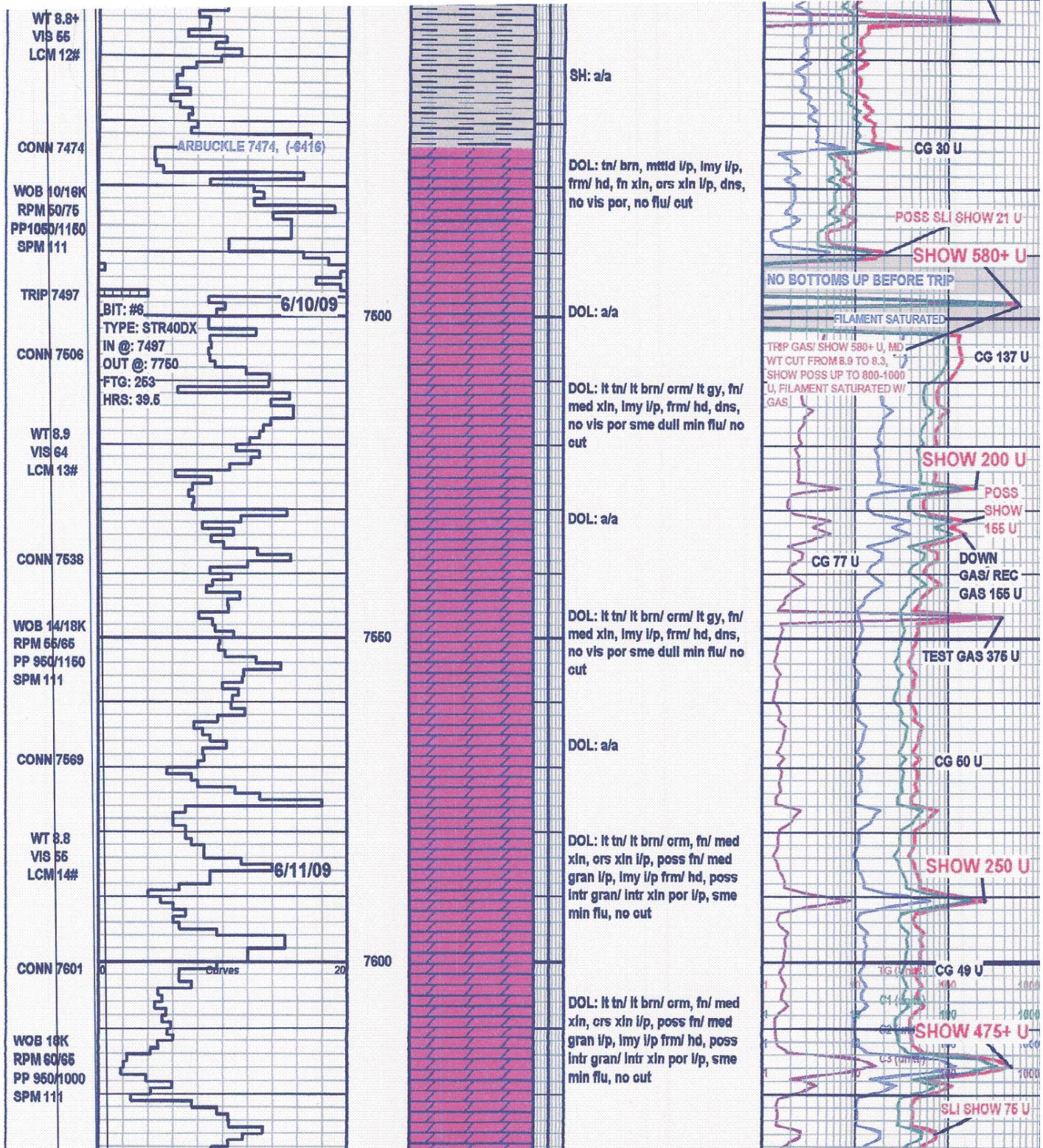


Fig 41

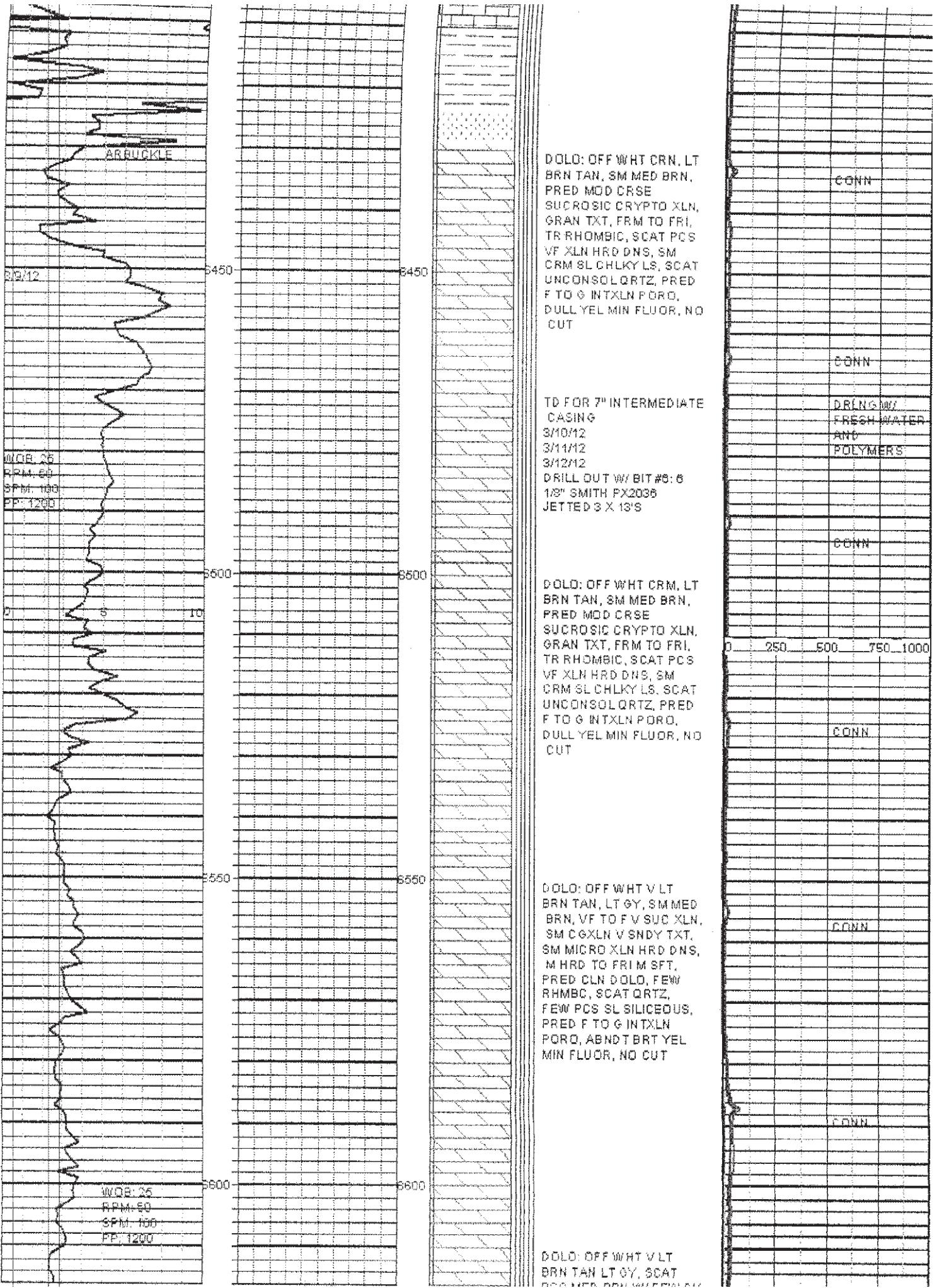
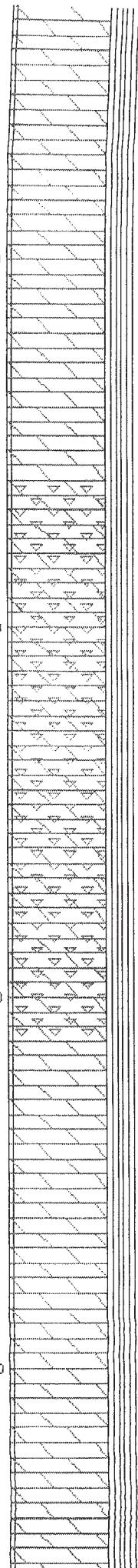
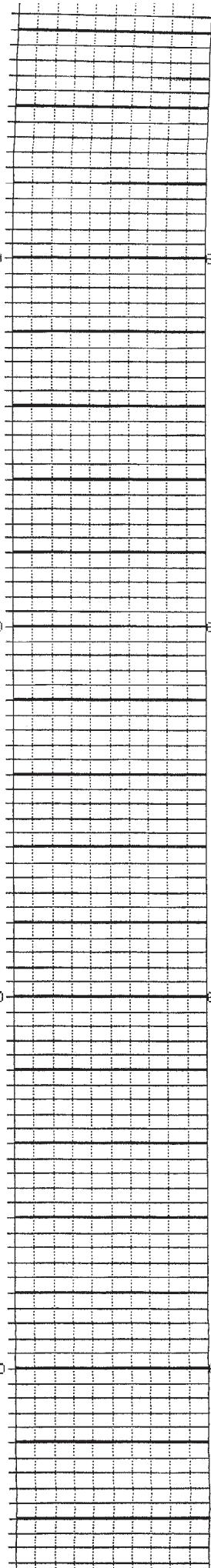
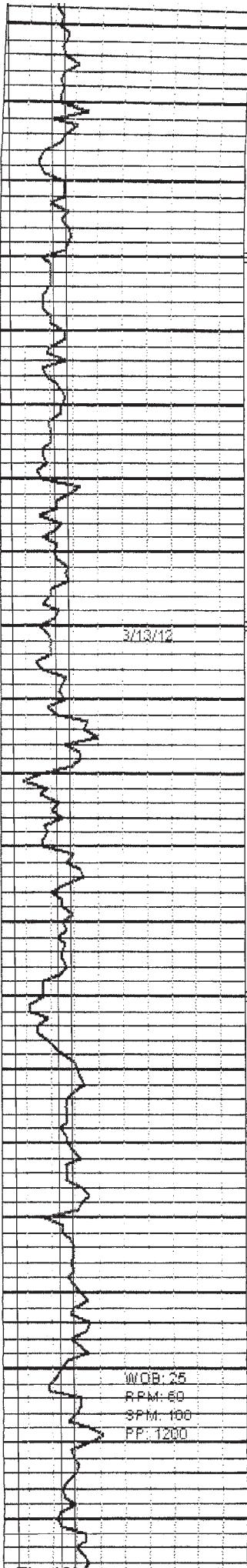


Fig 42a



FLO WED BRN WY FEW DN
BRN, PRED VF SL TO
MOD SUD CGXLN, SCAT
PCS MICRO XLN V HRD
DNS CLN DOL0, SM PCS
SL SIL, FEW SEC FRACS.
F INTXLN PORO, ABNDT
BRT YEL ORNG MIN
FLUOR, NO CUT

CONN

CEINN

GOIN

DOLD: OFF WHT V LT TAN
V LT GY V SL TRNSL,
SMKY GY TAN OPAQ, F
TO MOD CRSE V SUC
CRYPTO XLN, V AREN
SNDY TXT, FRITO FRM,
FEW RHOMBIC, SCAT
PCS V SIL IN MTRX,
ABNDT SMKY GY BLU TO
BRN TNT TRPLTC
CHERT, SCAT PCS W/G
VIS SEC FRACS, F TO G
IN TXLN PORO, ABNDT
BRT MIN FLUOR, NO CUT,
SCAT FREE QRTZ

conn

**INCREASE B.G.
500-600 u**

CONN

CONN

DODO: LT BRN TO MED
BRN, SM DKR BRN,
MICRO TO VF XLN, SL
SUC IP, M HRD, SL SIL IN
MTRX, PRED CLN HRD
DODO, PRED P INTXLN
PORO, SM SEC FRACTS,
DULL MIN FLUOR, NO
CUT

conn

Fig 42b

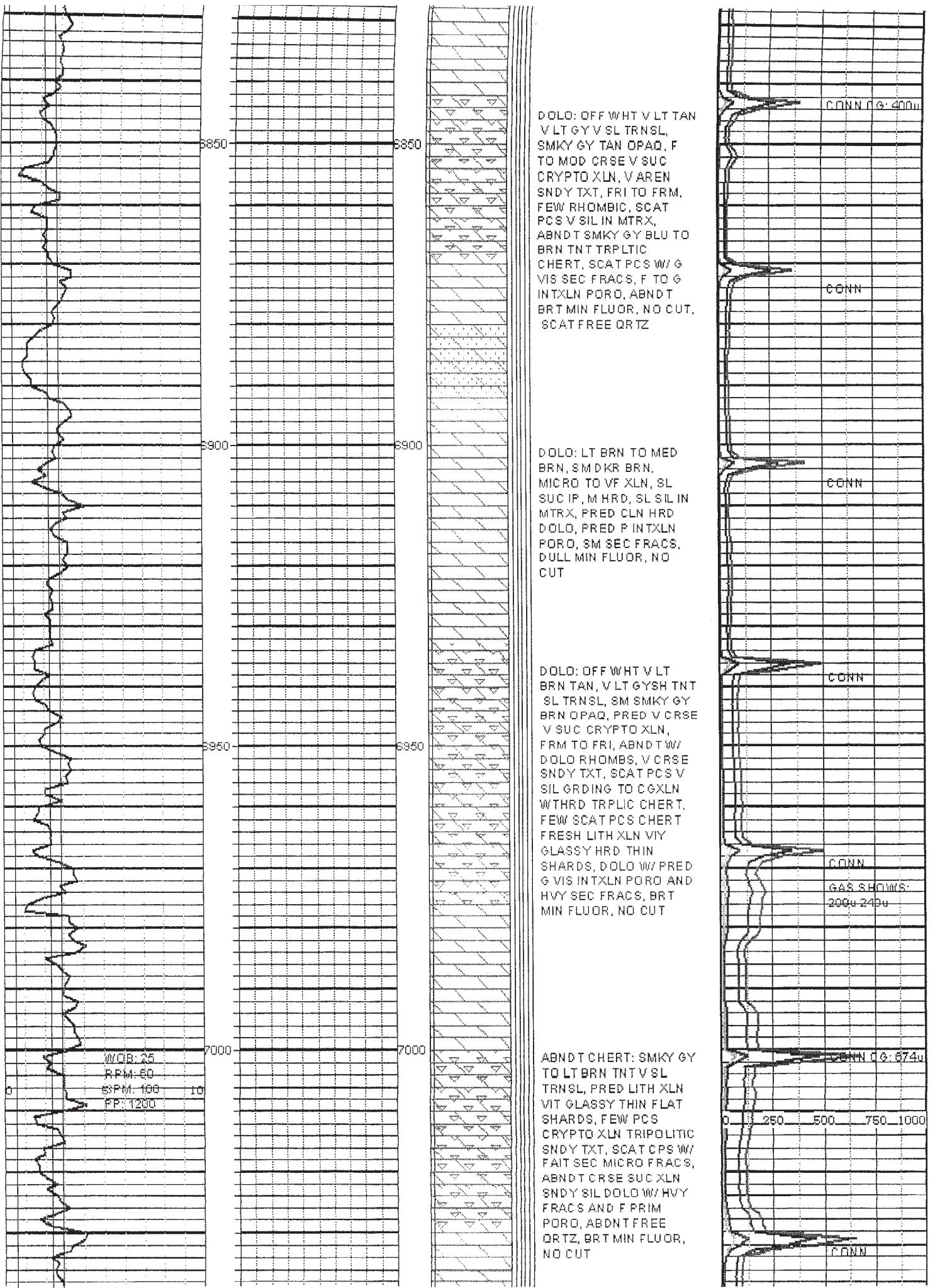


Fig 42c

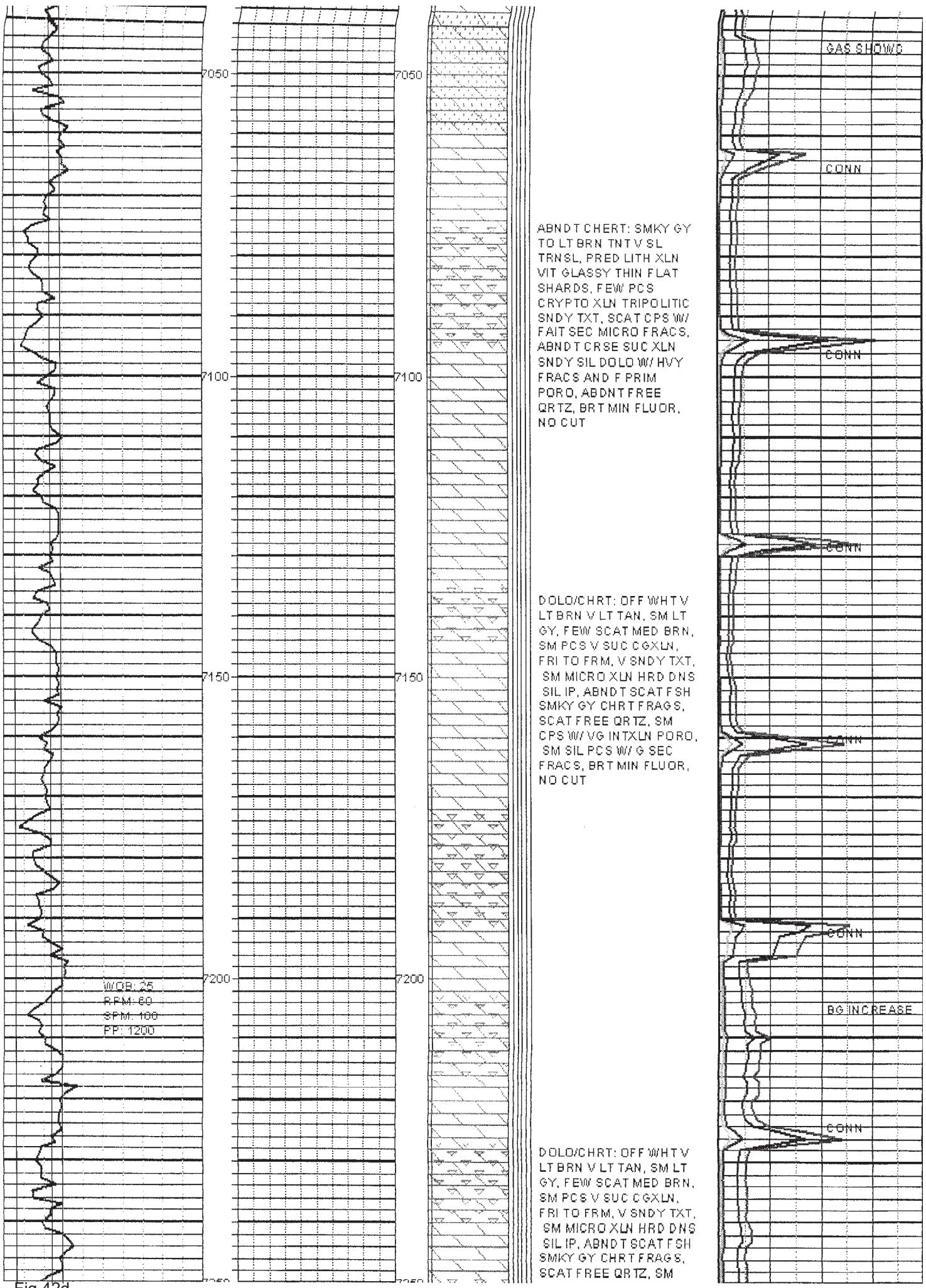


Fig 42d

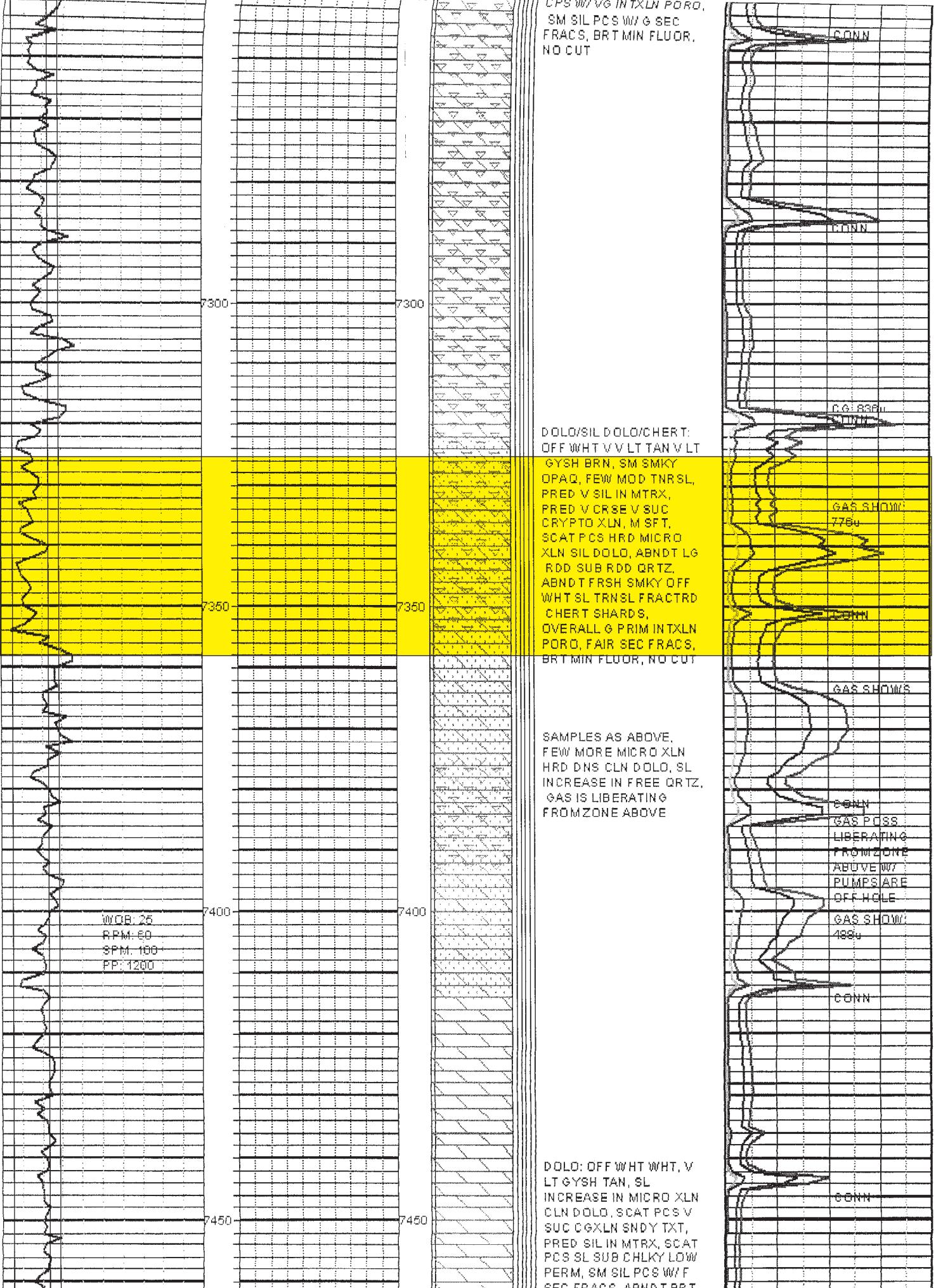
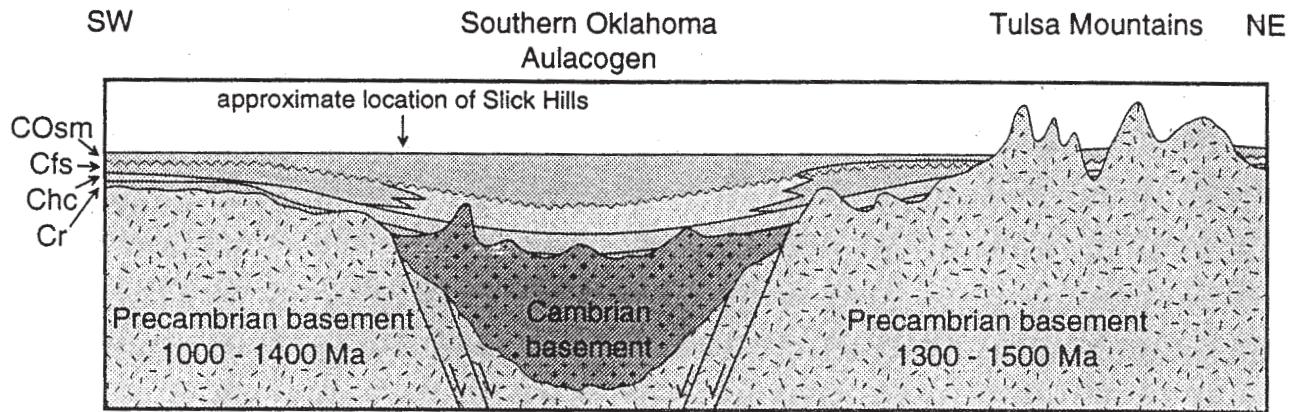
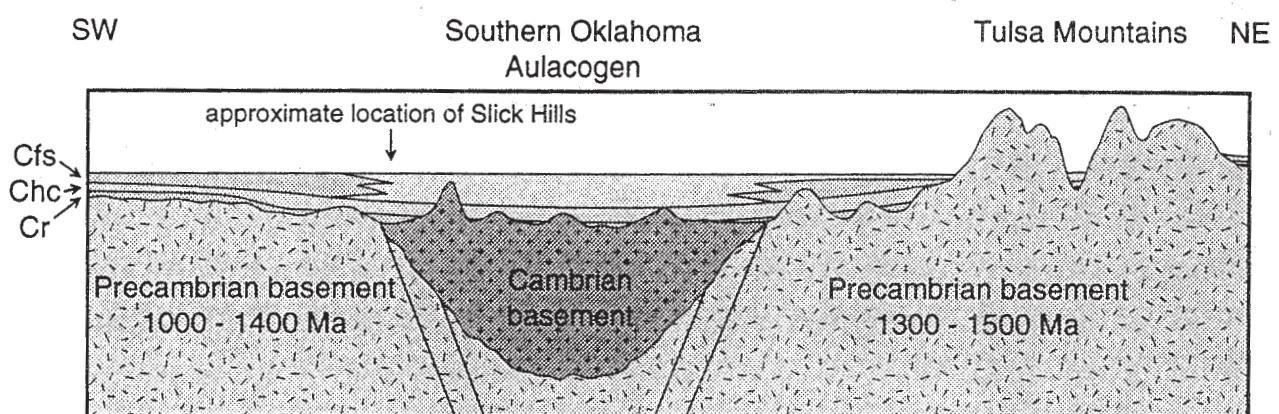


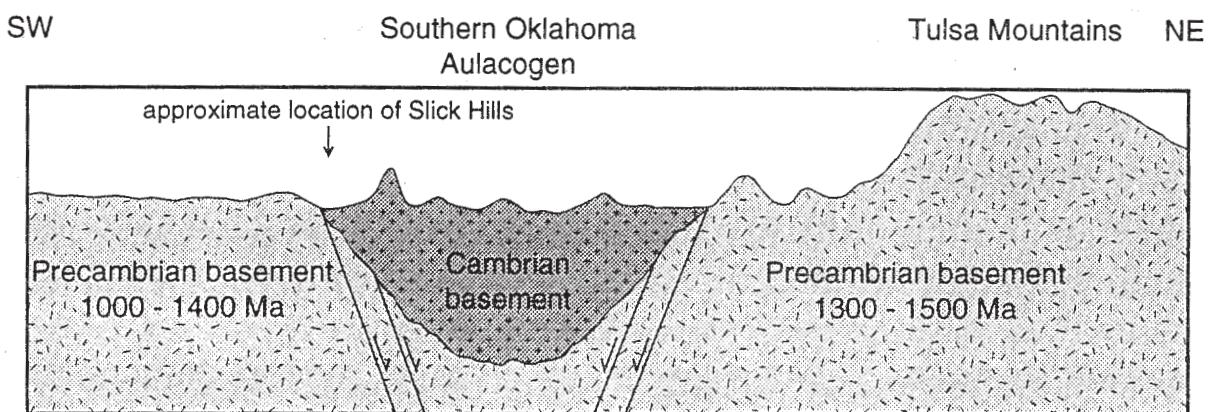
Fig 42e Excellent Gas Show base of Arbuckle



STAGE III - Reactivation of basin margin faults and deposition of Cambro-Ordovician Signal Mountain Formation in deeper waters? (Trempealeauan - Canadian)



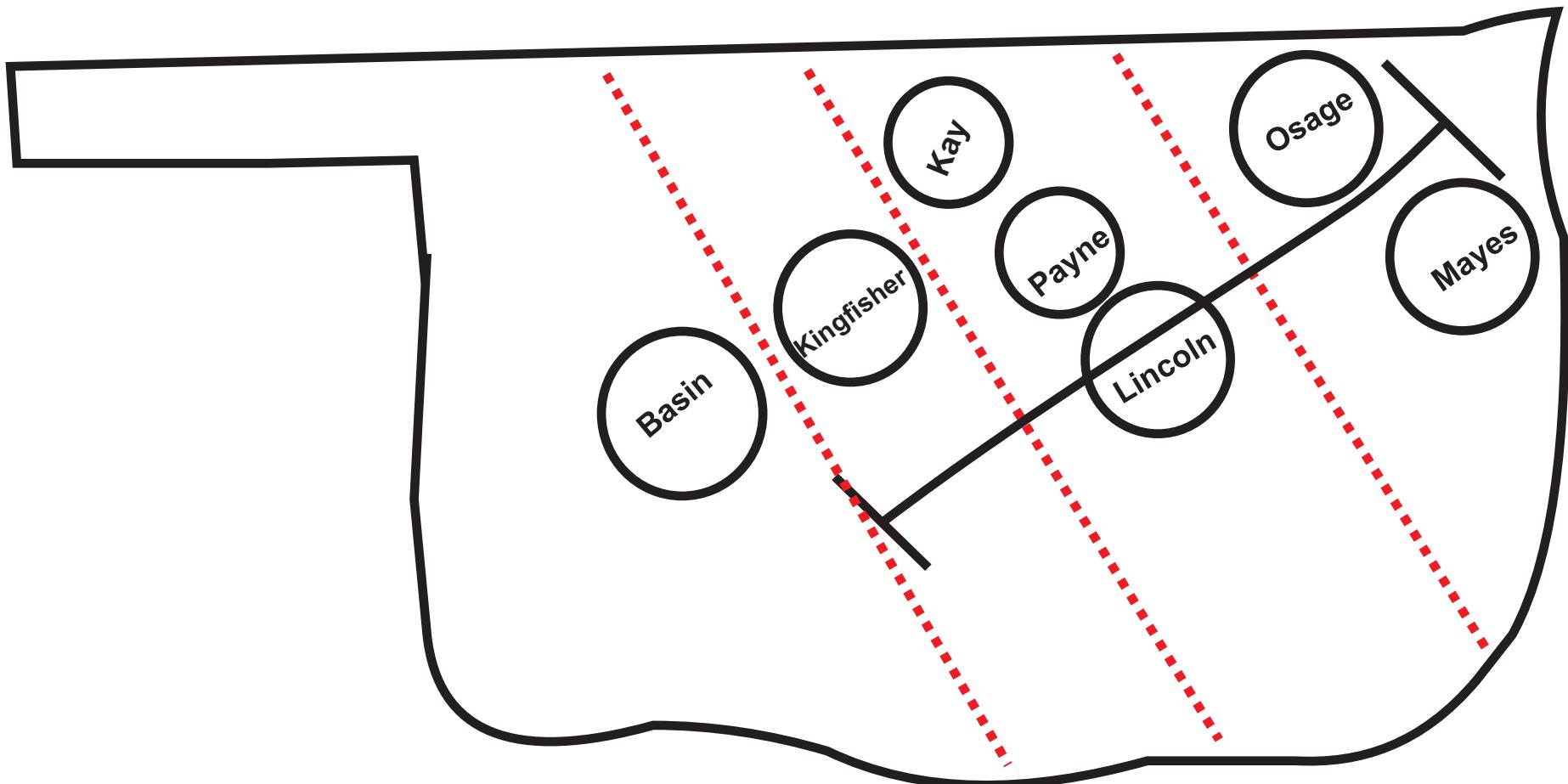
STAGE II - Deposition of Cambrian Timbered Hills Group and Cambrian Fort Sill Formation. (Franconian)



STAGE I - Development of rift basin. (Middle Cambrian)

Hosey and Donovan, 2000
Fig 13, Reprinted by
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Fig 43



AREAS EXHIBITING ARBUCKLE THICKS AND THINS

Fig 44

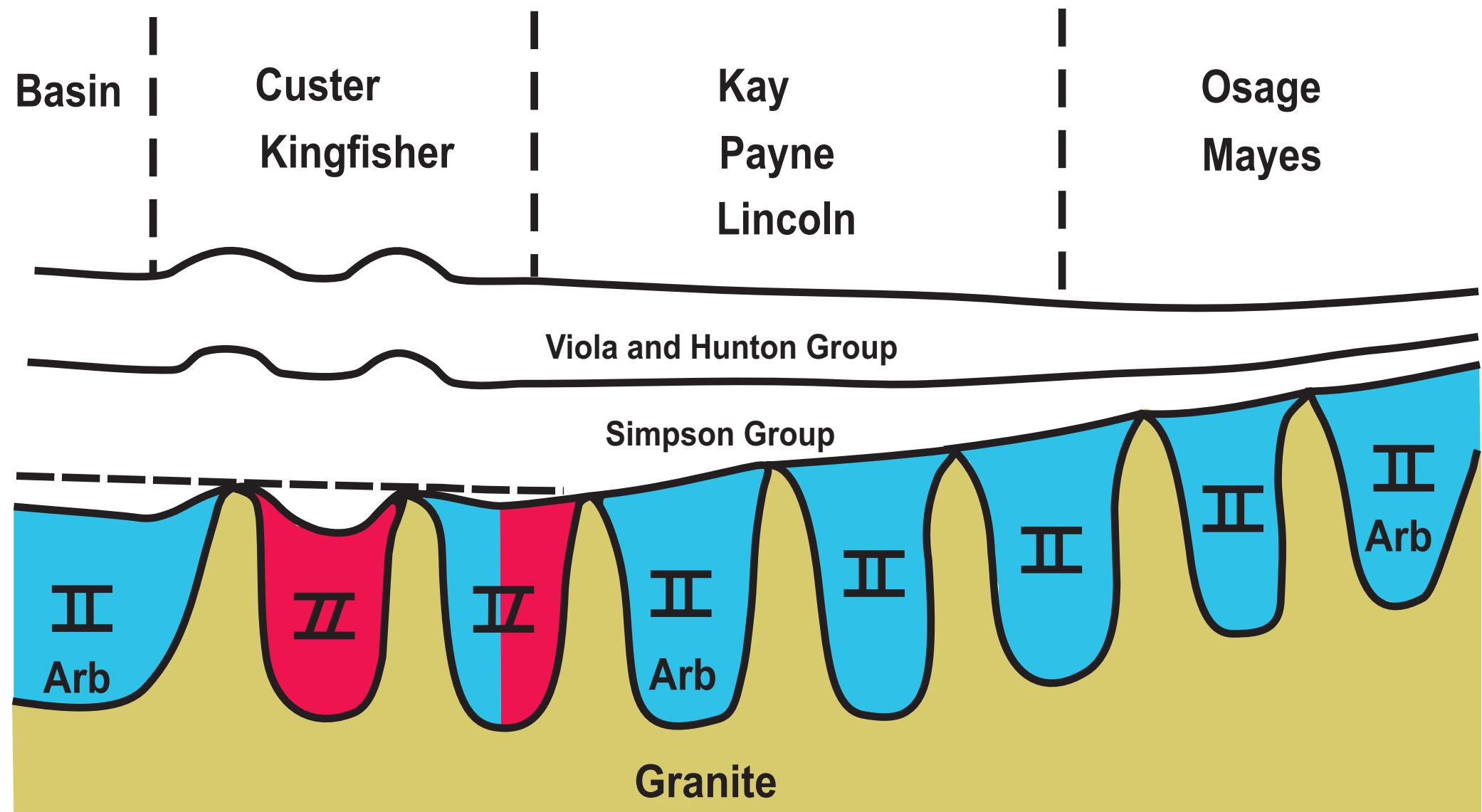


Fig 45

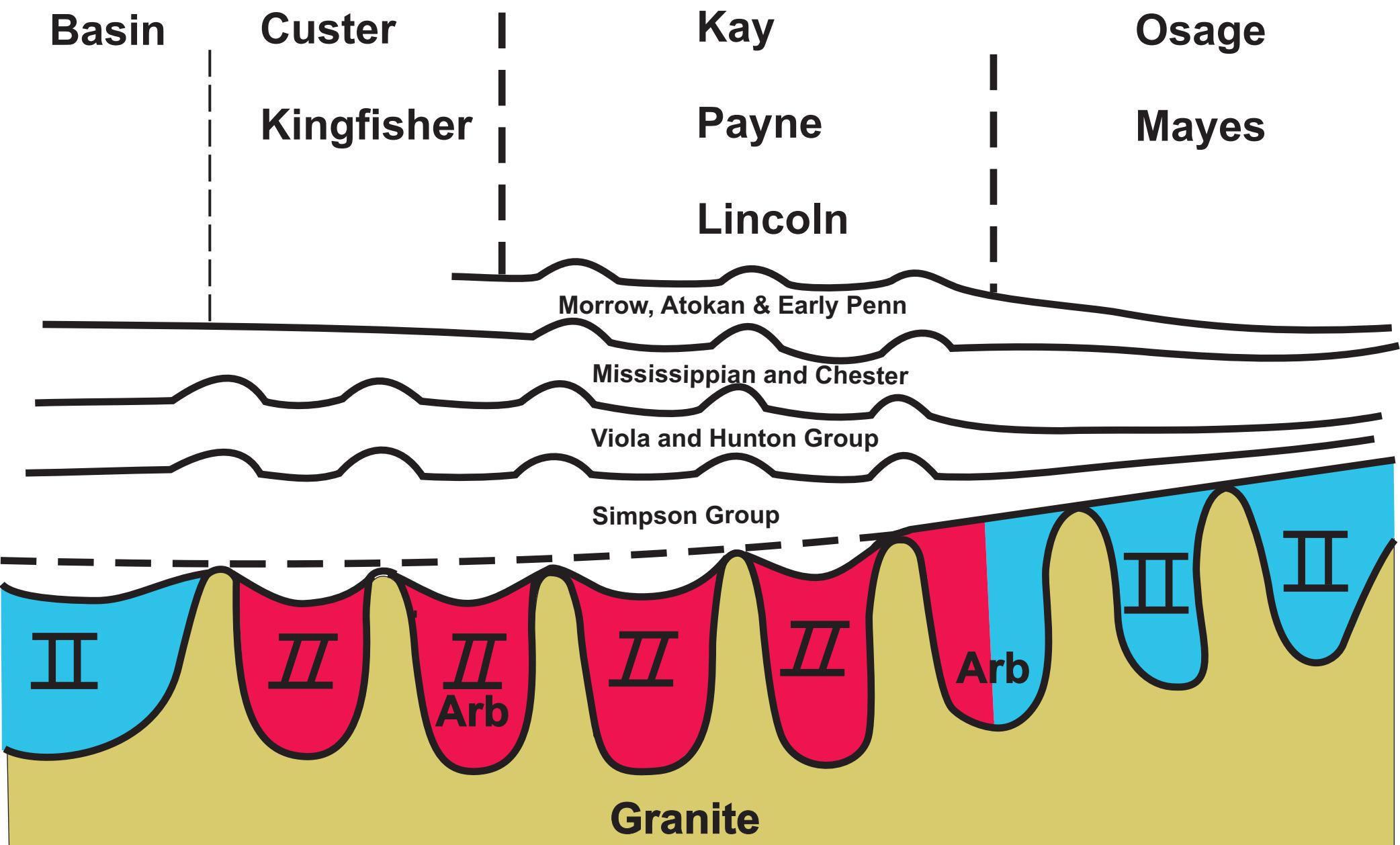


Fig 46

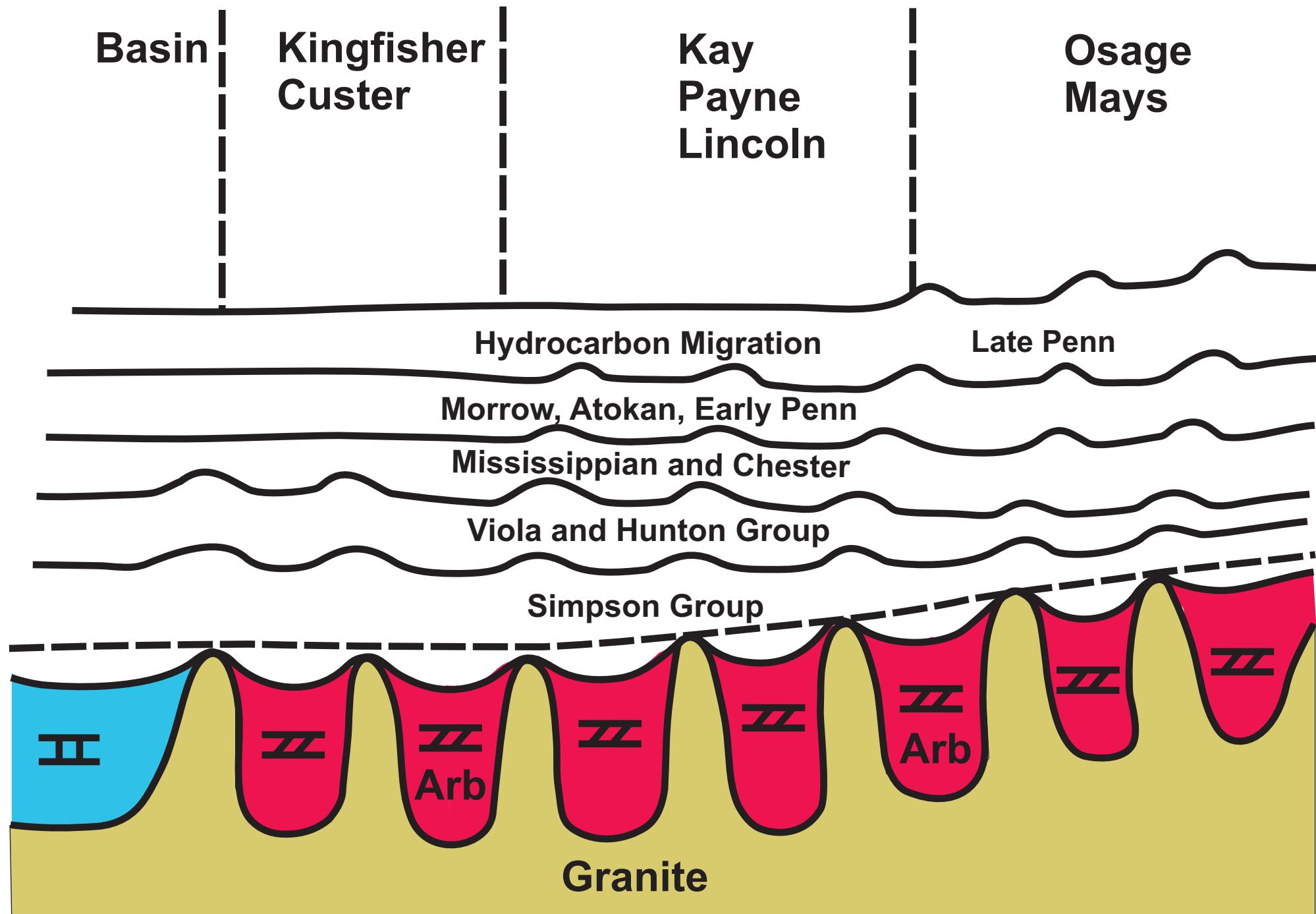


Fig 47

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